
The Arctic Subpolar gyre sTate Estimate (ASTE)

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Kirstin Schulz & CRIOS group**
(University of Texas, Austin)



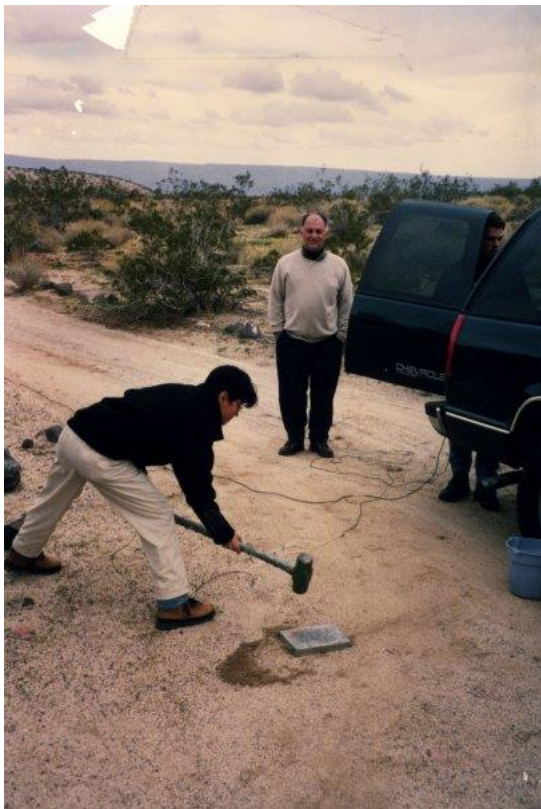
Estimating the Circulation and Climate of the Ocean
Summer School 29 May 2025



The University of Texas at Austin
**Oden Institute for Computational
Engineering and Sciences**



Who am I



From Vietnam

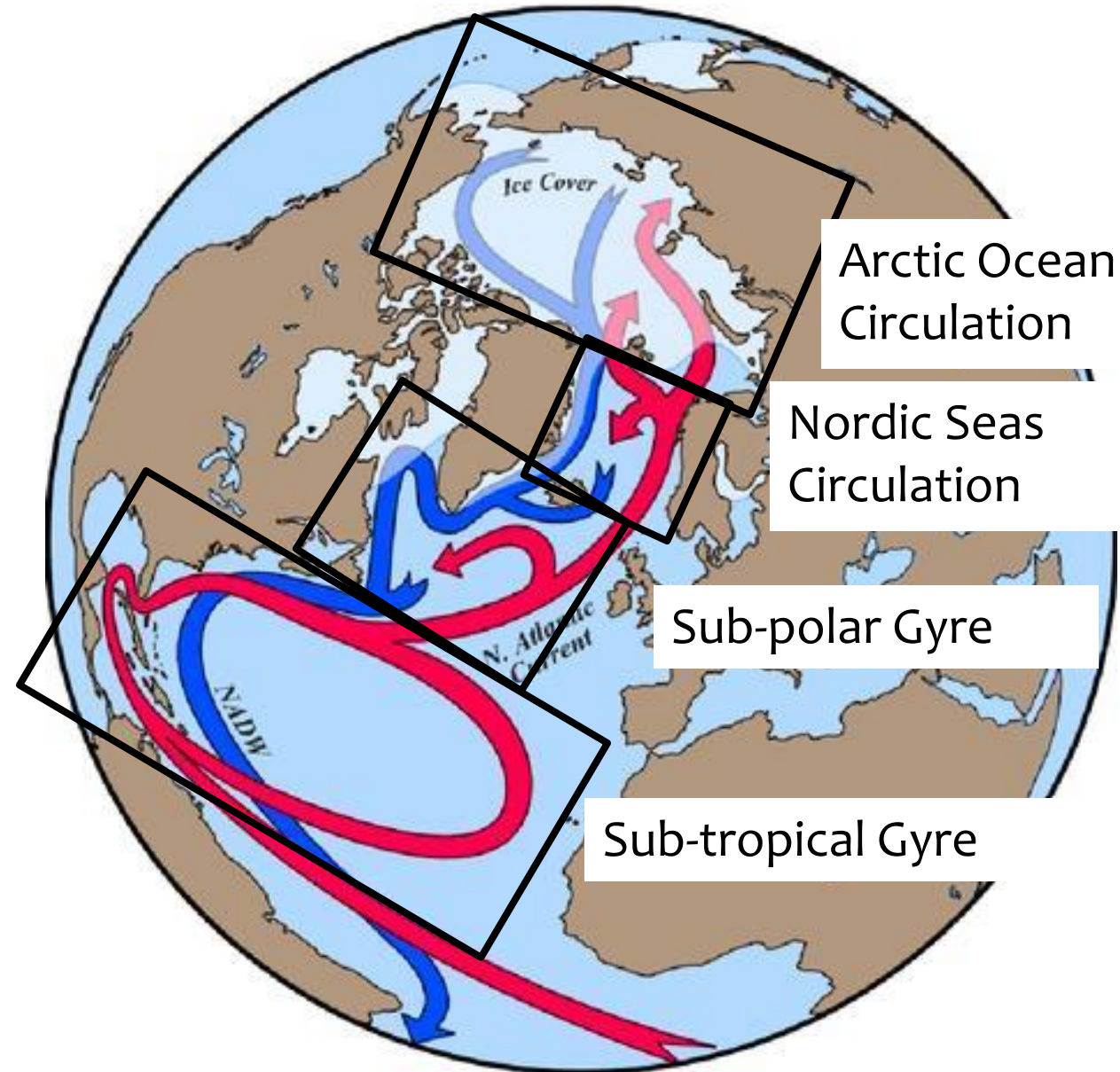
Undergraduate: UCLA , applied Geophysics

Ph.D. : MIT, Geophysics: satellite altimetry to detect change in Ice Sheet

Postdoc & the rest: ECCO



North Atlantic Meridional Overturning Circulation (MOC)



High latitude ocean circulation

Subpolar Gyre + Nordic Seas

- Circum-Greenland current
- Warm Atlantic Water
→ Greenland Fjords
→ into Arctic Ocean
- Dense water in Labrador and Nordic Seas
- Freshwater from Greenland Ice Sheet



[Straneo & Heimbach, 2013]

Arctic Ocean circulation

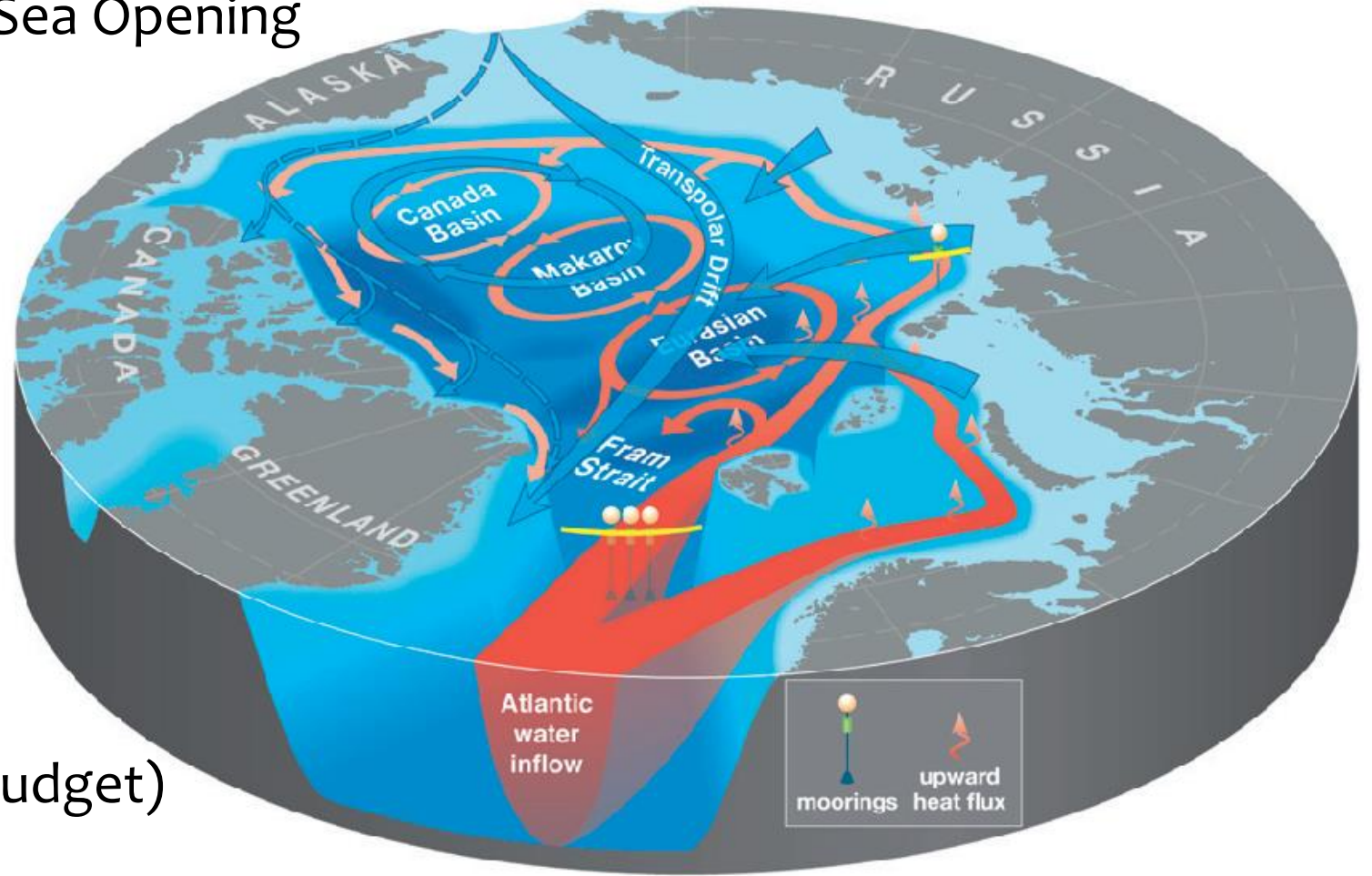
Atlantic Water (deep)

[Woodgate, 2012; Polyakov et al., 2013]

- Enters at Fram Strait + Barents Sea Opening
- subducts to mid-depth
- follow topography
- potential heat source to melt surface sea-ice

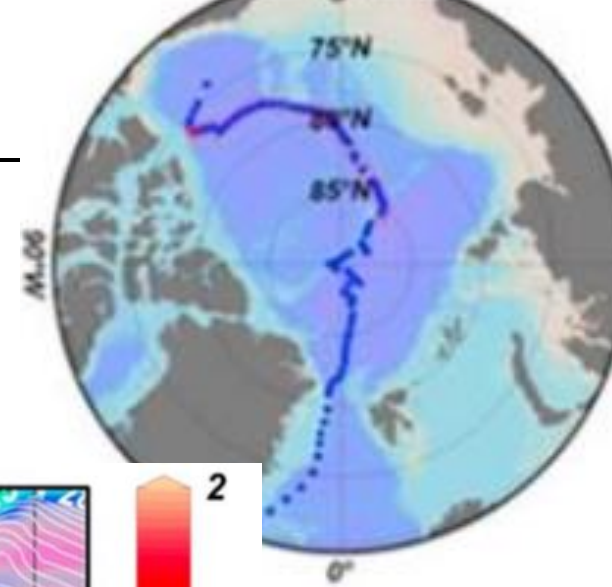
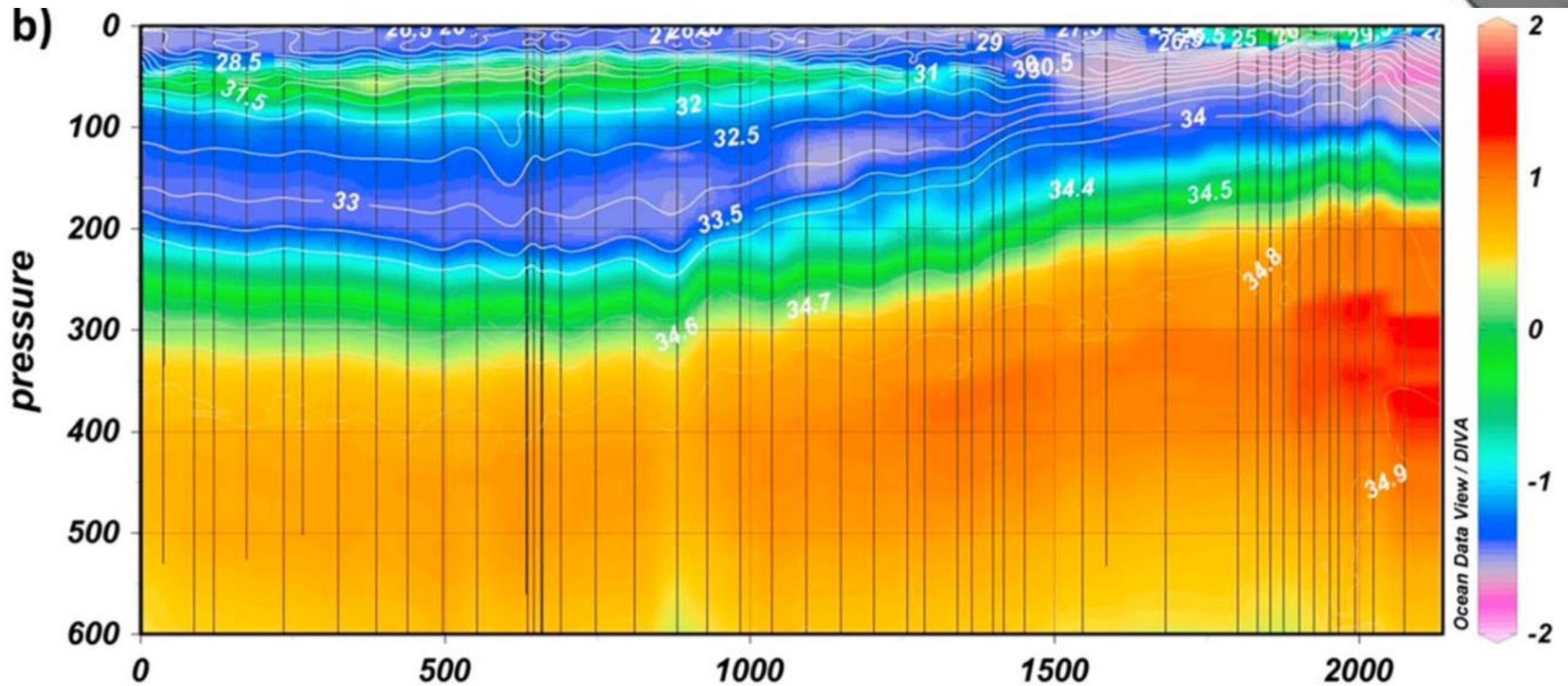
Pacific Water (mid depth & surface)

- Enters at Bering Strait
- Fresh water (33% of Arctic FW budget)
- Heat: near surface

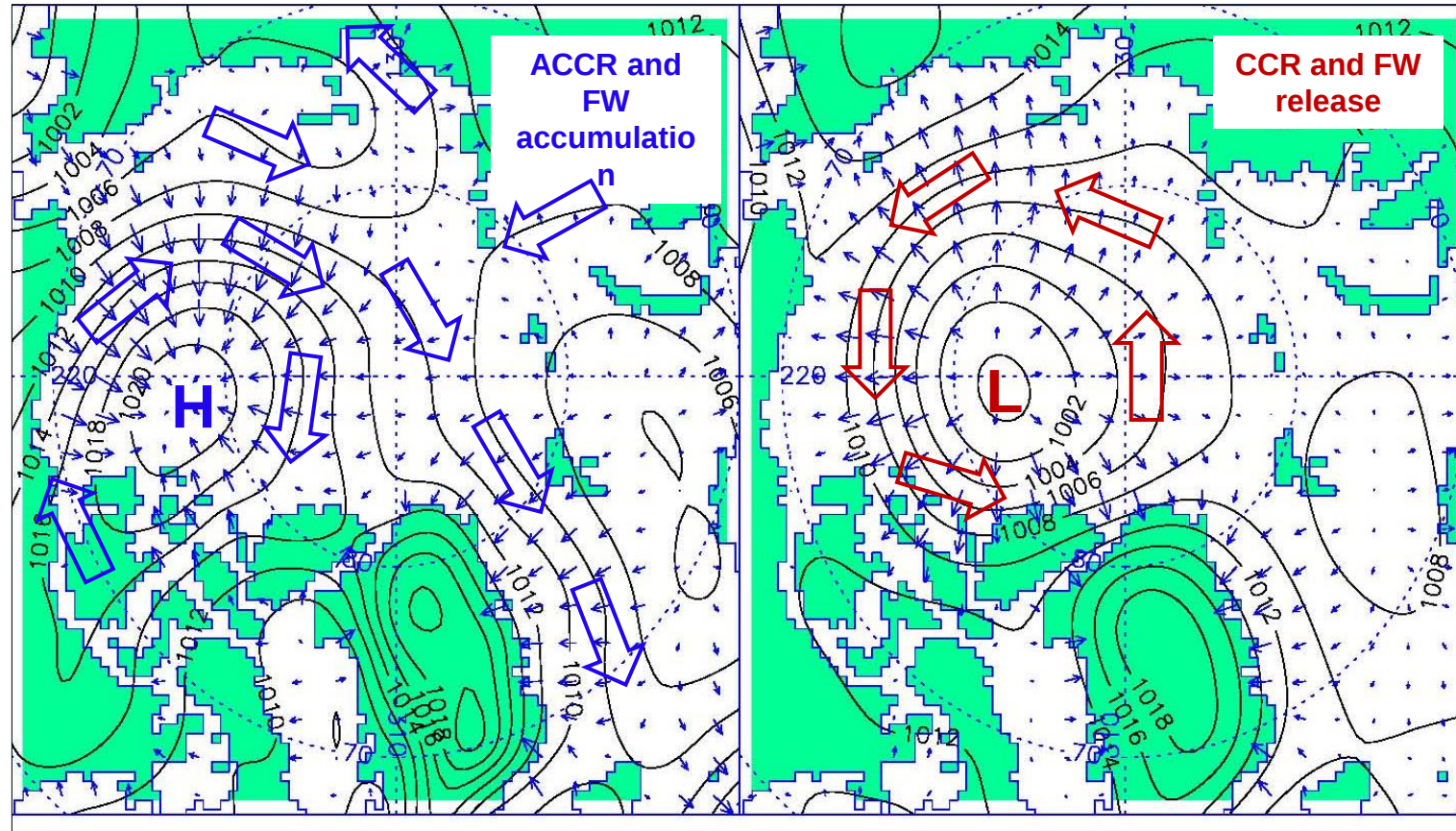


Arctic Ocean temperature structure

[Carmack et al., 2017]



Arctic Ocean: surface currents – wind driven



Panels show SLP (black lines, hPa) wind directions (large arrows) and Ekman transport (blue small arrows) typical for ACCRs (left) with Ekman transport converging; and CCRs (right) with Ekman transport diverging.

Challenges of (forward) modeling the Arctic

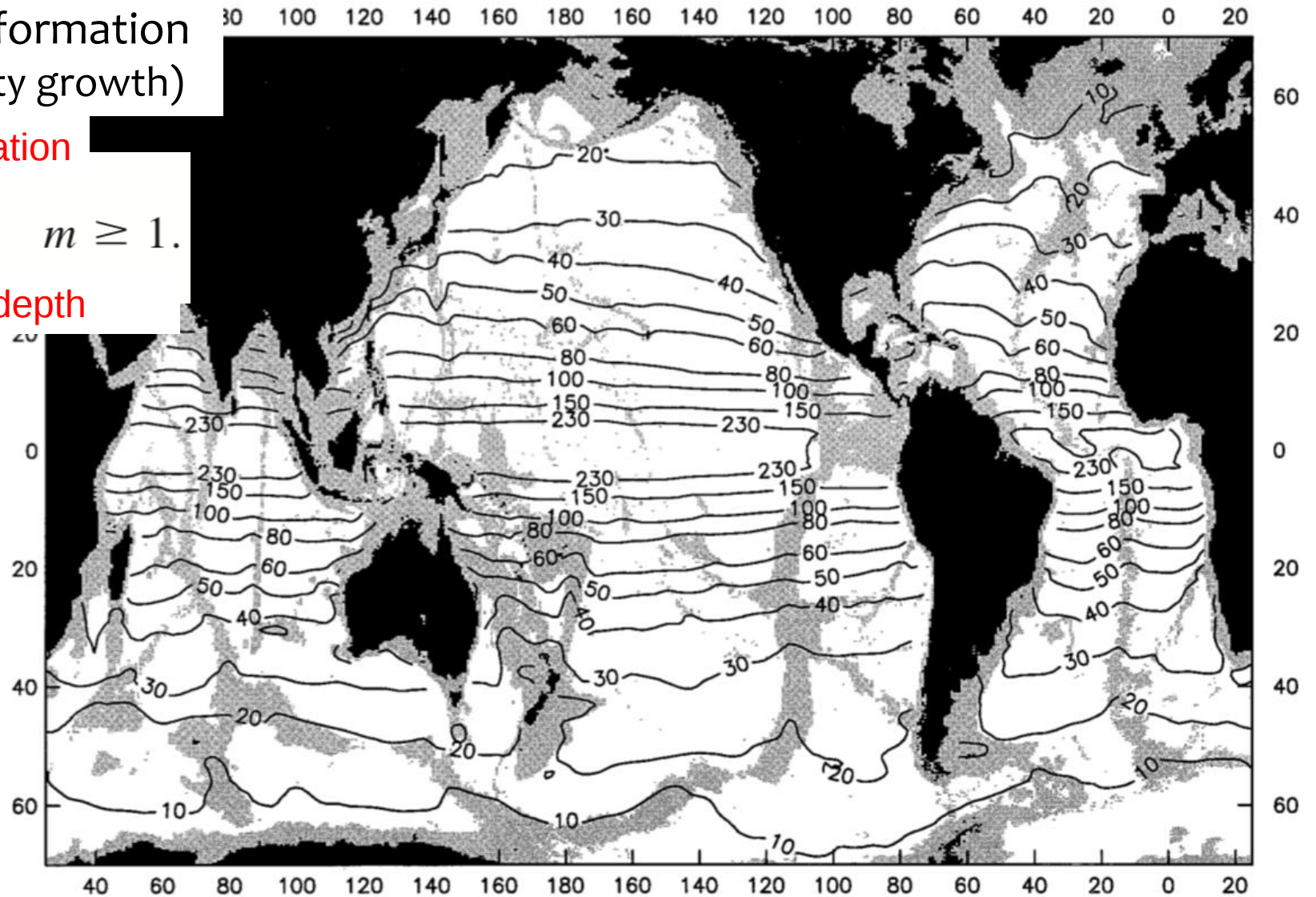
Small baroclinic radius of deformation
(lengthscale of fastest instability growth)

$$\lambda_m^{\text{WKB}} = \frac{1}{|f|m\pi} \int_{-H}^0 N(z) dz, \quad m \geq 1.$$

Coriolis param
($2\Omega\sin\theta$)

stratification

ocean depth



Chelton et al. [1998]

FIG. 6. Global contour map of the $1^\circ \times 1^\circ$ first baroclinic Rossby radius of deformation λ_1 in kilometers computed by Eq. (2.3) from the first baroclinic gravity-wave phase speed shown in Fig. 2. Water depths shallower than 3500 m are shaded.

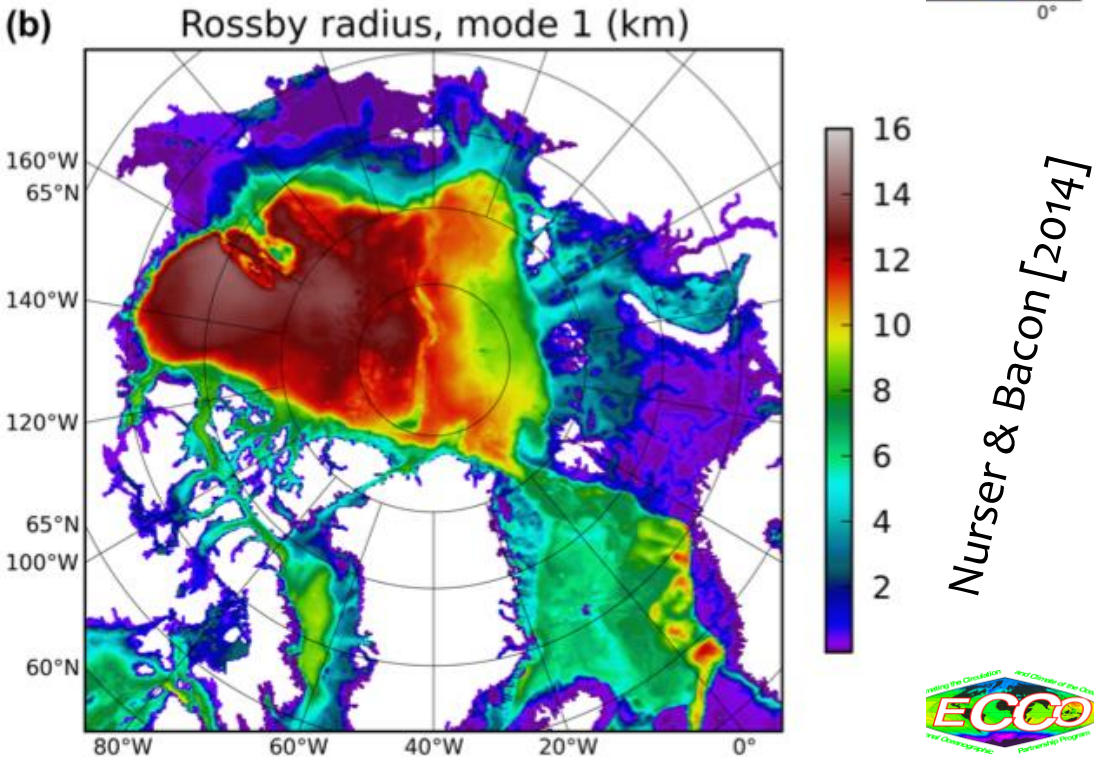
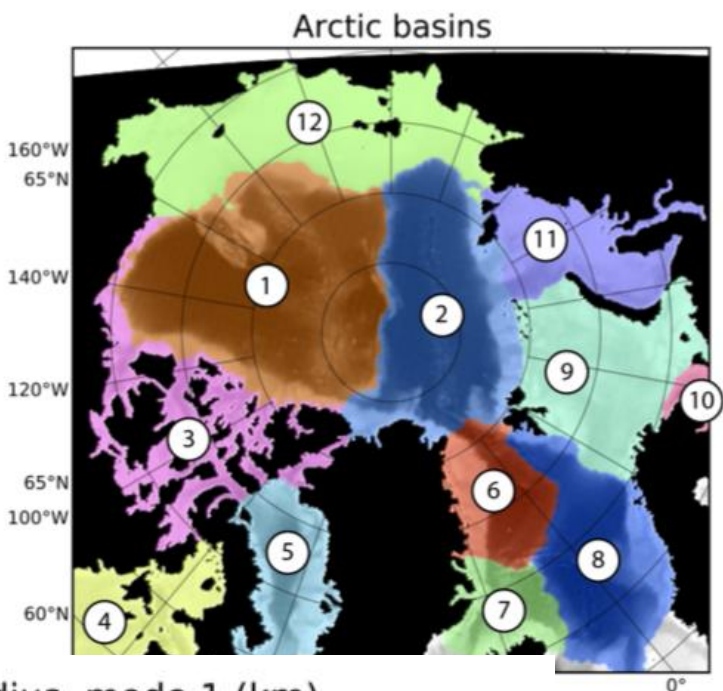
Challenges of (forward) modeling the Arctic

stratification: weak

$$\lambda_m^{\text{WKB}} = \frac{1}{|f|m\pi} \int_{-H}^0 N(z) dz, \quad m \geq 1.$$

large $\rightarrow$ $|f|m\pi$ $\leftarrow$ ocean depth: shallow

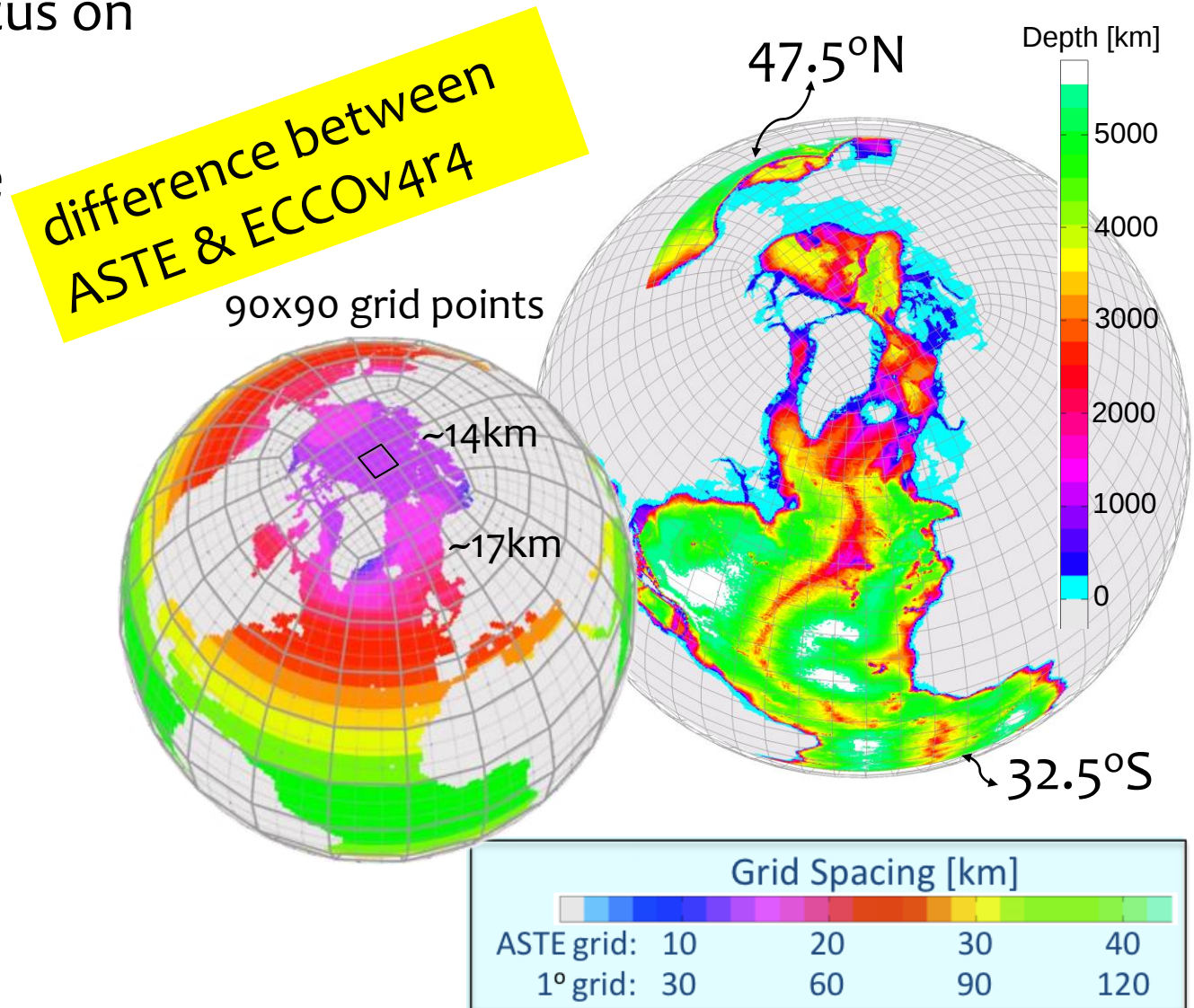
Region	Key	Mode 1 (km)			Mode 2 (km)
		Summer	Winter	Annual	Annual
Deep Arctic Ocean					
Amerasian Basin	1	11.2	11.0	11.1	5.2
Eurasian Basin	2	7.9	7.7	7.8	4.6
Environments of Canadian waters					
Canadian Archipelago	3	6.3	5.9	6.1	2.9
Hudson Bay & Foxe Basin	4	4.8	2.7	3.7	1.9
Baffin Bay	5	6.3	5.4	5.8	2.9
Nordic seas					
Greenland Sea	6	5.1	4.4	4.7	2.6
Iceland Sea	7	5.0	3.8	4.3	2.4
Norwegian Sea	8	7.2	6.5	6.9	3.0
Eurasian Shelf seas					
Barents Sea	9	3.3	1.9	2.5	1.3
White Sea	10	3.5	2.4	2.7	1.4
Kara Sea	11	4.6	3.0	3.8	1.9
East Siberian Shelf seas	12	3.7	2.5	3.1	1.4



Arctic Subpolar gyre sTate Estimate (ASTE)

[Nguyen et al., 2021]

- Higher resolution (1/2° based) with focus on Arctic and subpolar North Atlantic
- OBCs: ECCOV4R* 1° global state estimate
- Initial conditions:
 - T/S: spin up from WOA13v1
 - sea ice: spin up from PSC Jan/2002
- atmospheric forcings: JRA-55
- 3-D ocean mixing parameters:
 - initial guesses
- 2002-2017
 - starts of ICESat & GRACE
 - better coverage of Argo
 - begin of ITP deployment (2004)



Primary data constraints:

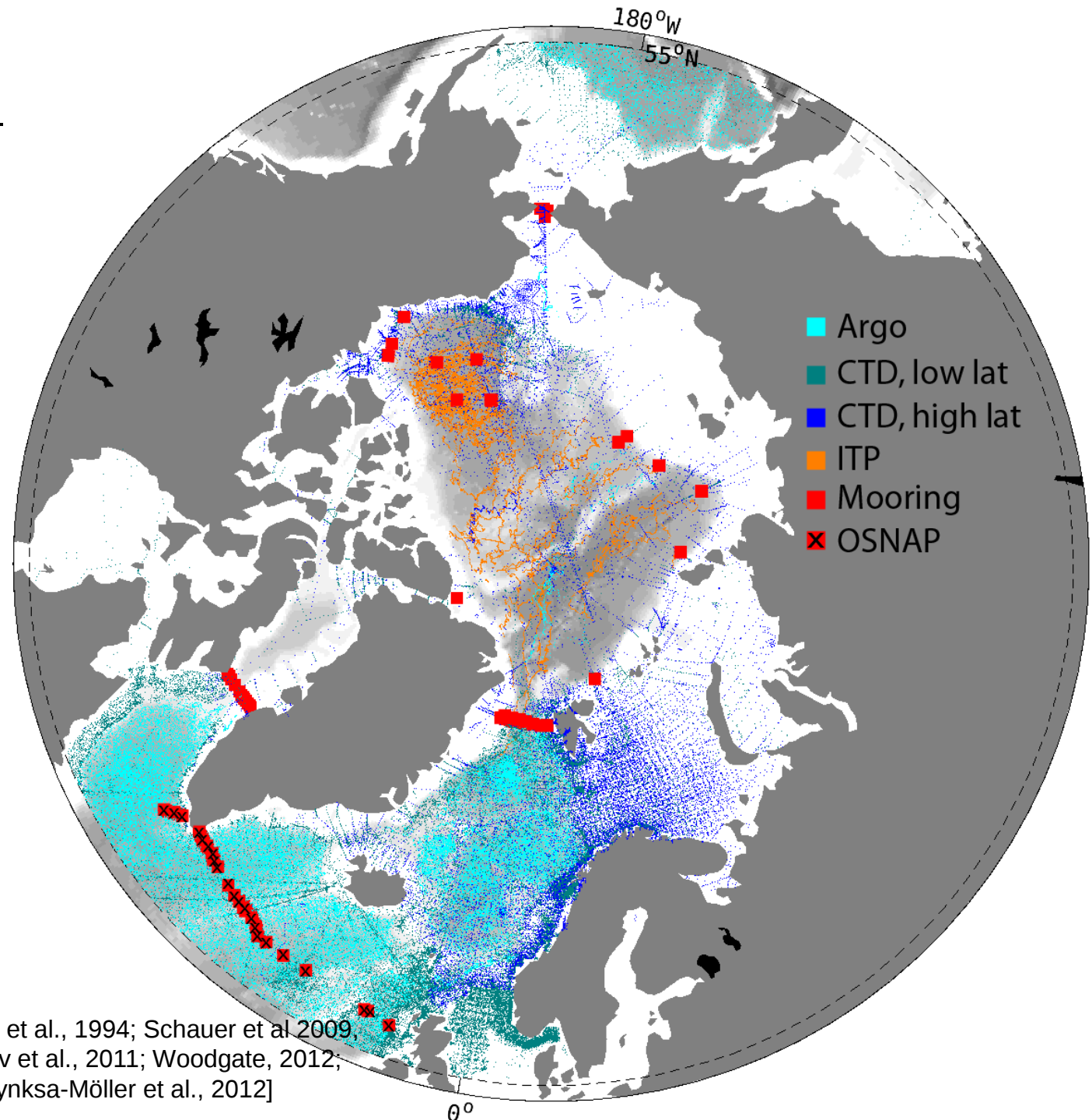
Low latitudes:

- ECCOV4r* standard obs. (SSH, SST, Argo, XBT, CTDs, seal, mean DMT)
- Line-W, Gulf Stream

High latitudes:

- Ice concentration (OSISAF)
- Ice thickness (SMOS+ CryoSat, ICESat), ice drifts
- moored T/S/U/V & hydrographic section, e.g., ITP, NABOS/CABOS, BGEP/BGOS, WHOI & ICES databases
- Arctic main gateways
e.g., Fram, Davis, Bering straits

[Rudels et al., 1994; Schauer et al 2009, Aksenov et al., 2011; Woodgate, 2012; Beszczynska-Möller et al., 2012]



ASTE control space

Initial Greens Function's params
optimized from Nguyen et al. 2011:

- Sea ice drags
- Sea ice albedos
- Internal turbulence mixing $O(10^{-7})$

Important Forward Parameter(s):

- **Viscous dissipation (circulation)**

Adjustable parameters (control space) :

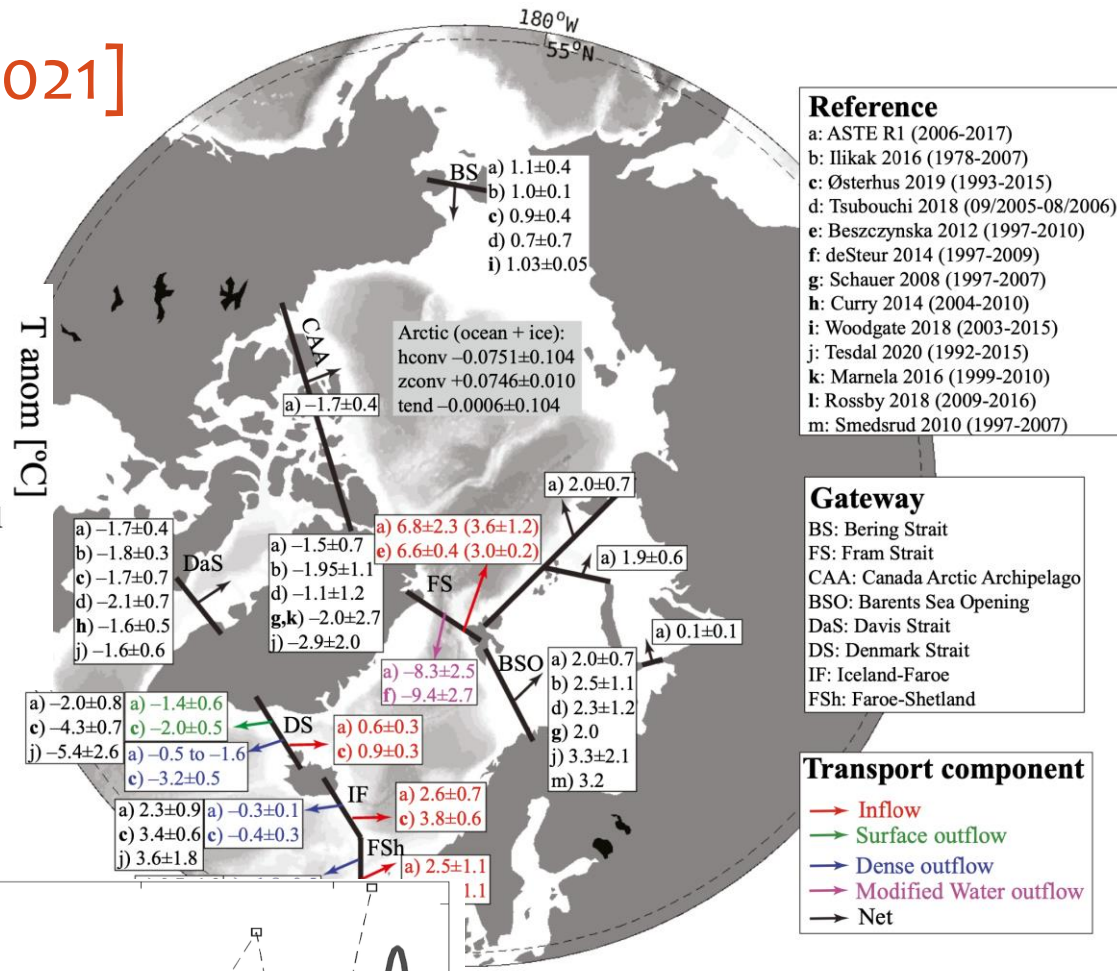
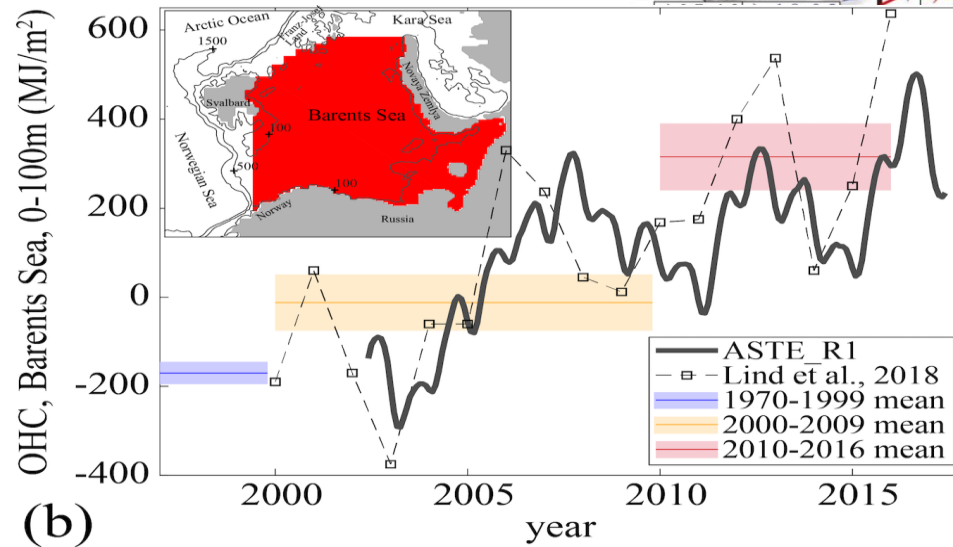
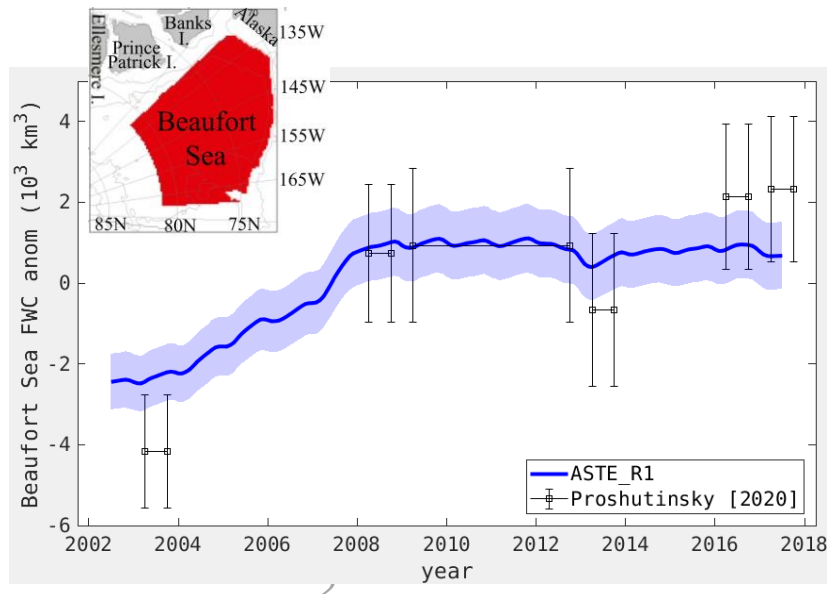
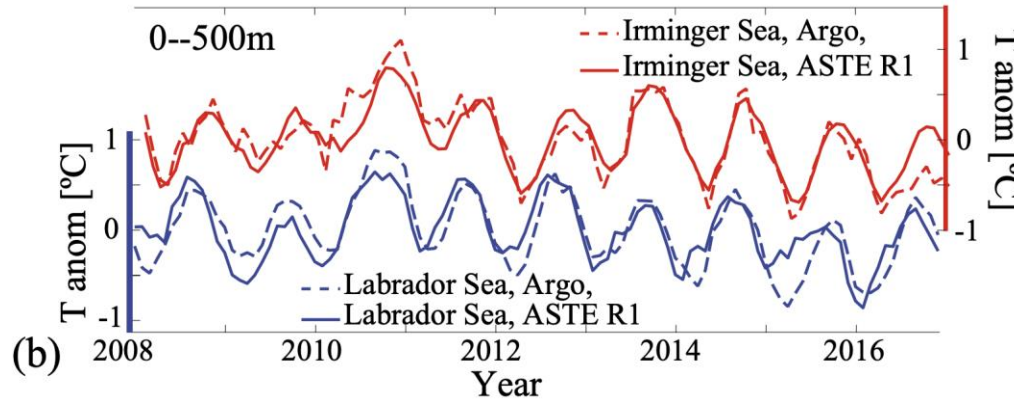
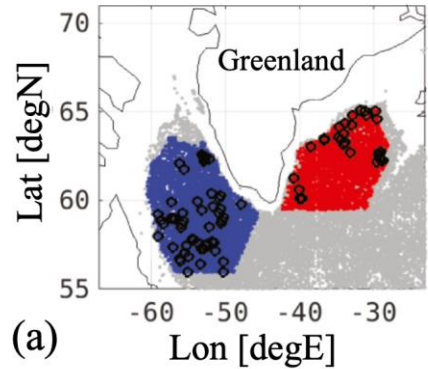
- initial T/S conditions
- time-varying atmospheric forcings
- **3-D internal turbulence mixing**
- Unresolve mesoscale eddies
 - **Redi (isopycnal diffusion of tracers)**
 - **GM (bolus transport)**

Requires:

- Good knowledge of regional oceanography
- Good initial guess of ocean/ice state and input parameters

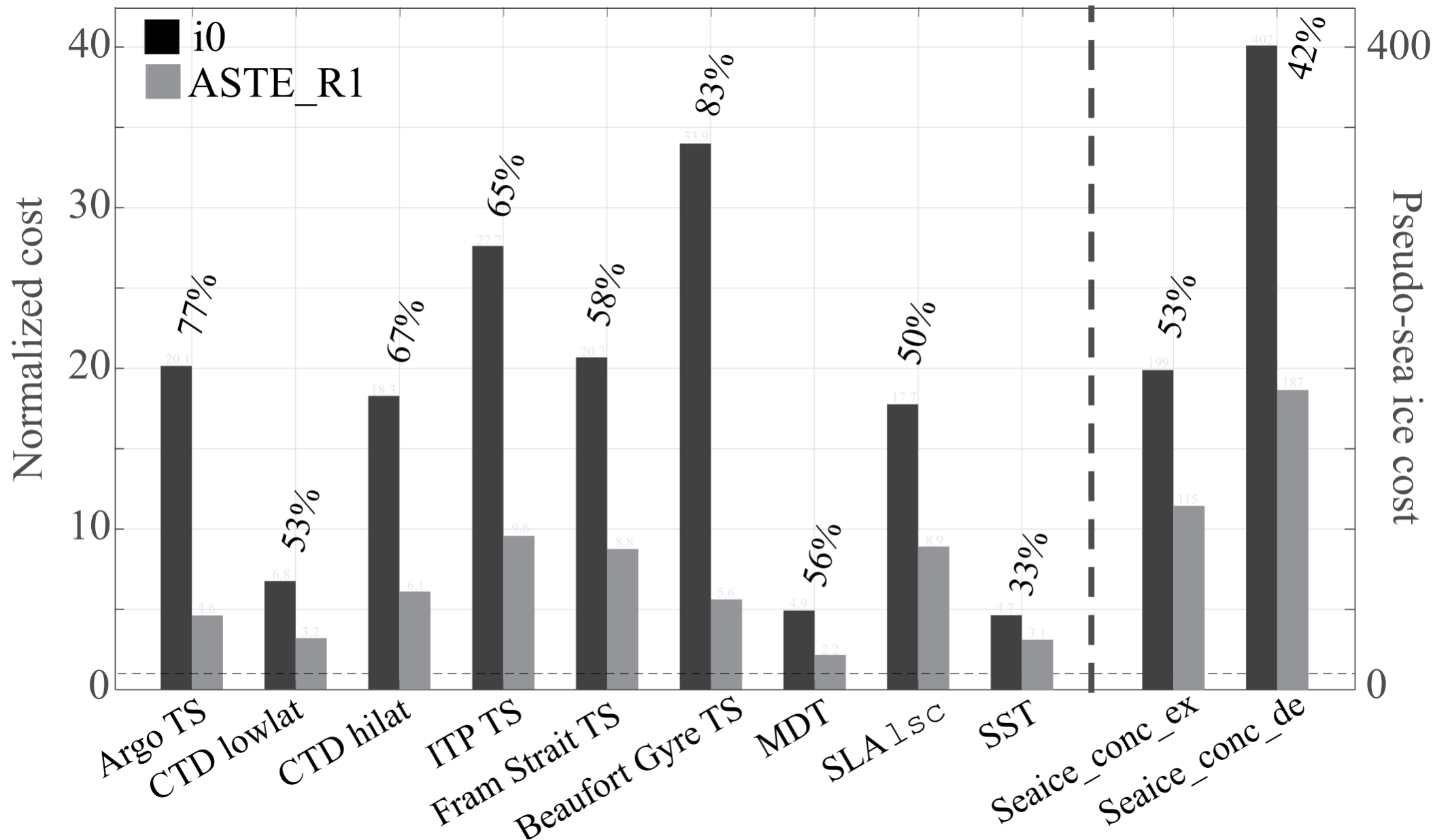
ASTE Release 1 2002–2017 [Nguyen et al., 2021]

Extensive analyses on misfits, bias, budgets

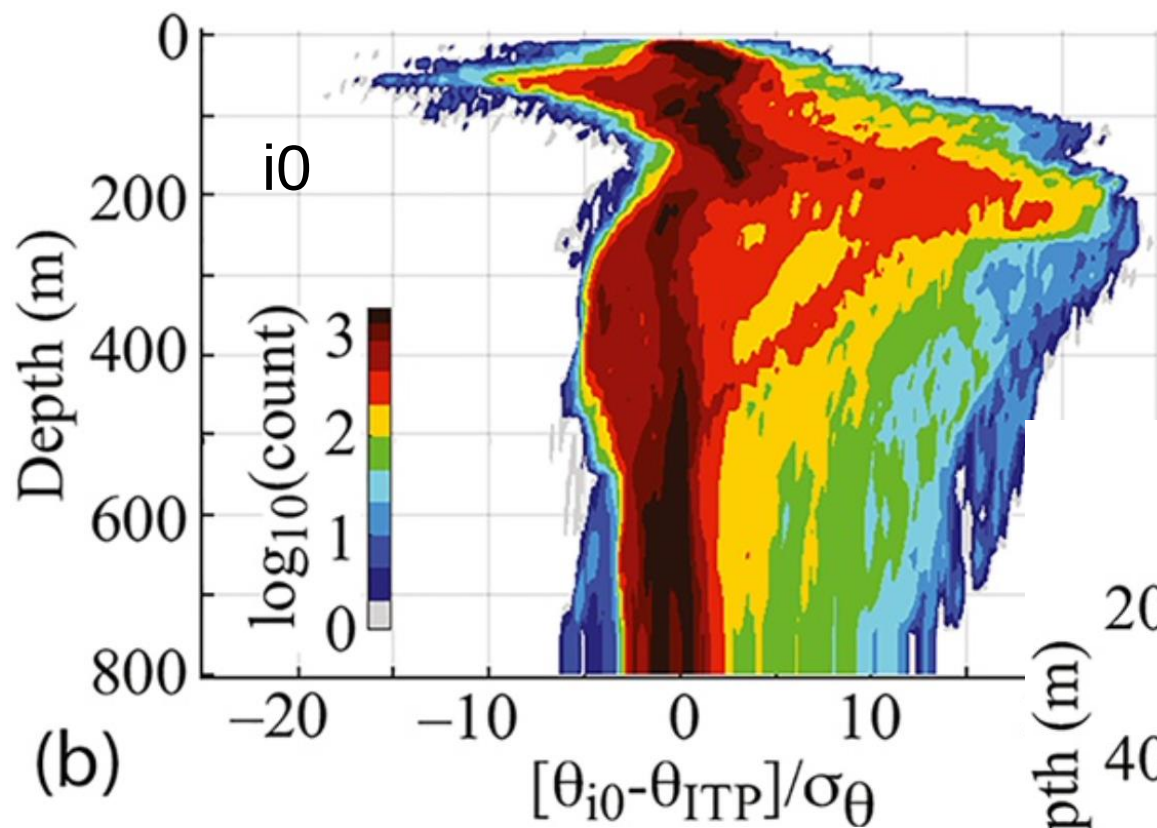


• Data at Arctic Data Portal (arcticdata.io) & UT ECCO-mirror site
<https://web.corral.tacc.utexas.edu/OceanProjects/ASTE/Release1/>

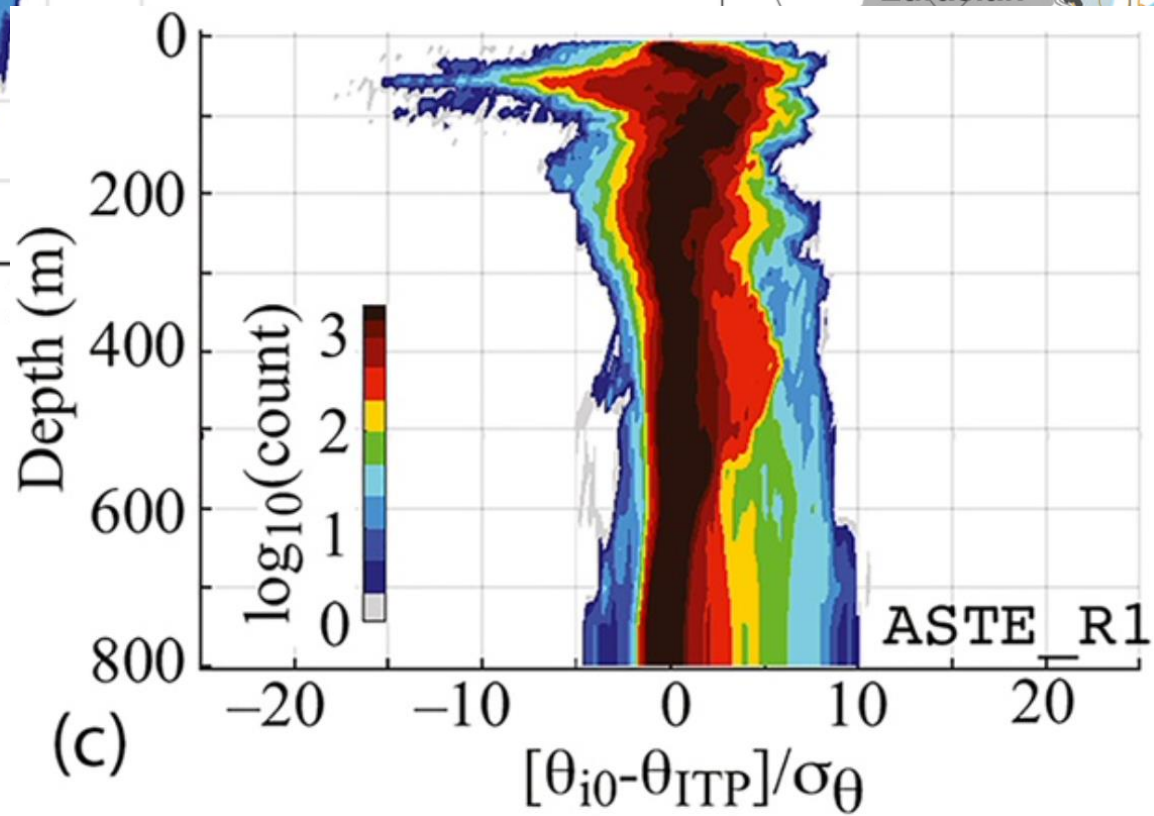
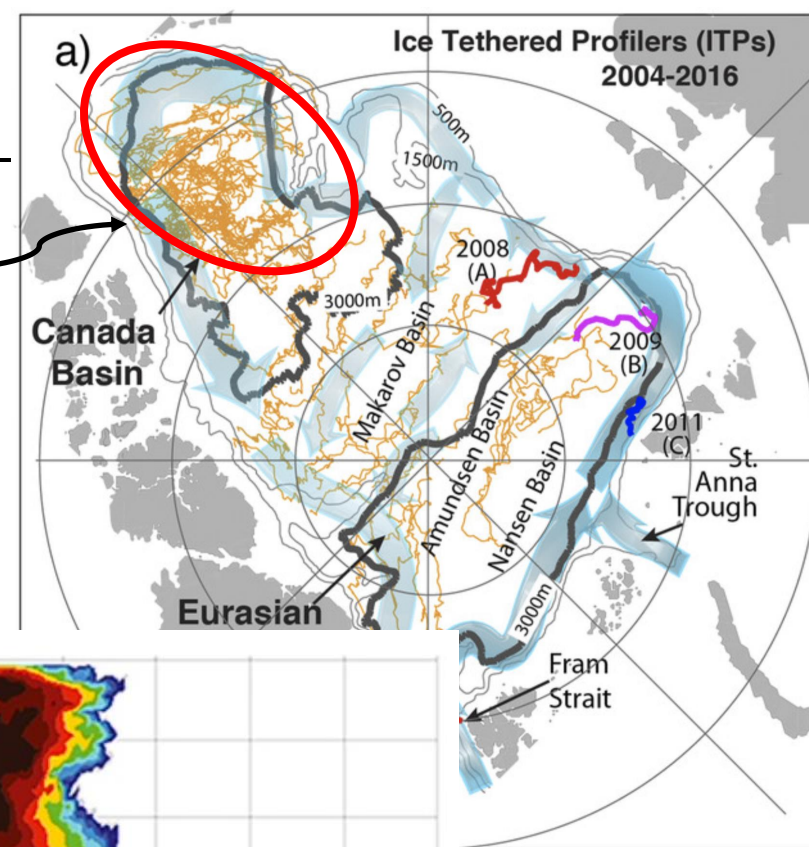
Misfit reductions: Assessment of most impactful datasets



Ice-Tethered Profilers (ITP)

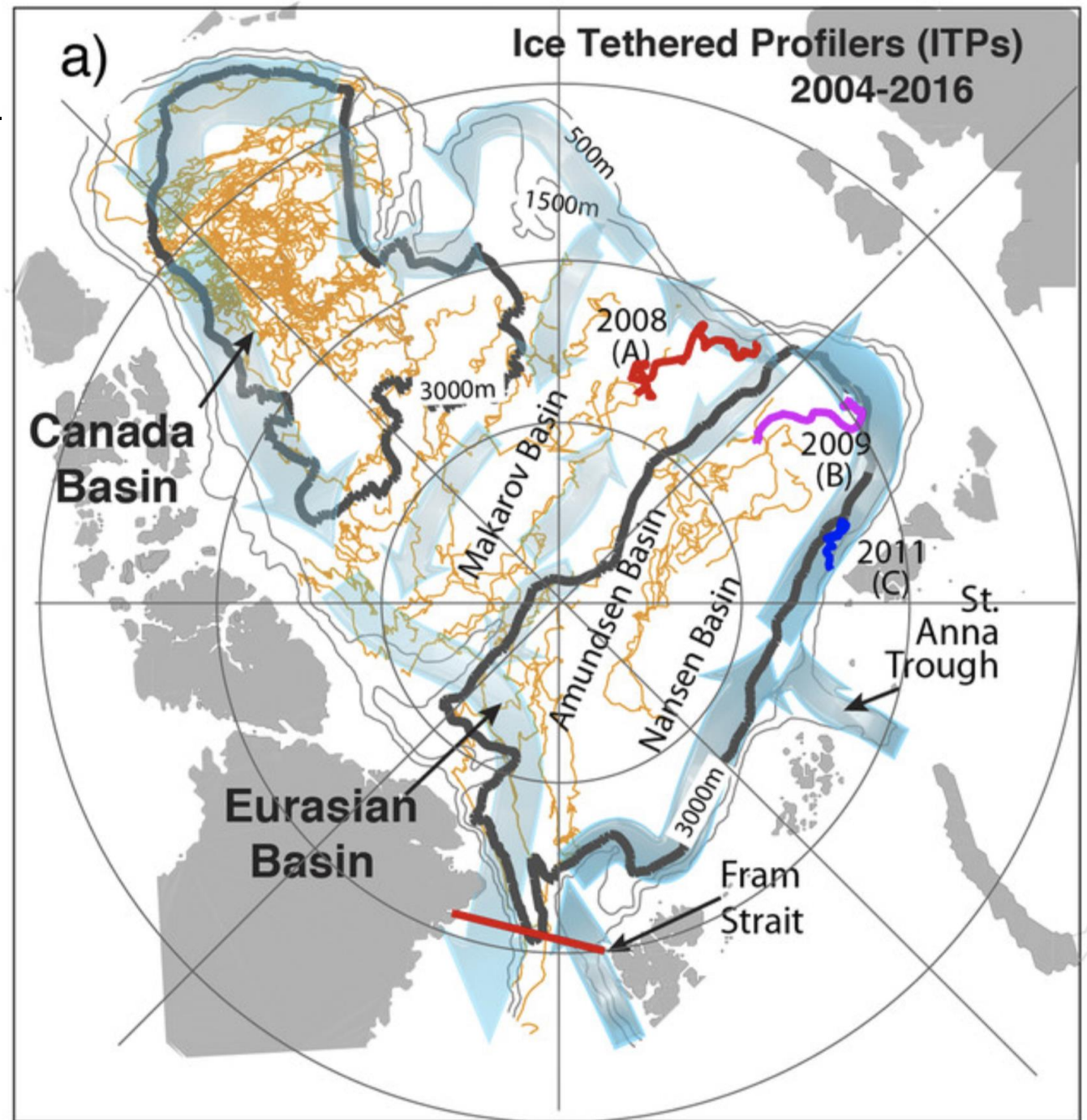
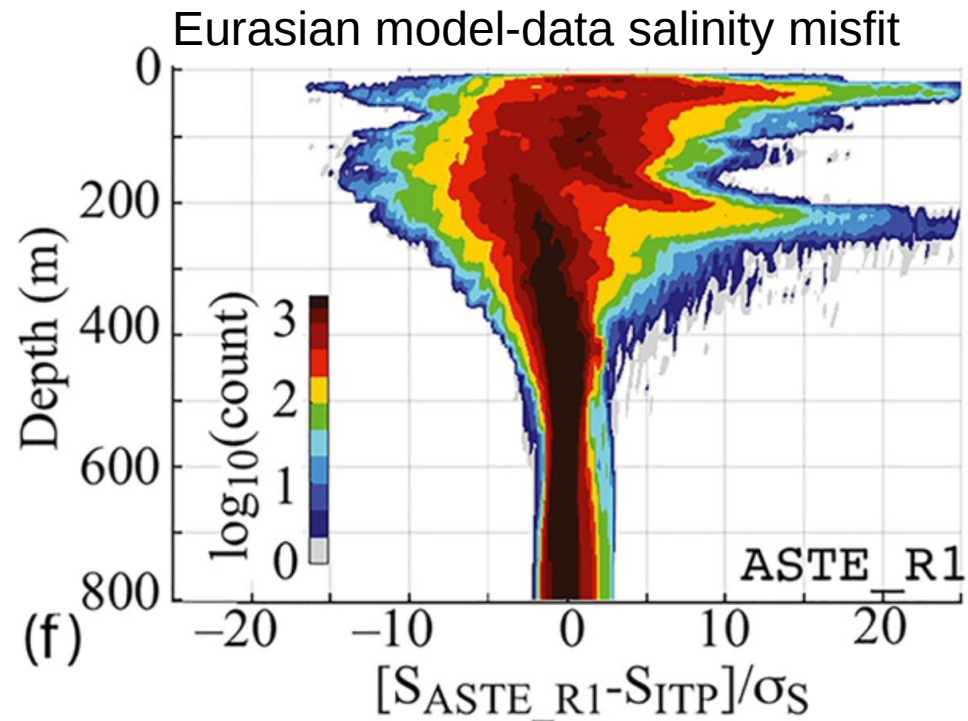


ITP data in the
Canada Basin
interior



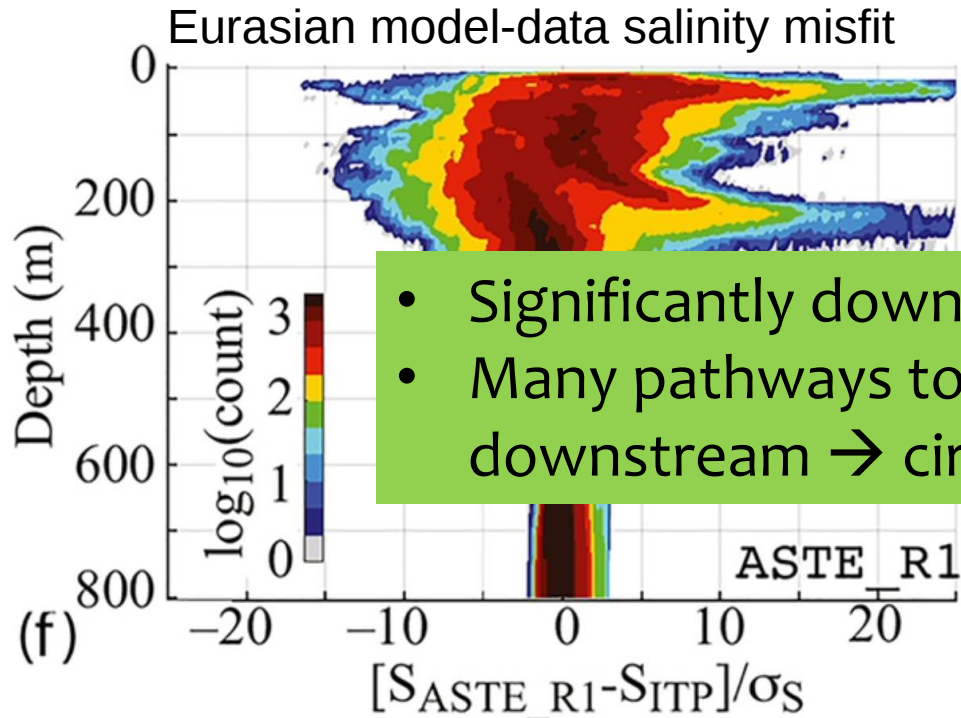
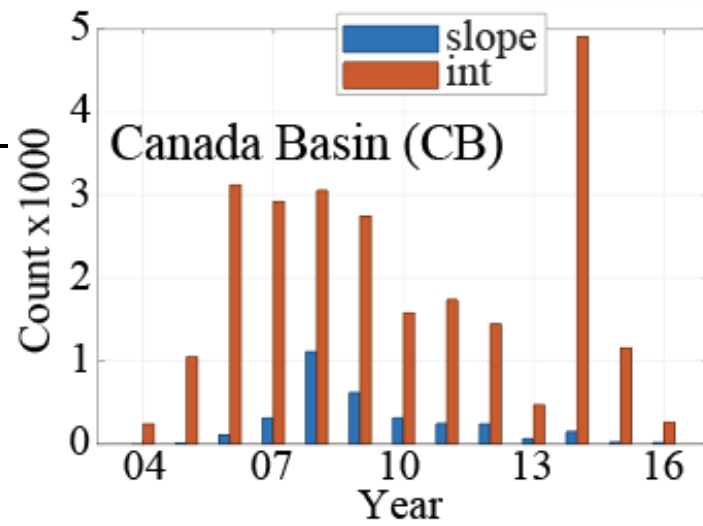
Data distribution

- Inhomogeneous

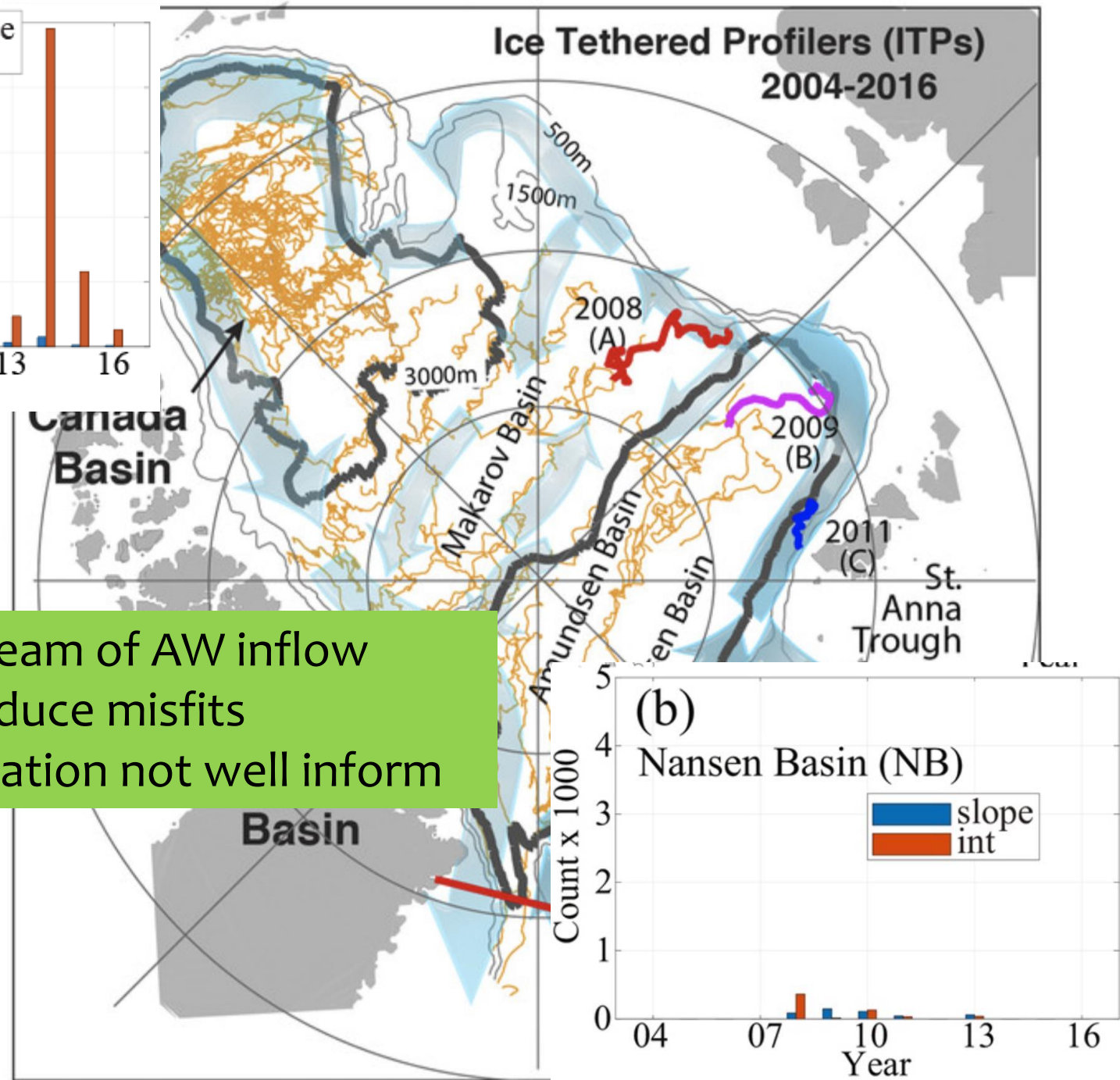


Data distribution

- Inhomogeneous



- Significantly downstream of AW inflow
- Many pathways to reduce misfits downstream $\rightarrow$ circulation not well inform



From idealized / theoretical: constraint on circulation

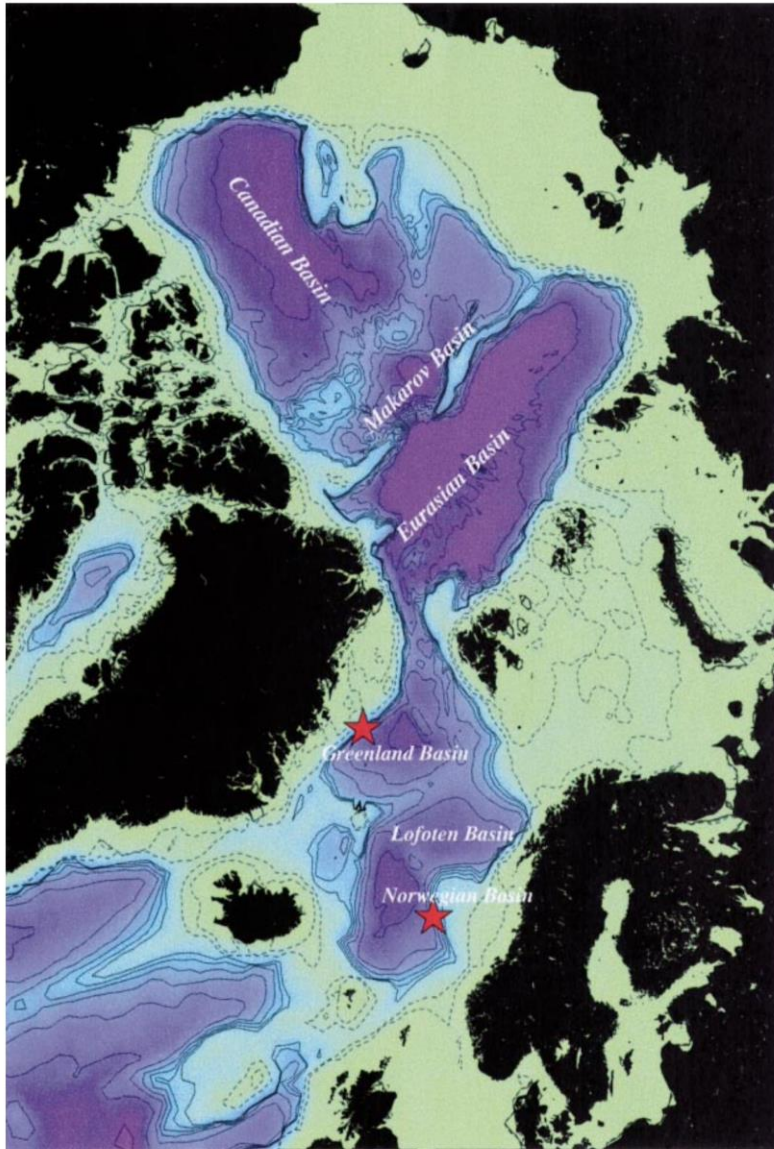
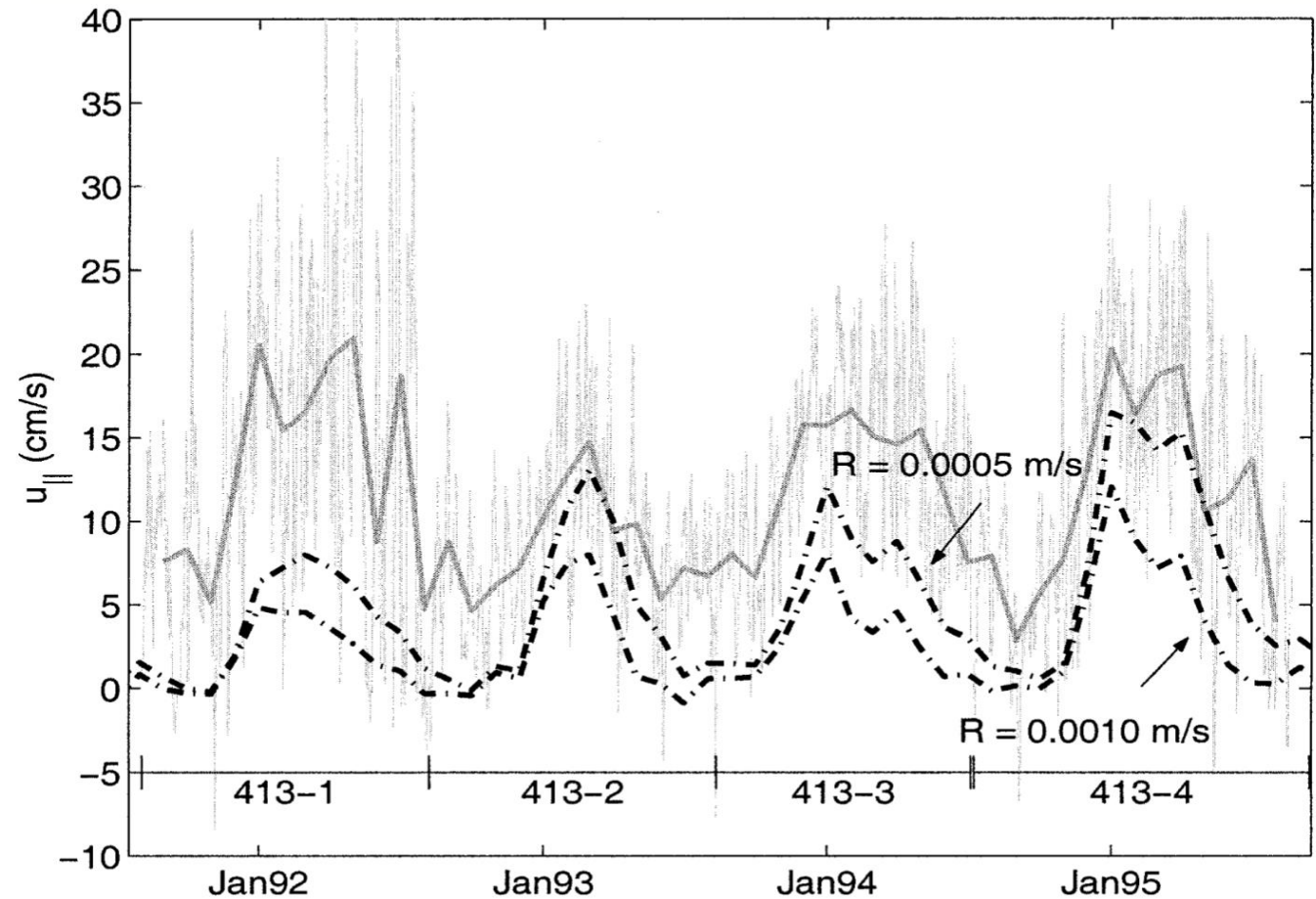


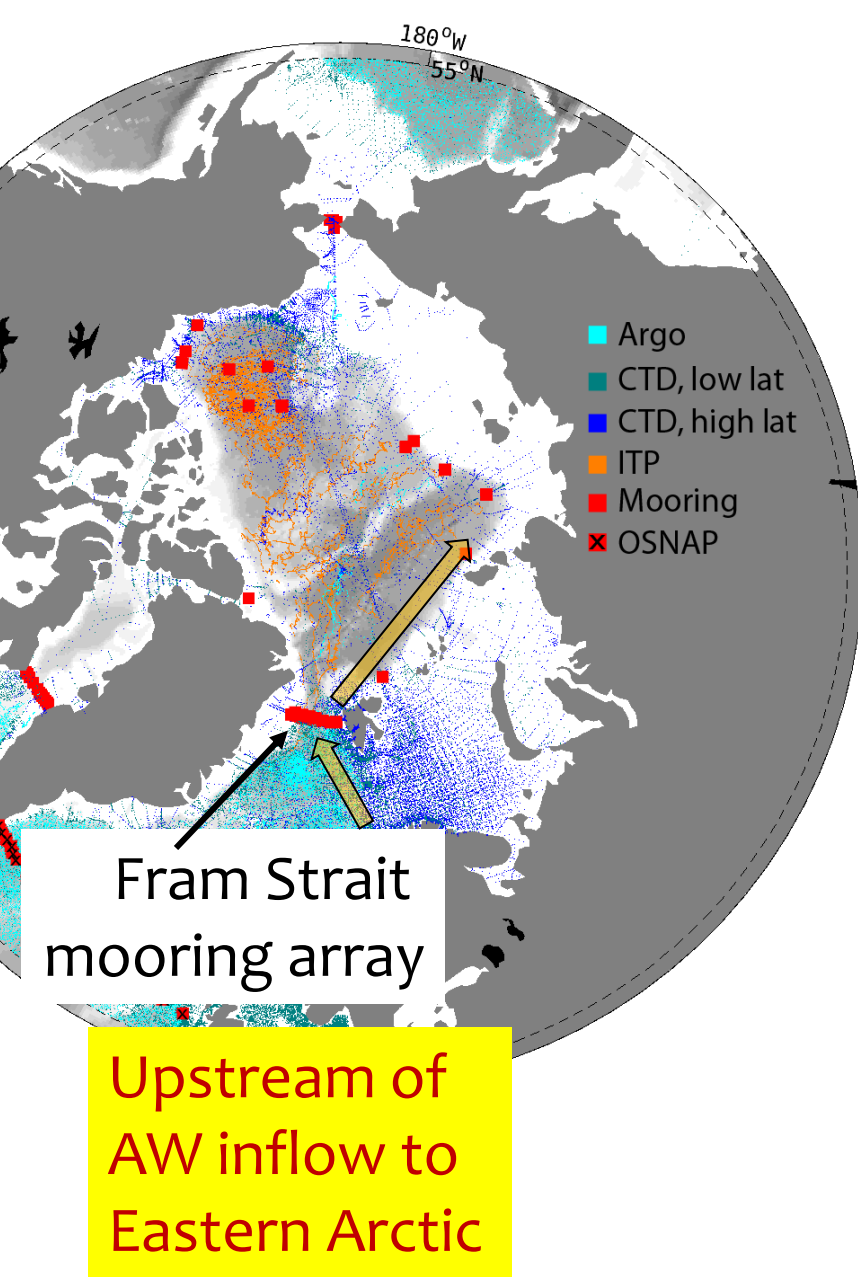
FIG. 3. The f/H field (lines) and bathymetric field (shades) for the Nordic seas and Arctic Ocean. The location of two moorings discussed later are shown with stars. Note the f/H contours approximately follow the isobaths.

western moorings (numbers 410–413), the principal components of the currents align closely with the local tangent of the f/H contours. This component explains more than 85% of the total variance when currents are

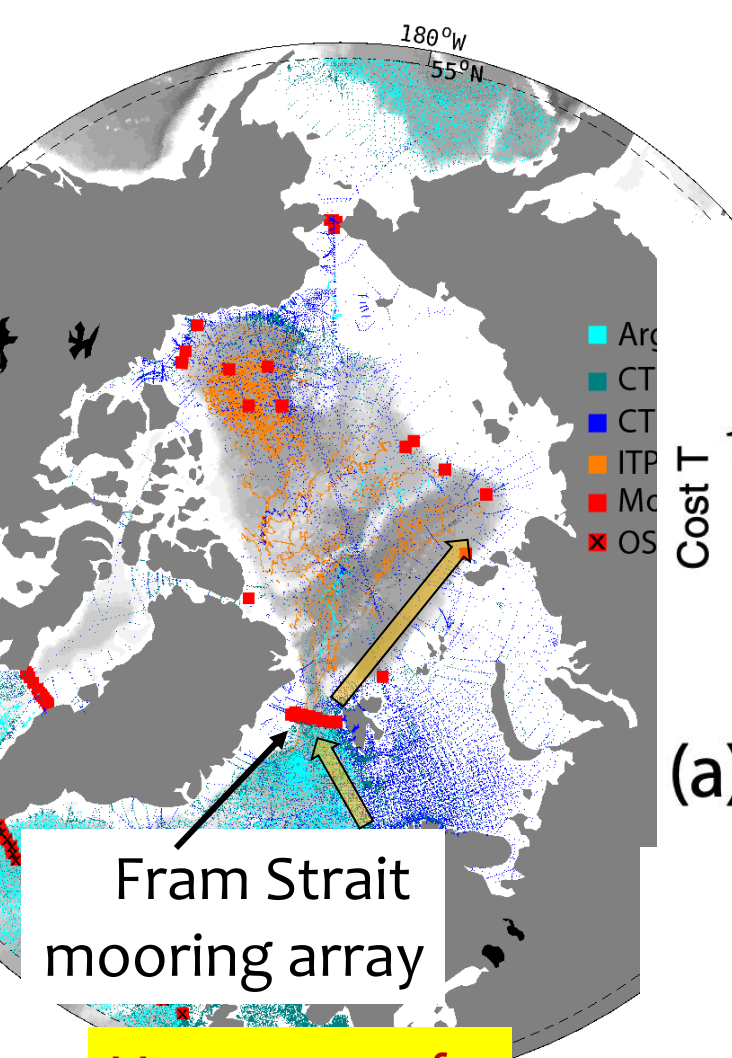


Isachsen et al., JPO 2003

Fram Strait Mooring Array

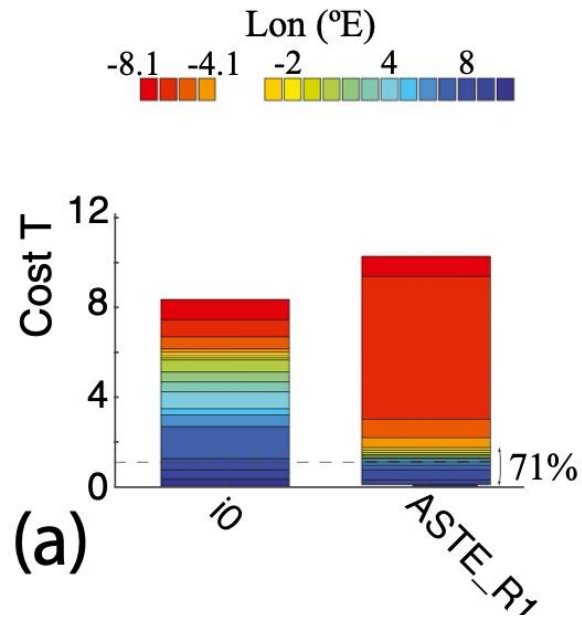


Fram Strait Mooring Array

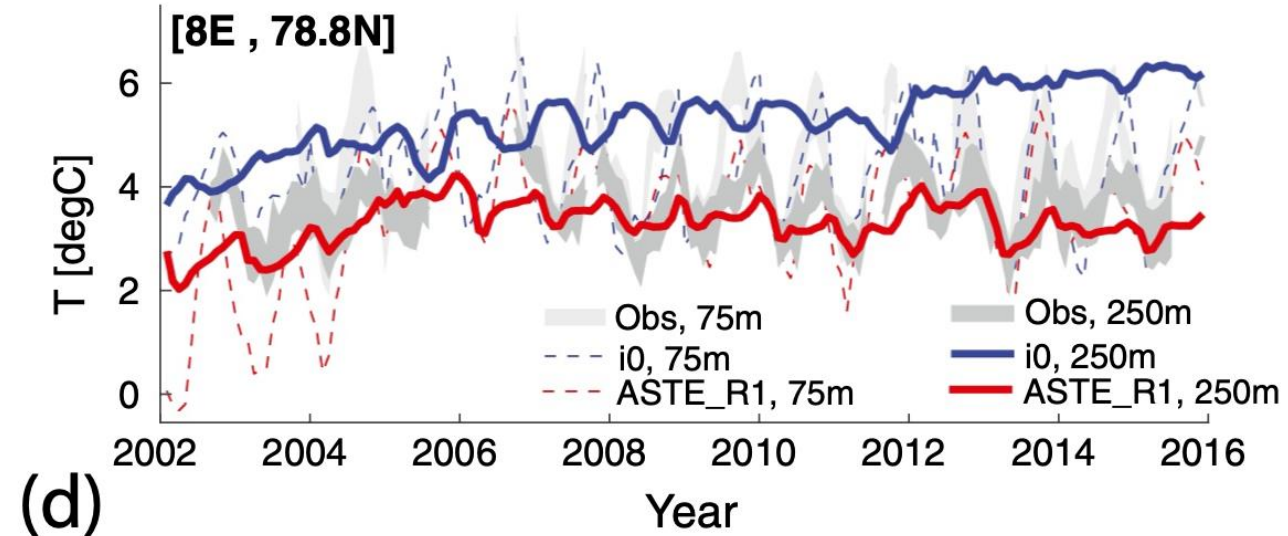


Fram Strait
mooring array

Upstream of
AW inflow to
Eastern Arctic



(a)



(d)

Parameter estimation:

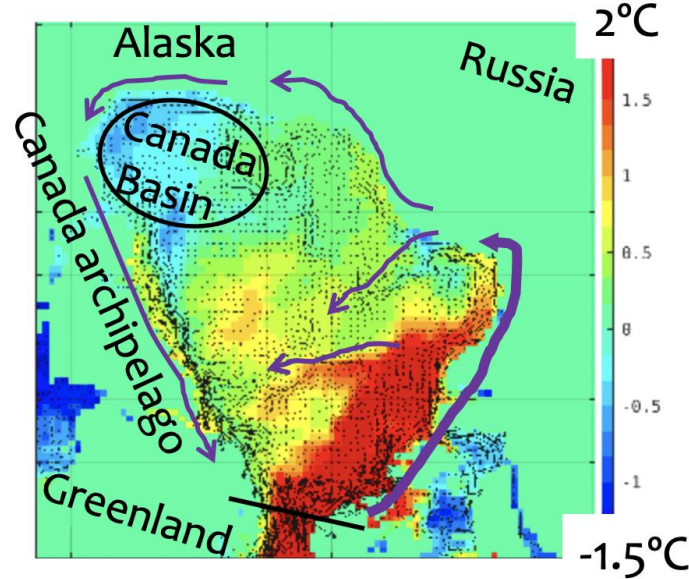
Uncertain parameters:

- Viscous dissipation (circulation)
- 3-D internal turbulence mixing
- Unresolve mesoscale eddies (Redi)

When all parameters are highly uncertain
→ non-unique pathway to reduce indirect misfits

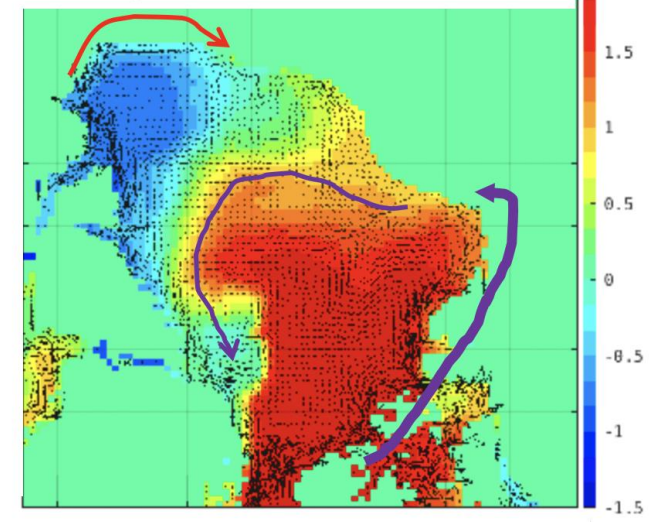
Core AW ~225m

(a)

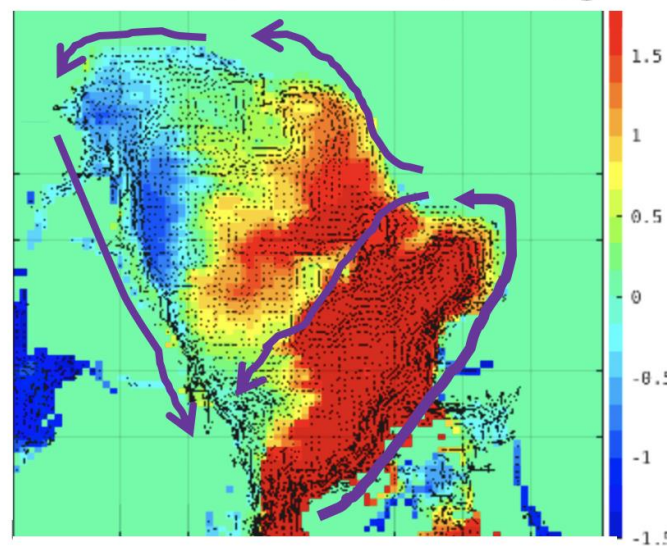


(b)

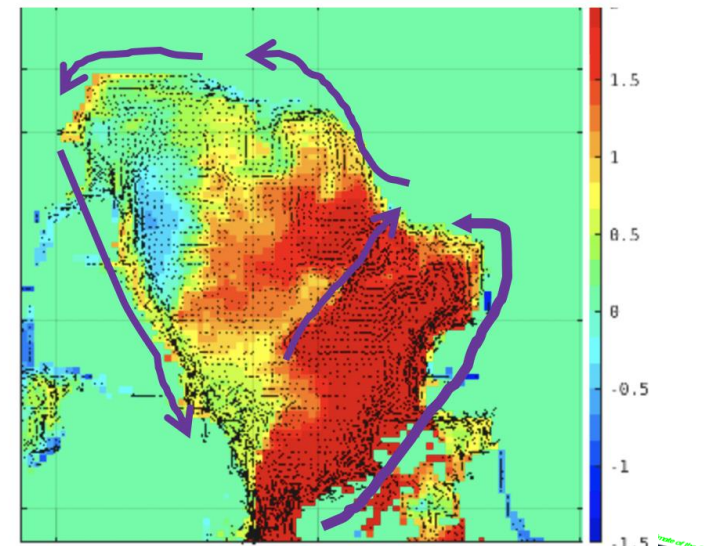
Higher Ah viscosity



(c) Same as b, lower kapredi/gm



(d) Same as c, lower diffkr



Mixing at depths

Staircases:

Observed throughout the
Arctic ocean interior

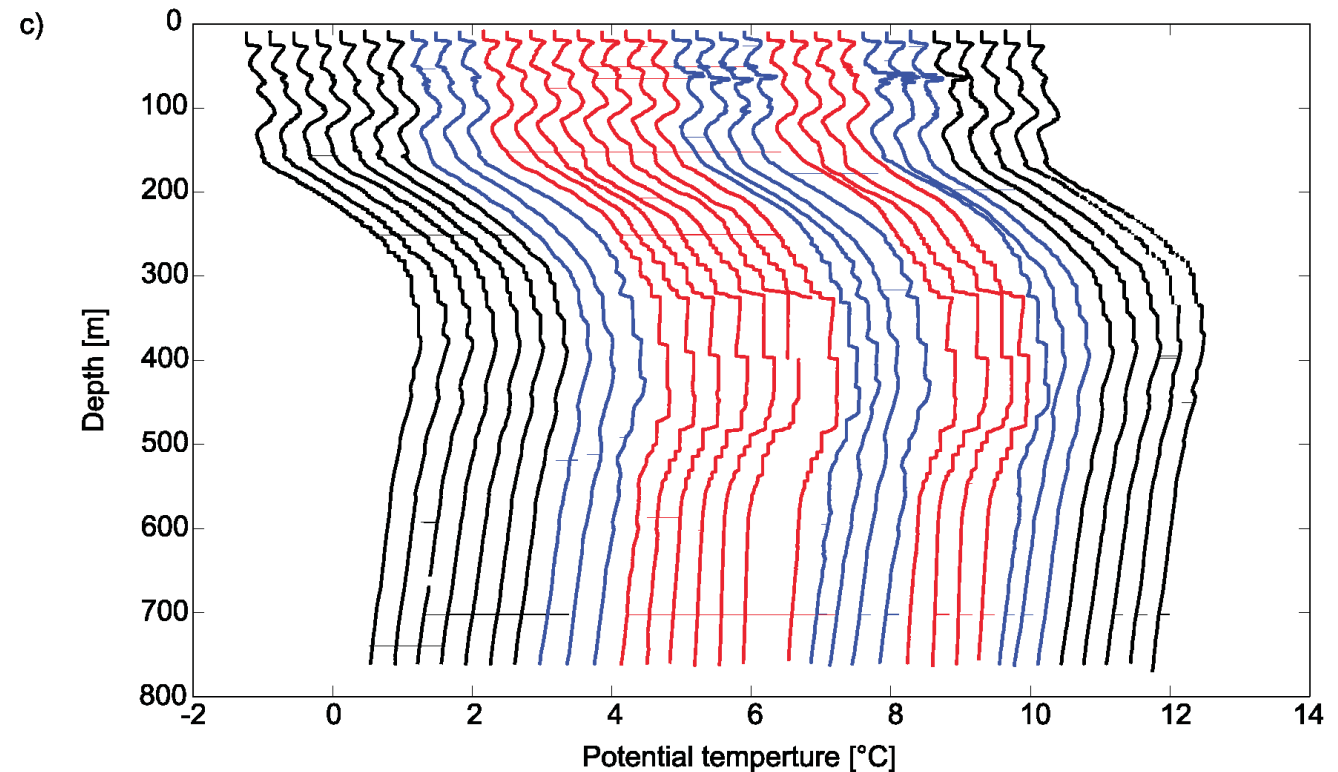
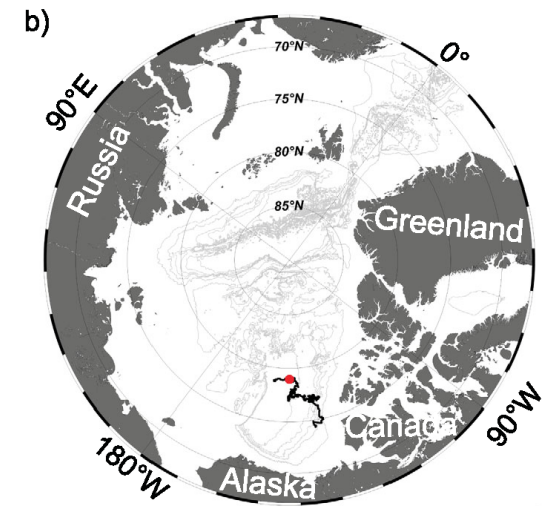
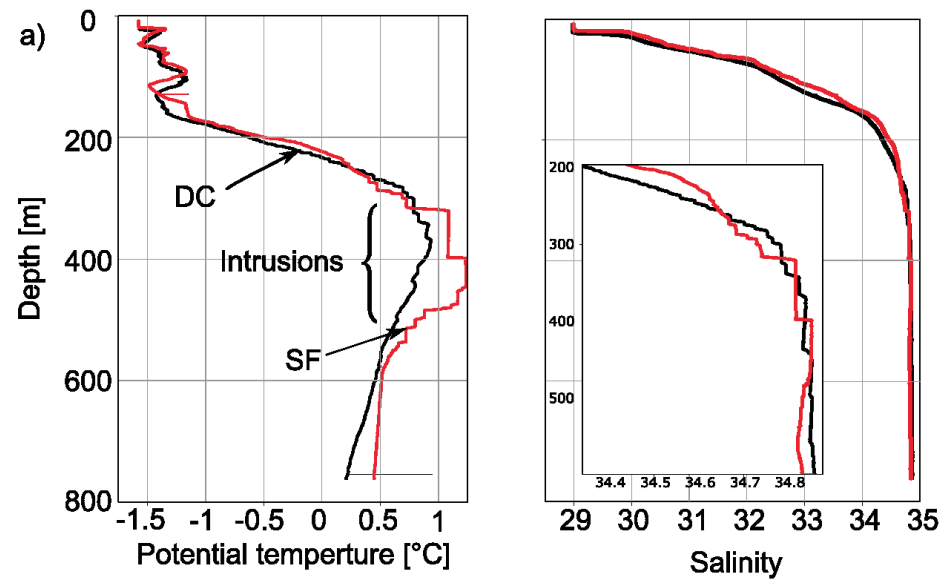
Lateral extent ~5000km

Processes:

- Diffusive convection
- Salt fingering

(molecular level)

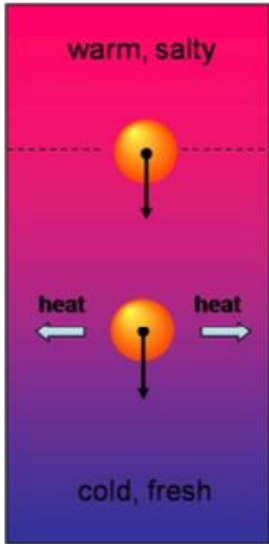
“quiet ocean?”



Arctic Ocean: interior mixing

Double diffusion: stratified fluid, two components having different molecular diffusivities
Sea water: temperature diffuses $\sim 100\times$ faster than salt

Case 1: Salt fingering:



Schematic of the salt fin

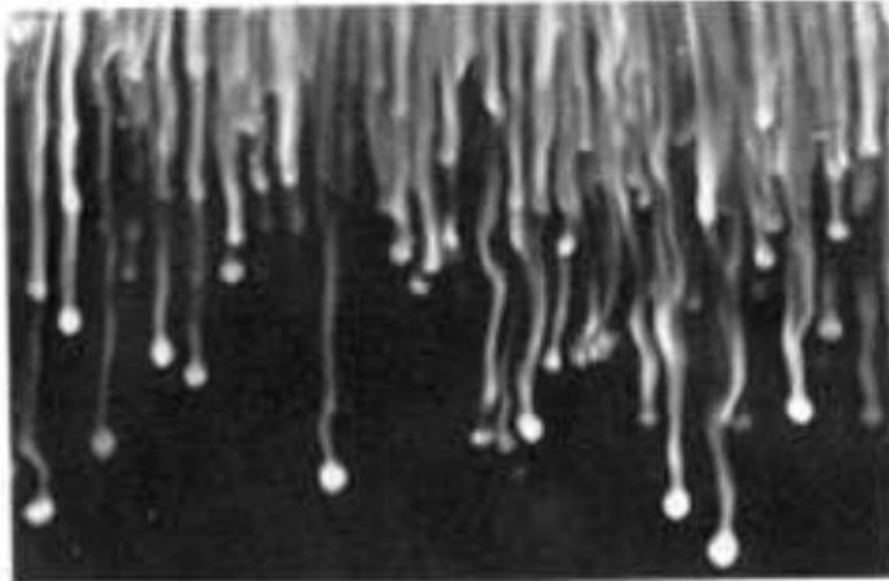
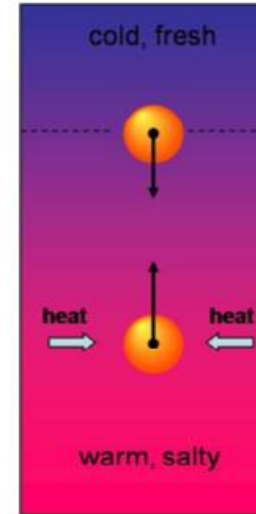
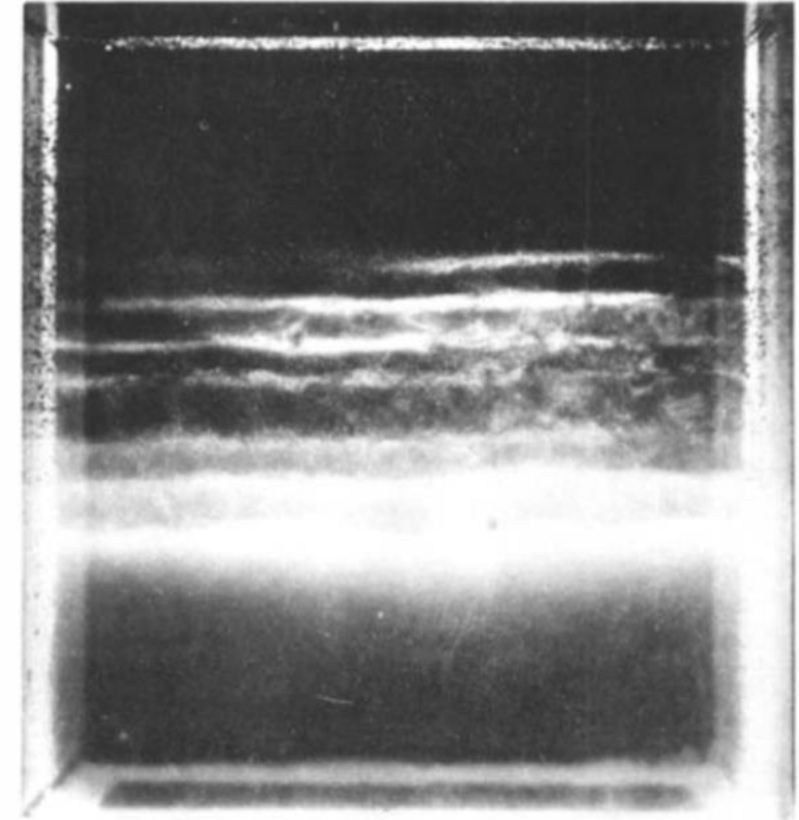


Figure 2: Salt fingers in a laboratory experiment [3].

Case 2: Diffusive convection:

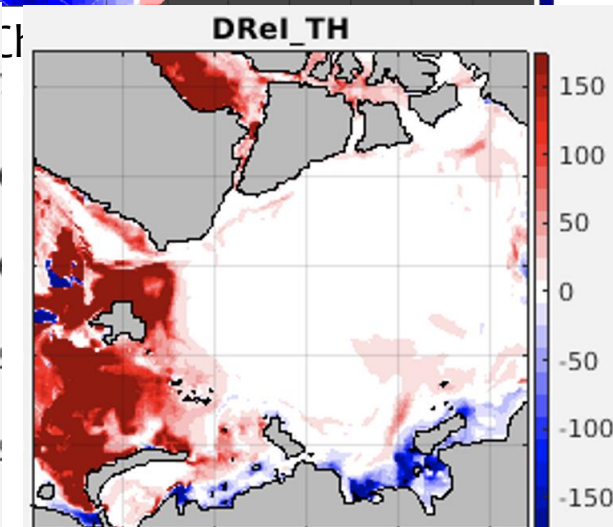
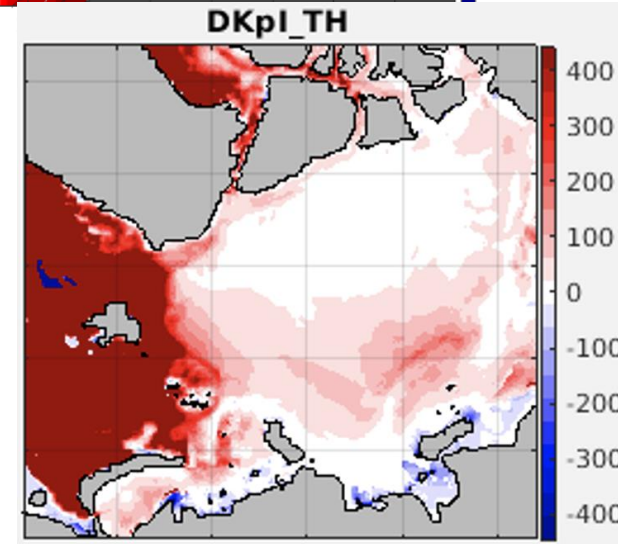
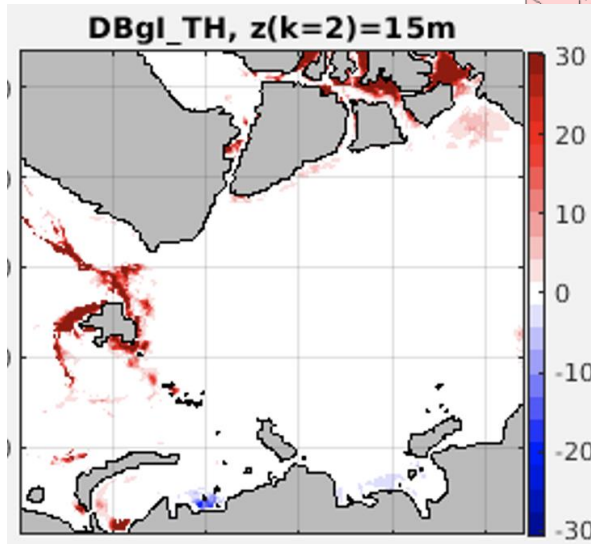
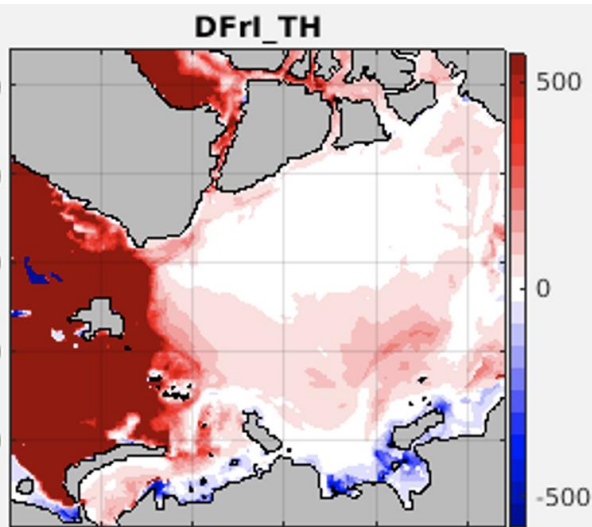
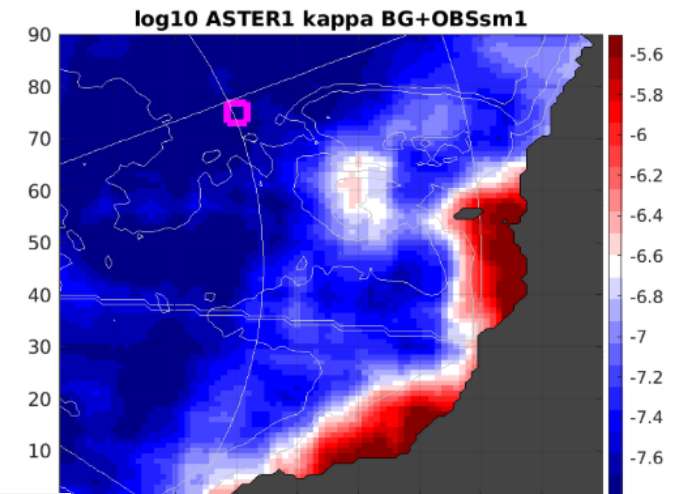
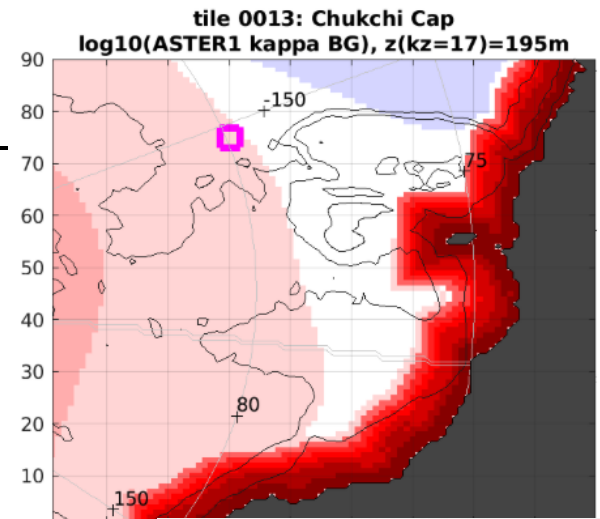


Schematic of the diffusive con



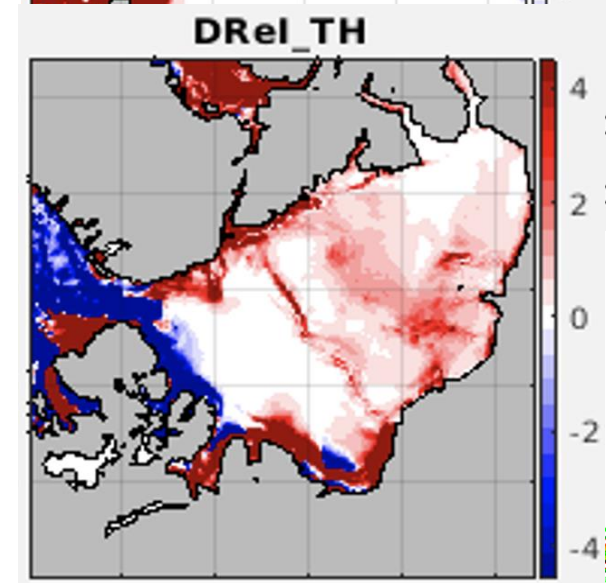
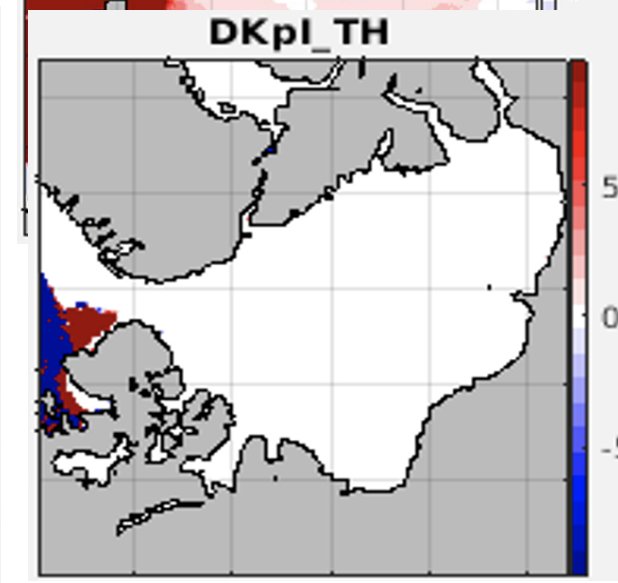
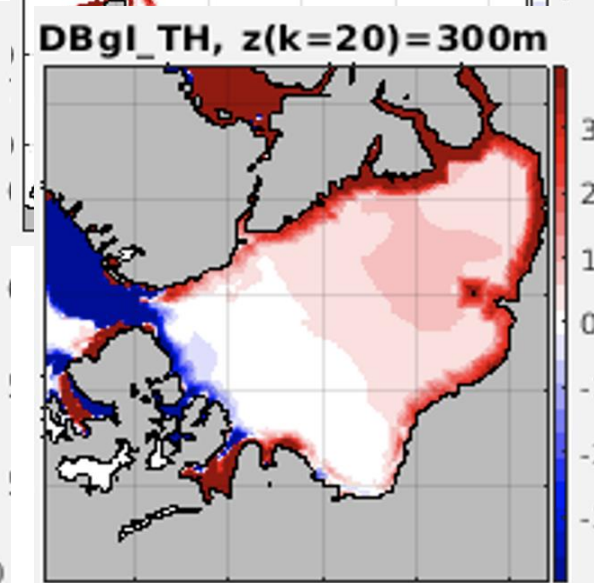
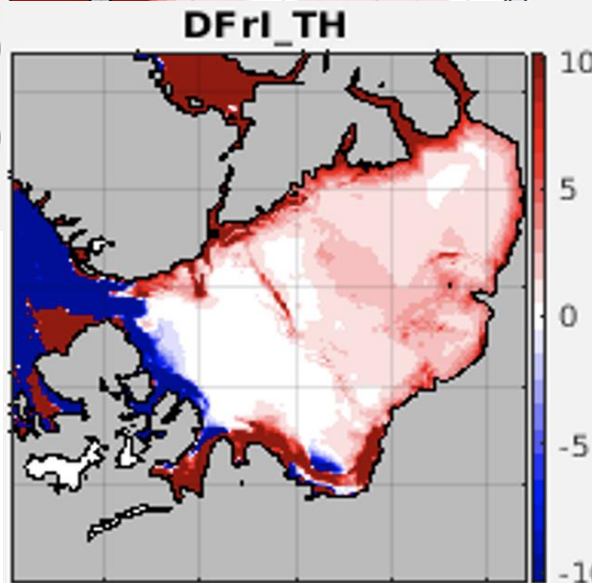
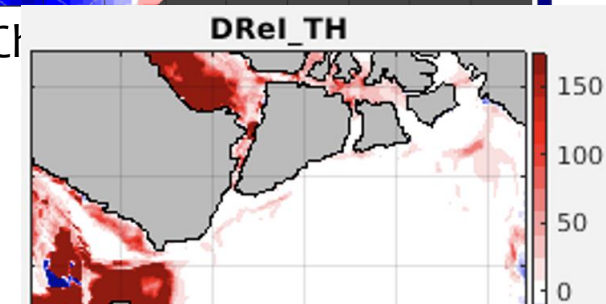
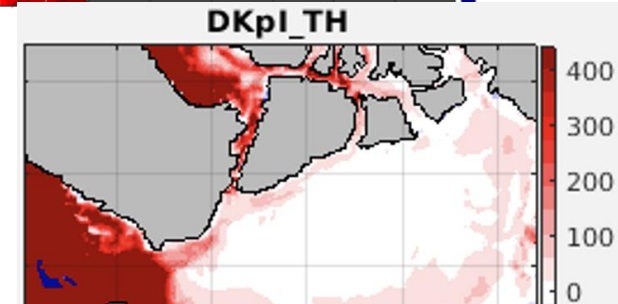
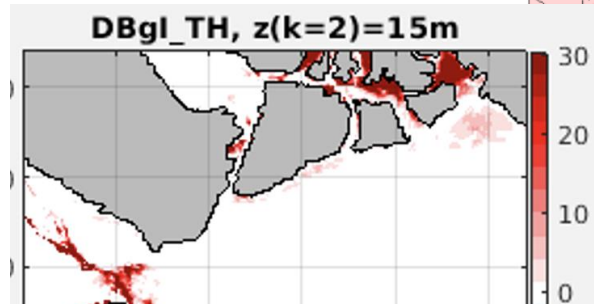
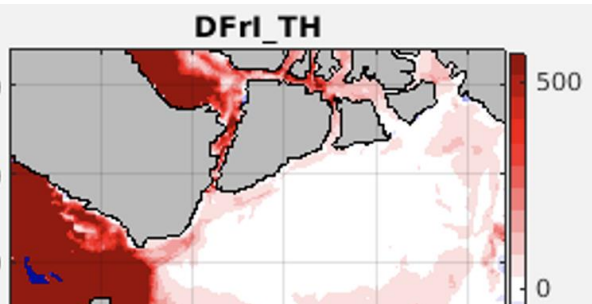
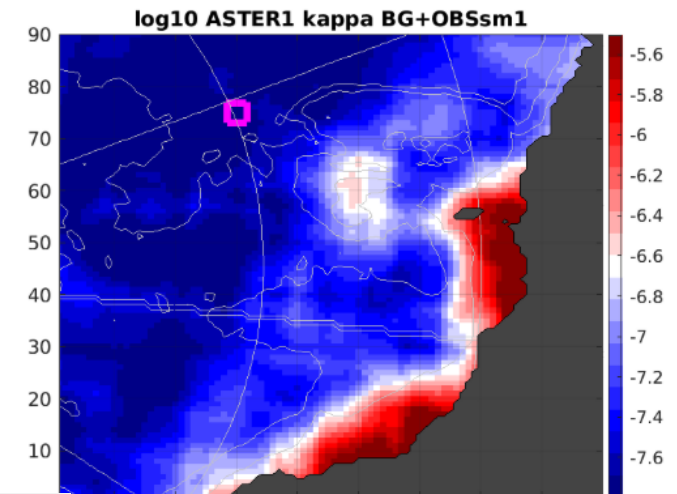
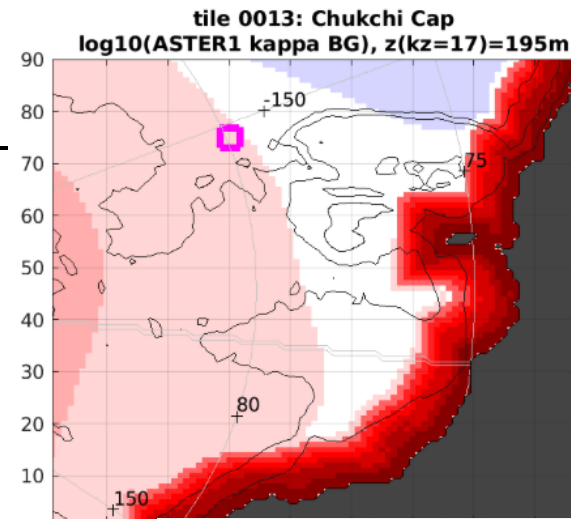
Improving ASTE:

Diffusive flux budget
(Aidan Parfett, UBC)



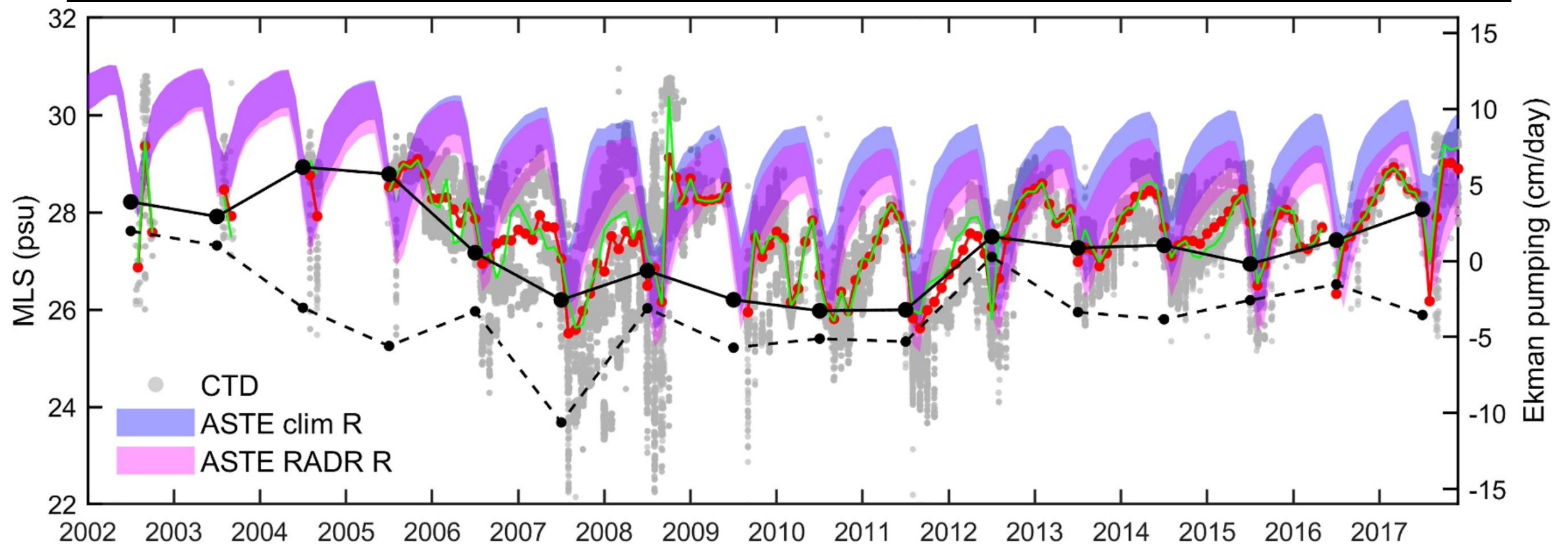
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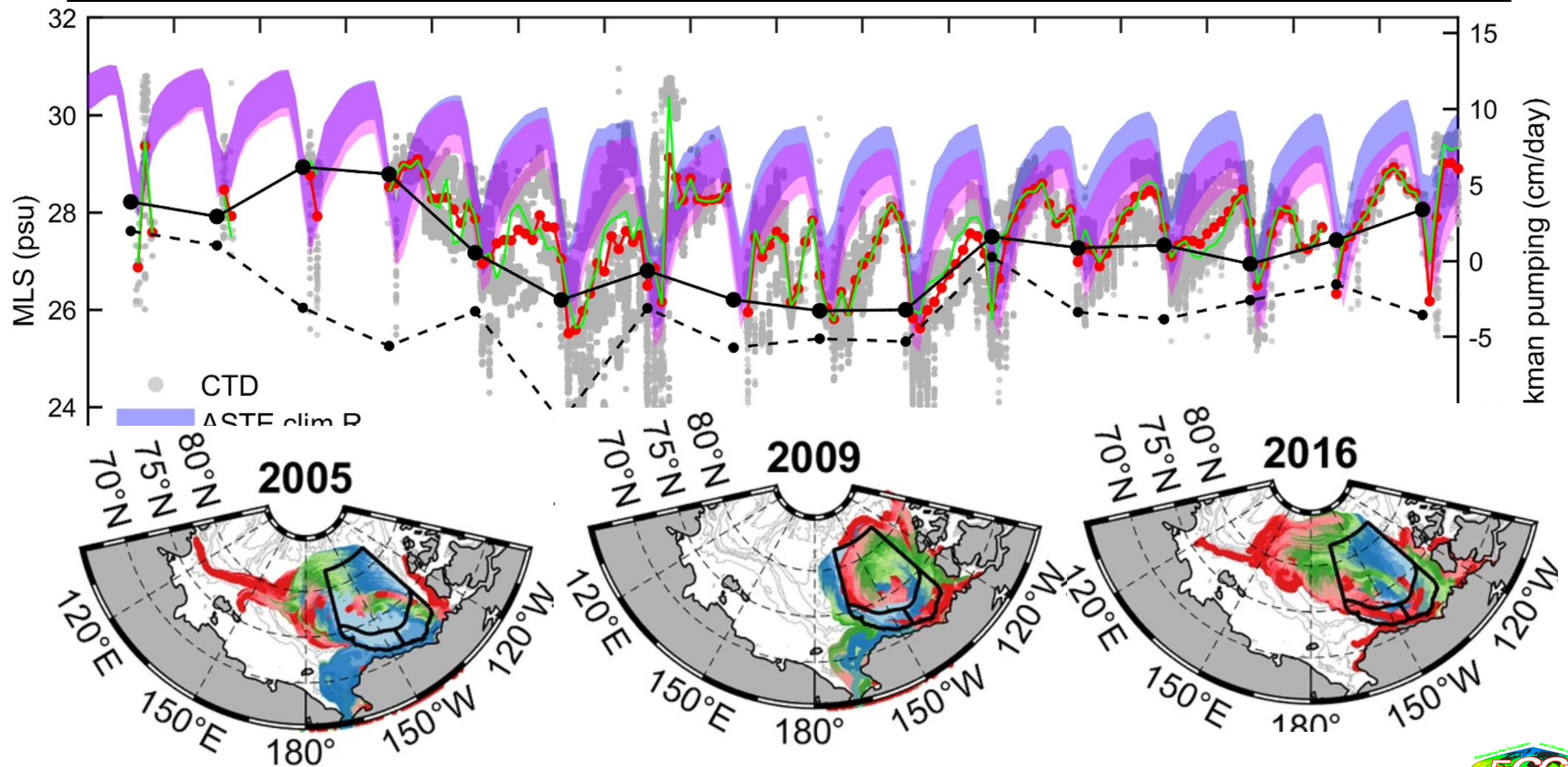
Improving ASTE: ML salinity budget in Beaufort Sea

Zhong et al., 2024



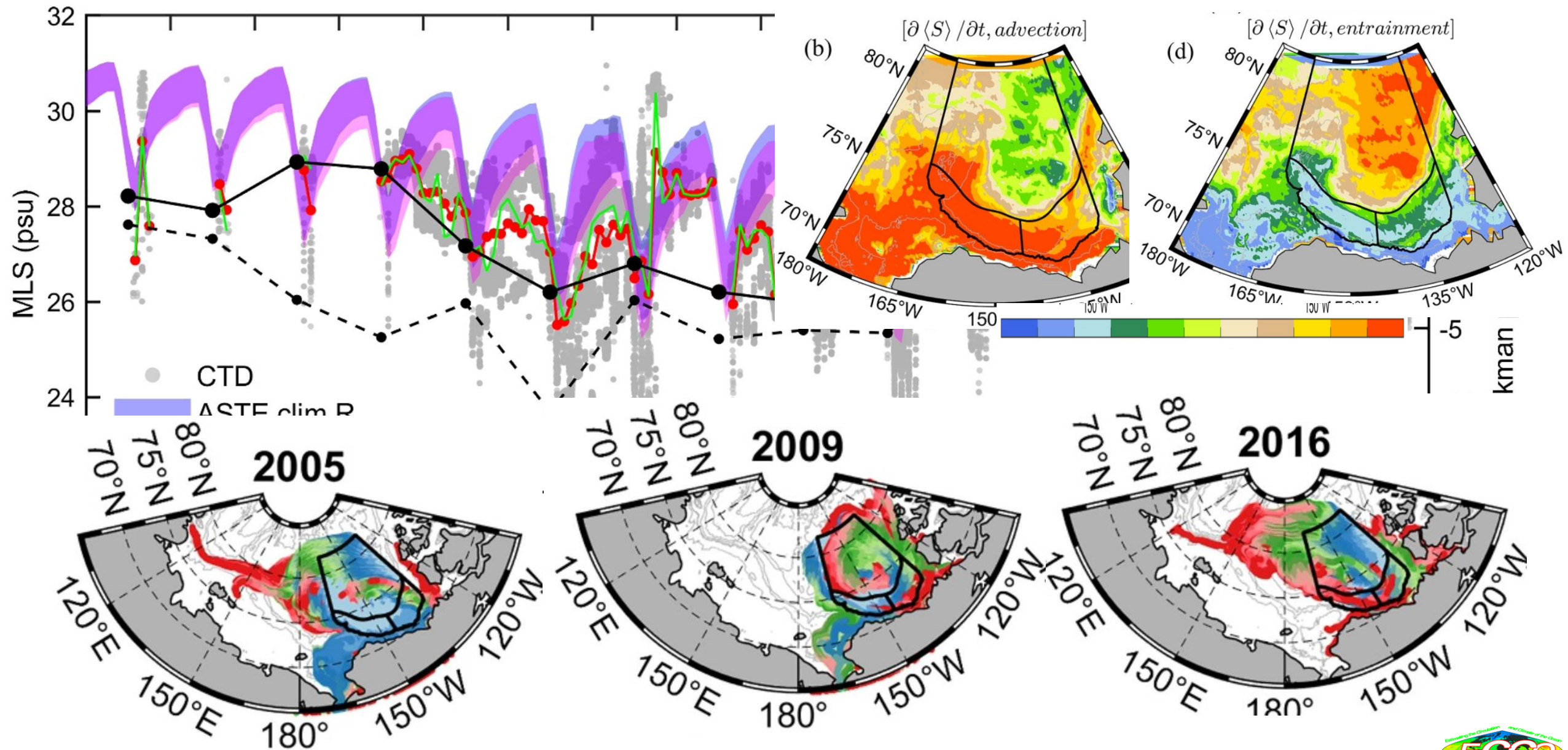
Improving ASTE: ML salinity budget in Beaufort Sea

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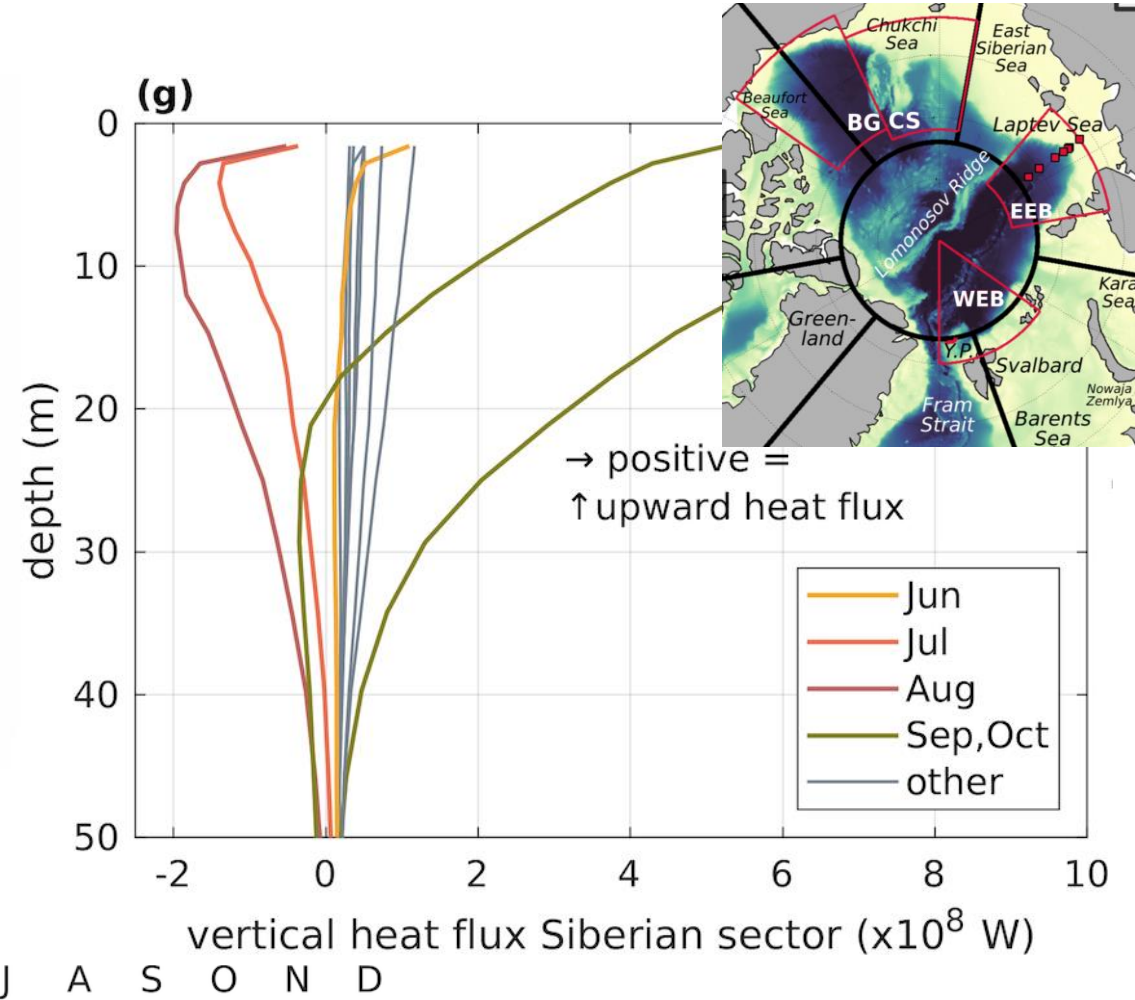
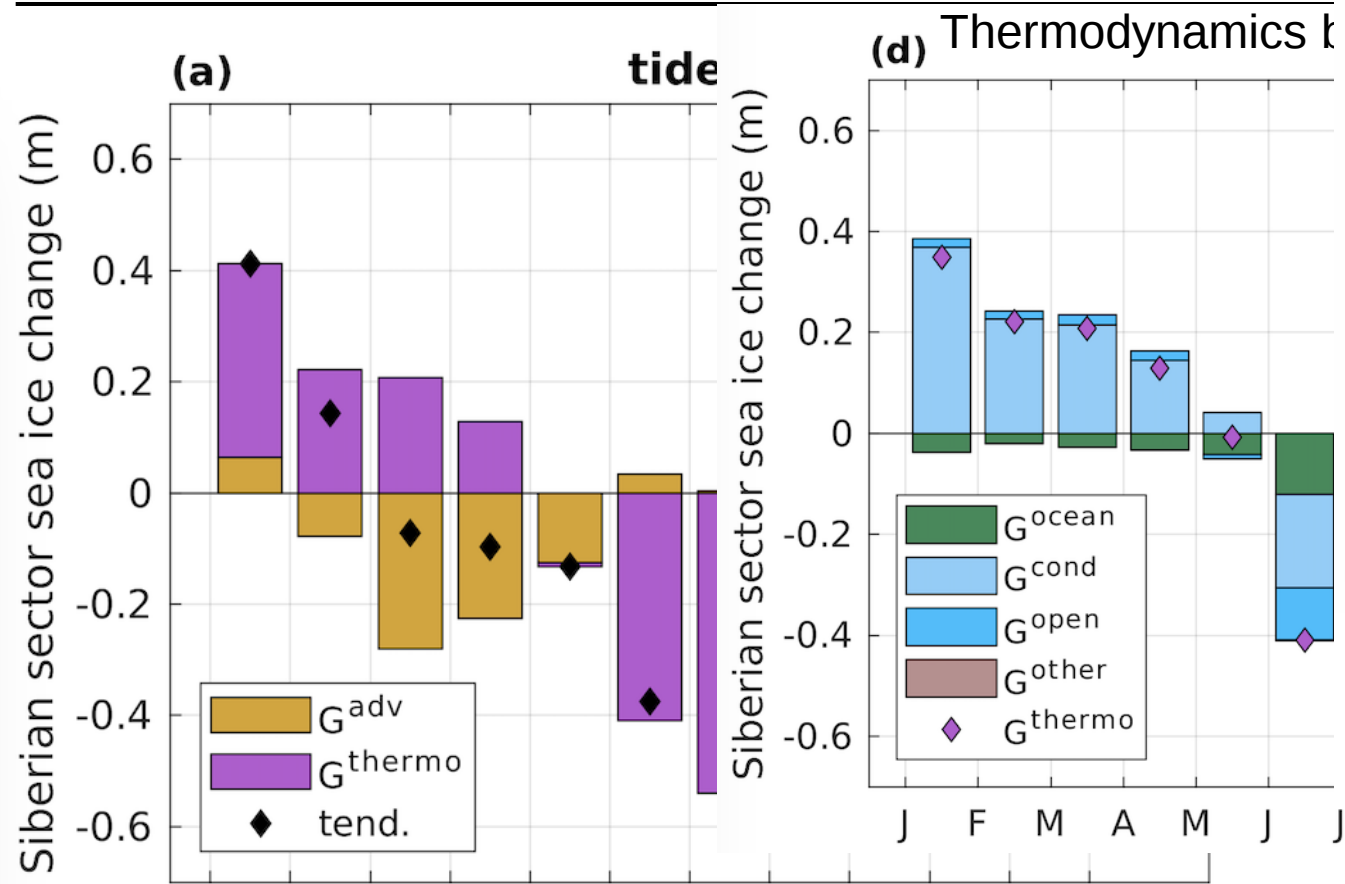


Improving ASTE: ML salinity budget in Beaufort Sea

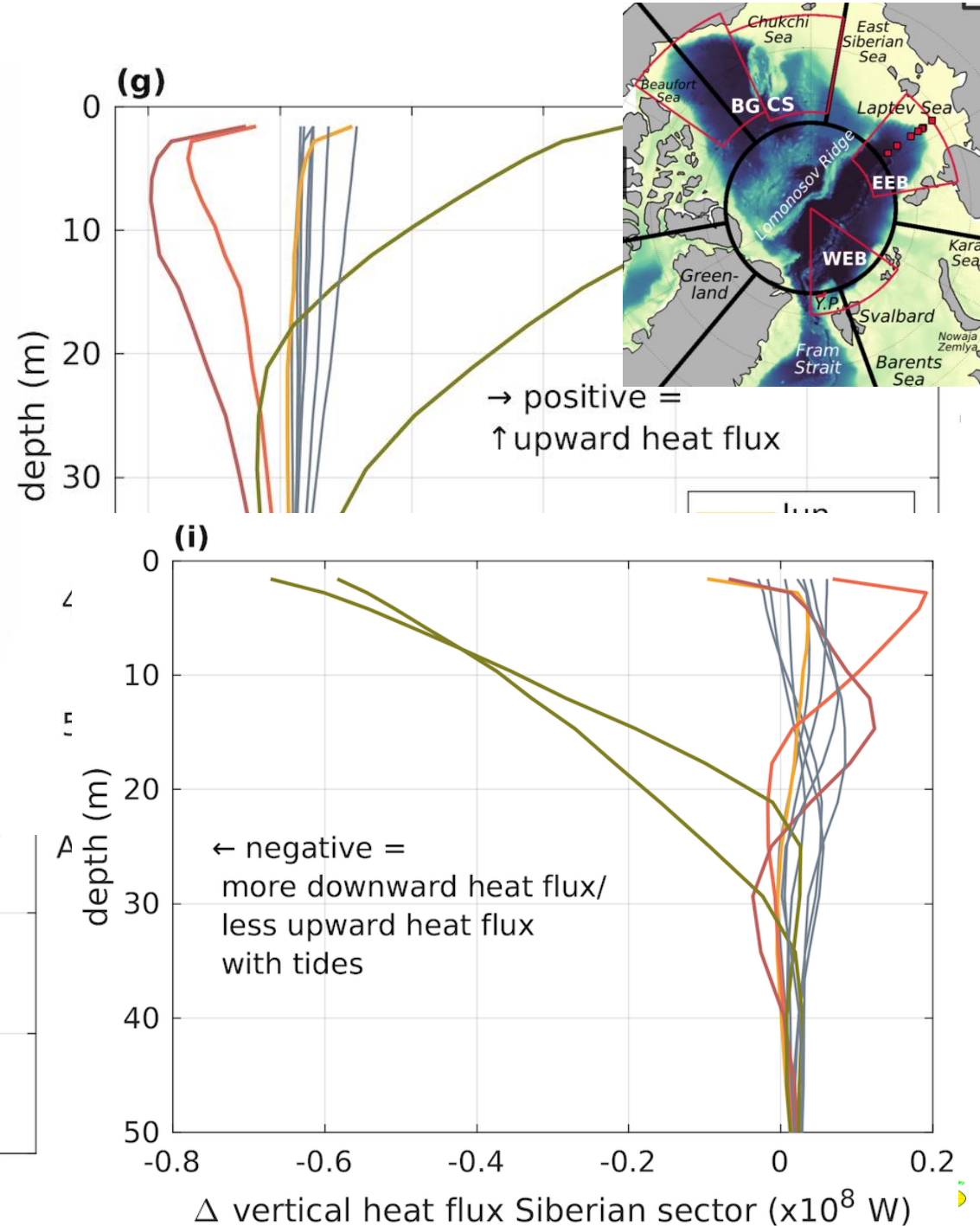
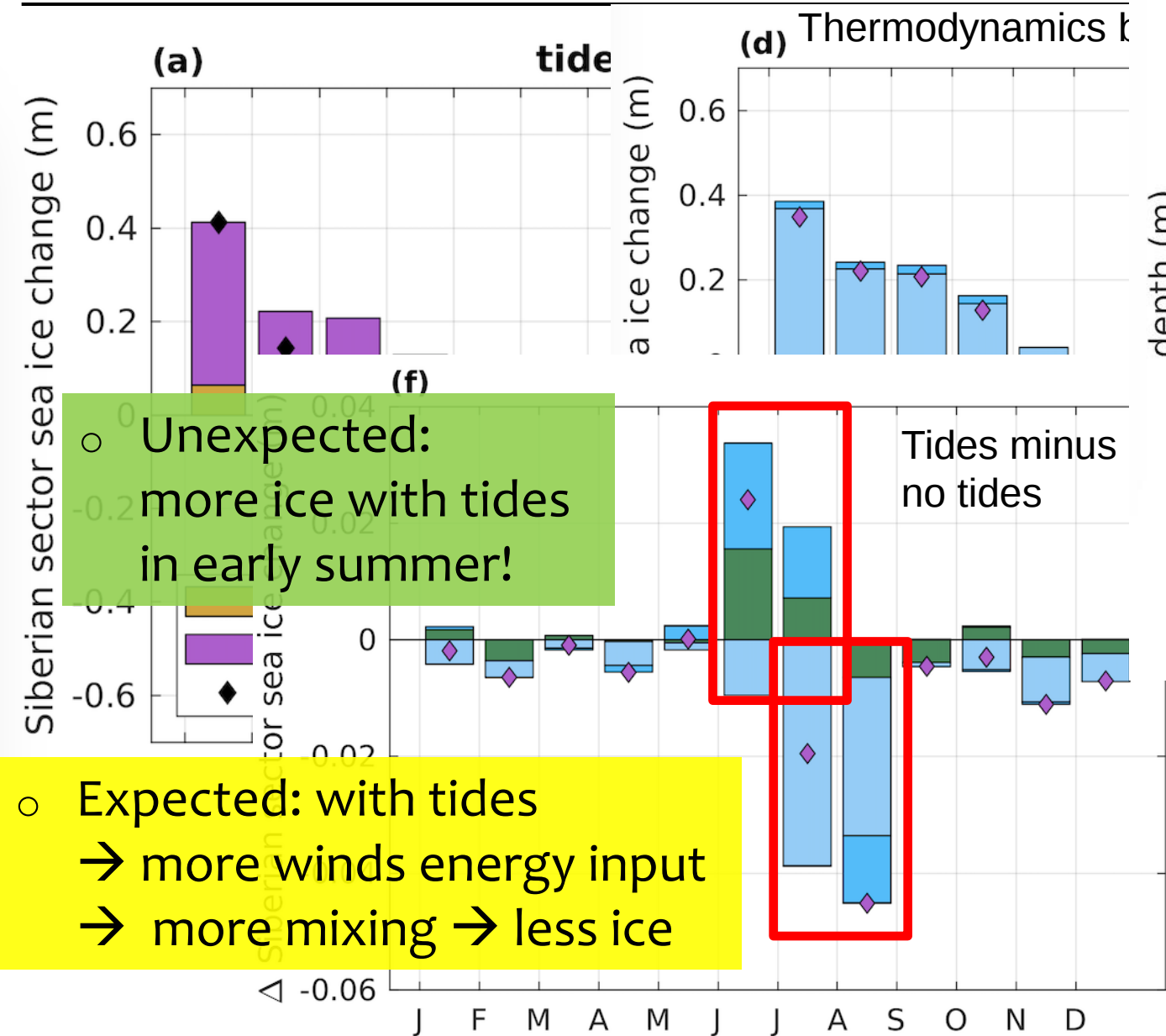
Zhong et al., 2024



Improving ASTE: Tides and sea ice Schulz et



Improving ASTE: Tides and sea ice Schulz et

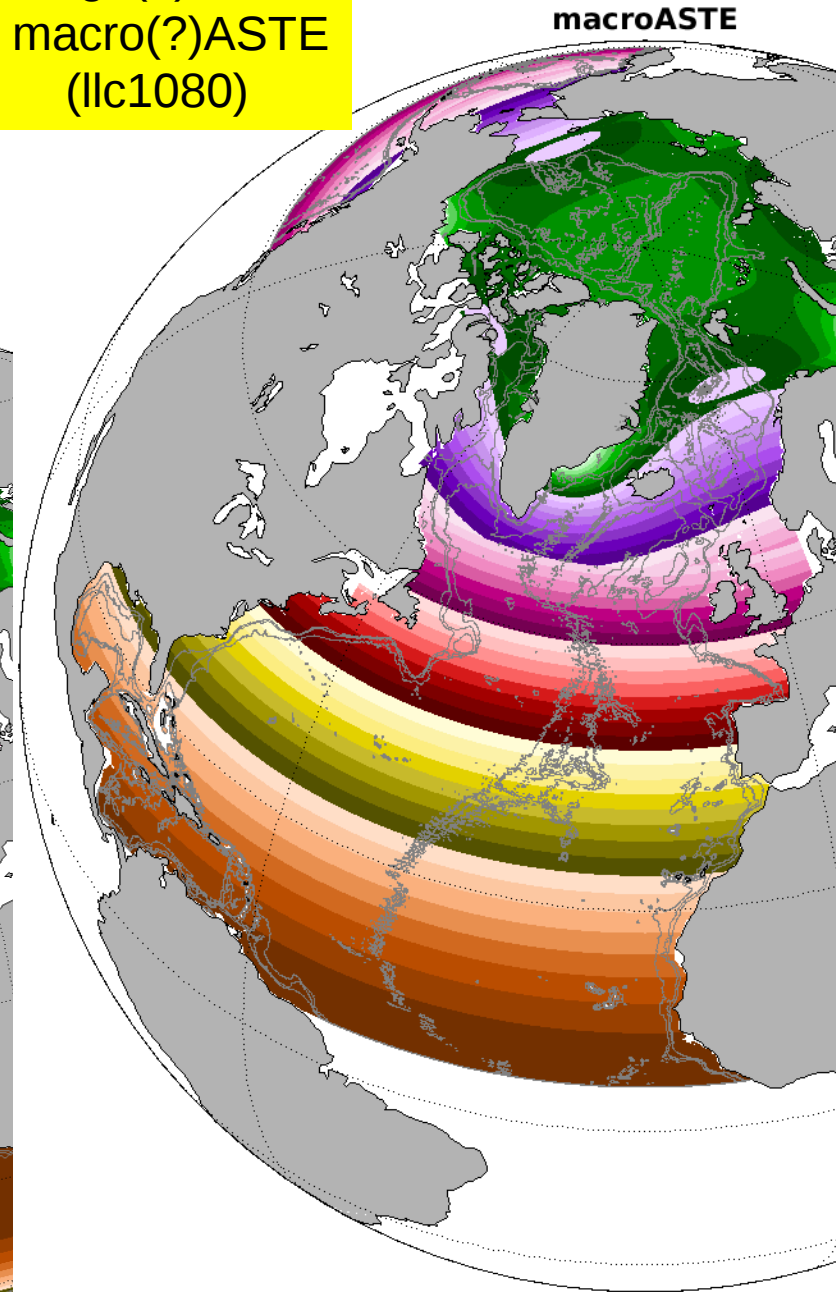
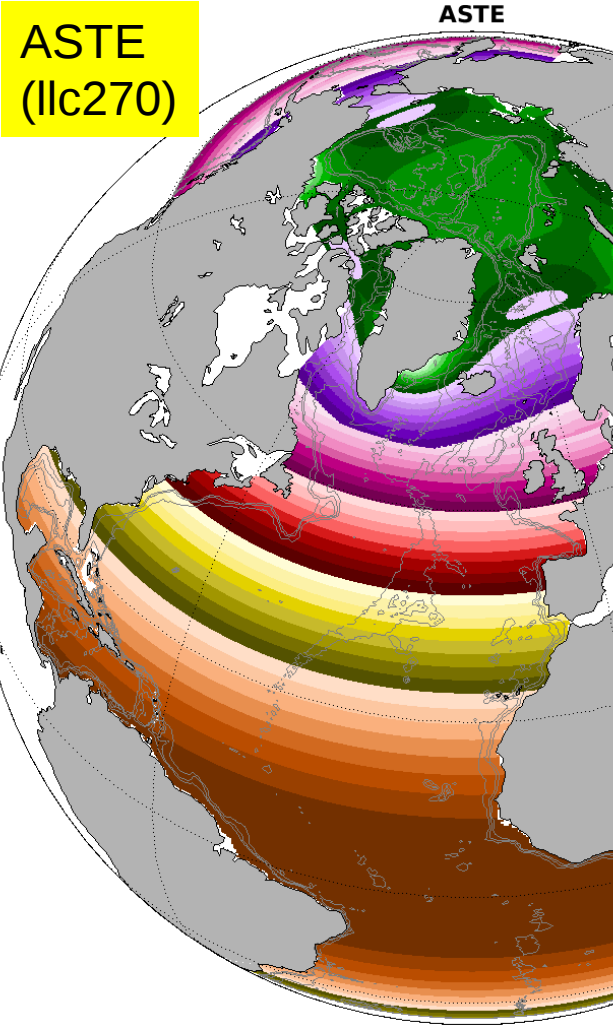


ASTE Extension (to end of 2023)

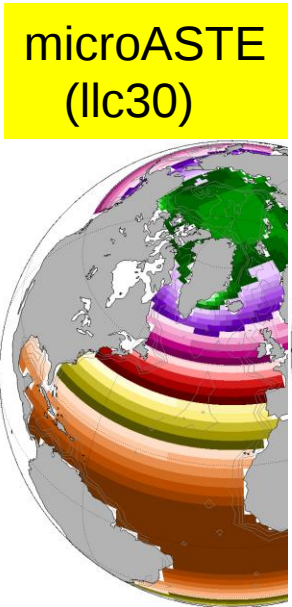
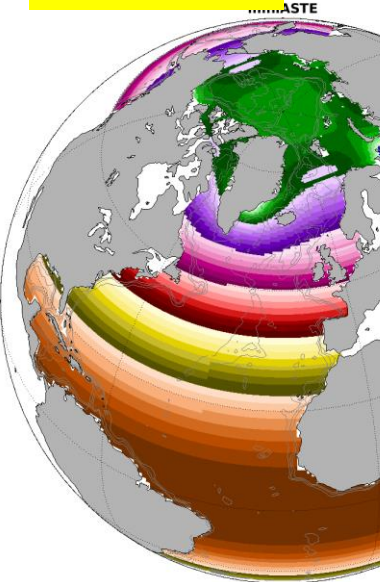
- Inputs update
 - Forcing
 - River runoff
 - Mixing maps
- Data updates
- Toward highres
- WMT

mega(?)ASTE
macro(?)ASTE
(llc1080)

jumbo(?)ASTE



miniASTE
(llc90)



170,60,16,3.5

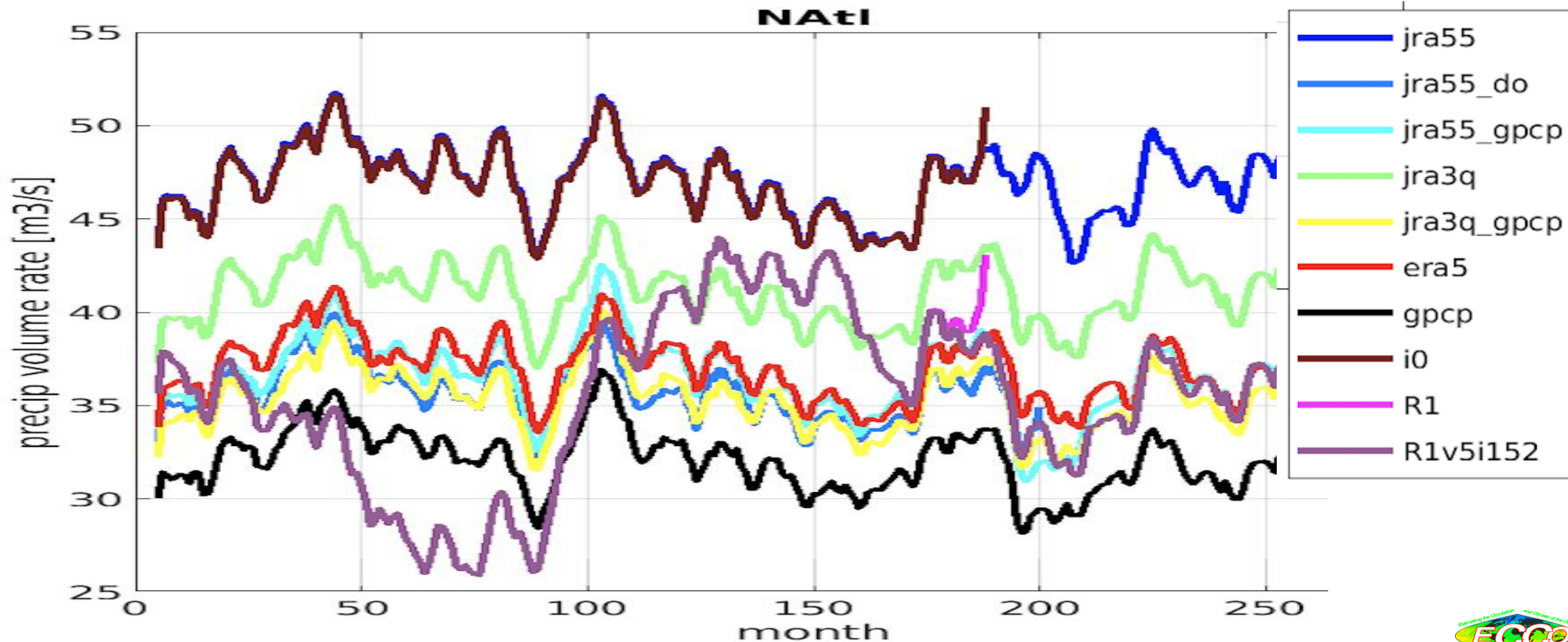
270,90,30,7.5

360,130,40,10

450,150,50,12

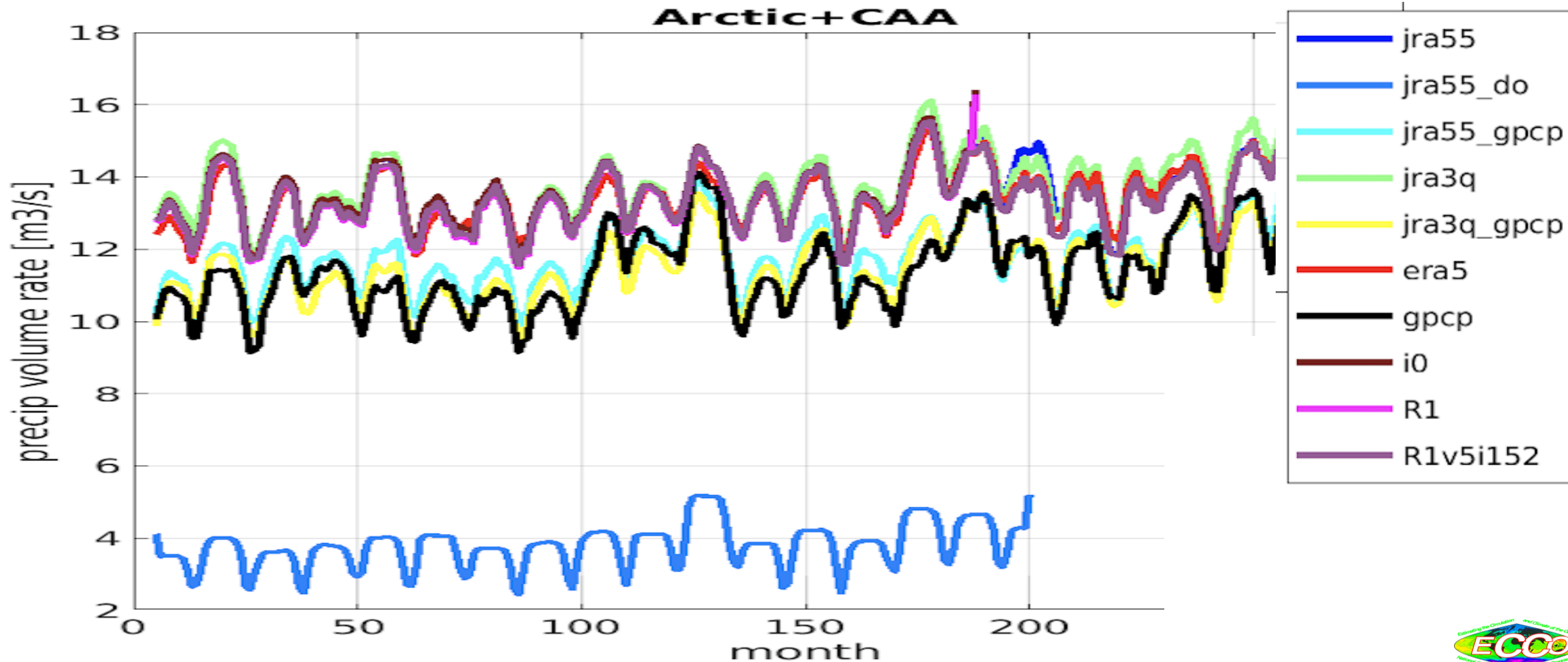
ASTE Extension

- to end of 2023
- Update to forcing: JRA-3Q [https://jra.kishou.go.jp/JRA-3Q/index_en.html]



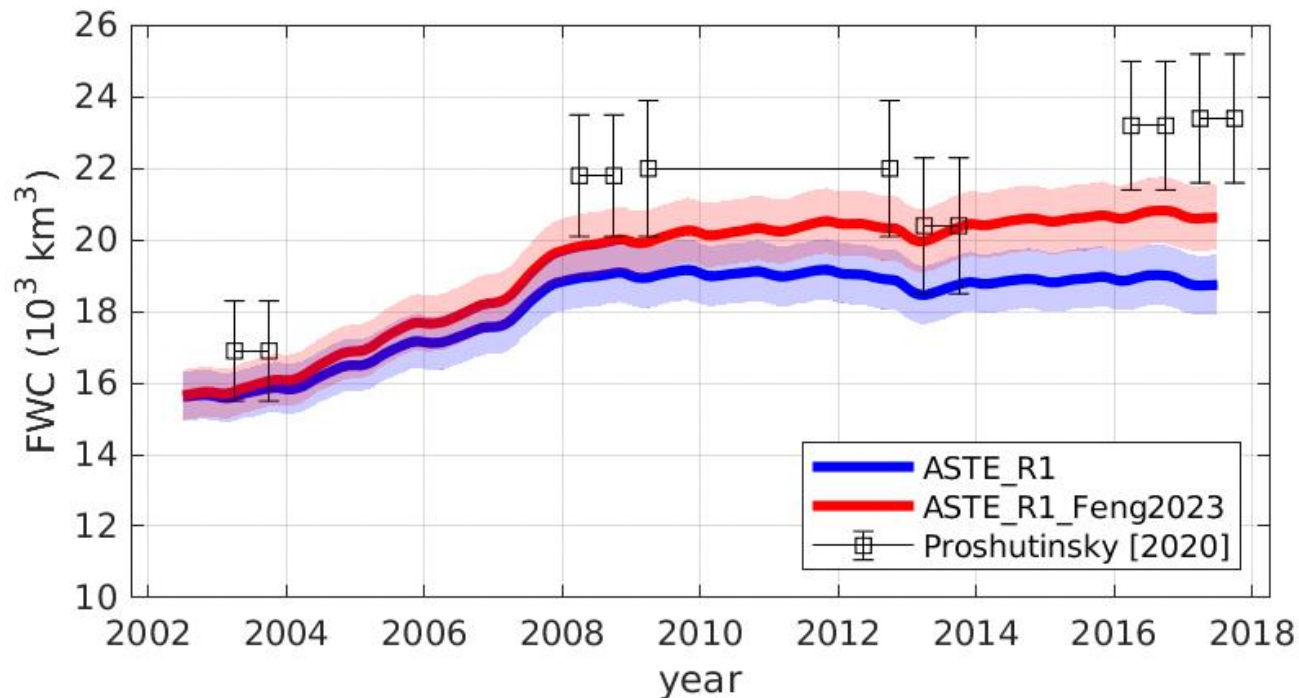
ASTE Extension

- to end of 2023
- Update to forcing: JRA-3Q [https://jra.kishou.go.jp/JRA-3Q/index_en.html]



ASTE Extension – river runoff (with help from Wenli Zhong)

- Updating river runoff: [Feng et al. 2021]
- pan-Arctic Remotely-sensed Arctic Discharge Reanalysis (RADR) for 1984-2018
- Assimilated product, of 9.18mil discharge obs from 227 mil river width measurements from Landsat images



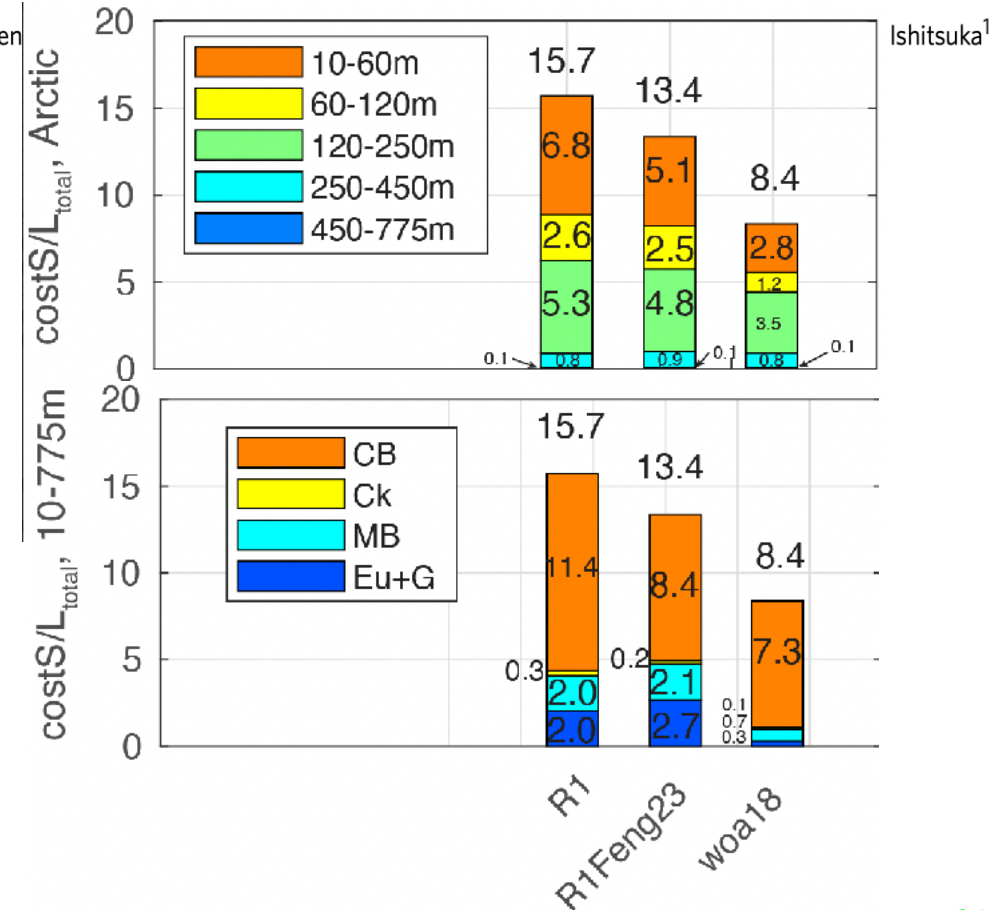
ARTICLE

<https://doi.org/10.1038/s41467-021-27228-1>

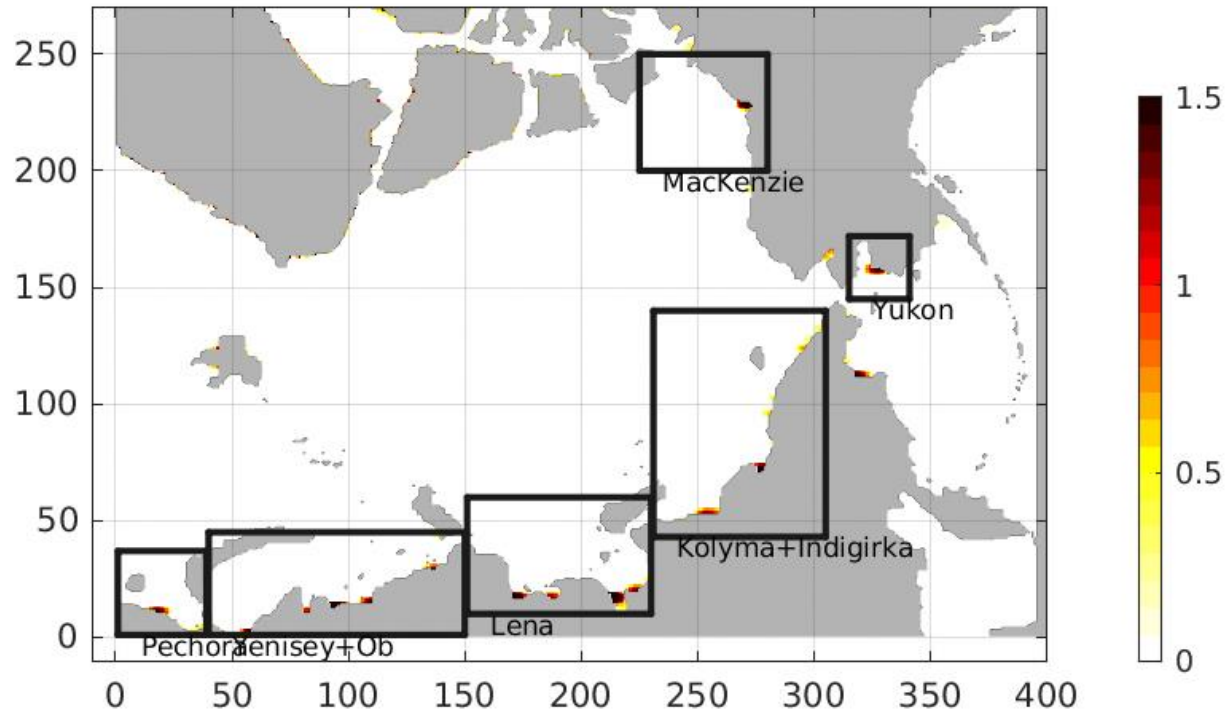
OPEN

Recent changes to Arctic river discharge

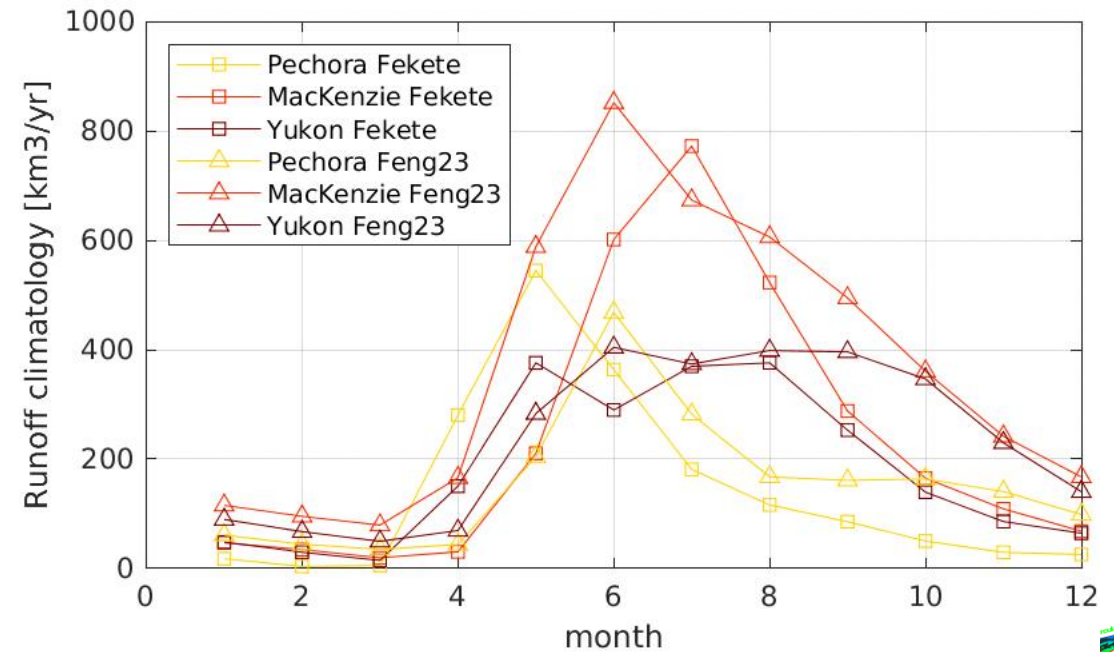
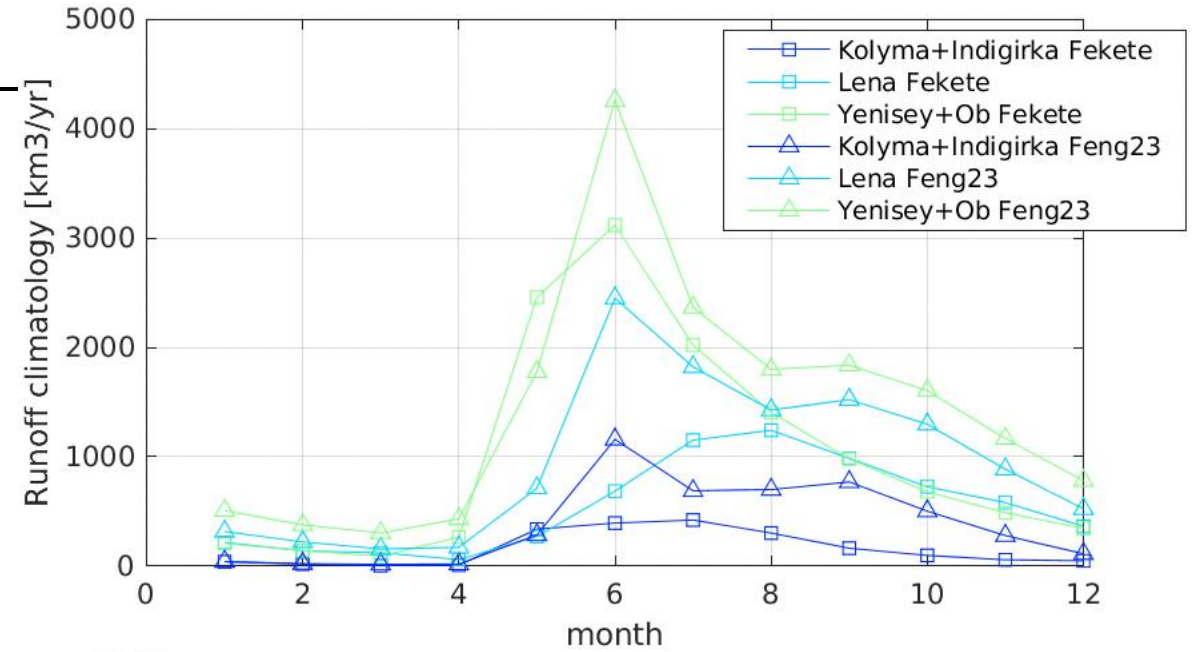
Dongmei Feng



ASTE Extension – river runoff



Note: change in seasonal cycle



ASTE Extension – data sets

Thanks Ian for satellite
obs, Matt for Argo

- MOSAiC
- Mooring arrays across GSR: <http://www.oceansites.org/tma/gsr.html>
- A-Twain data: [Våge16,PerezHernandez17,Renner18]
<https://data.npolar.no/dataset?page=1&q=A%20Twain&location=arcticOcea>
<https://openpolar.no/Record/npolardata:oai:npolar.no:dataset%2Fd3a5adc2-aa97-4934-893d-83427f7134ad>
- Davis Strait data: not avail after 2015, already sent email to C. Lee to ask.
- GIS ridge: <http://kogur.whoiedu/php/index.php>, http://kogur.whoiedu/Historical_Data/ [Våge13, 18]
- Iceland Greenland Project (IGP, Huang21): <https://web.whoiedu/all0118/> , (only ctd, not underway vel)
- UDASH: [Behrendt17, Behrendt18], 1980--2015, <https://doi.pangaea.de/10.1594/PANGAEA.872931>
- International Council for the Exploration of the Seas (ICES): <https://data.ices.dk/>
- Faroe-Island data: <http://www.envofar.fo>
- GreenlandScotlandRidge: <http://www.oceansites.org/tma/gsr.html> (transp avail, contact people for indiv)
- FramStrait [in/out]flow (start here, already downloaded publication from year 2023)
<https://www.pangaea.de/?q=von+Appen%2C+Wilken-Jon&f.pubyear%5B%5D=2022&f.project%5B%5D=FRAM>
<https://www.npolar.no/en/projects/fram-strait-arctic-outflow-observatory/#toggle-project-dataset;>
<https://data.npolar.no/dataset/c4d80b64-25f6-4afd-b392-696430c3fd14>
- NABOS (Igor Polyakov)

ASTE Jupyter notebook

https://github.com/crios-ut/aste/blob/master/aste_llcreader_example.ipynb

aste / aste_llcreader_example.ipynb

↑ Top

Preview

Code

Blame

1088 lines (1088 loc) · 787 KB

Raw



<https://doi.org/10.1145/3311790.3396656>

```
In [1]: import numpy as np
import warnings
import matplotlib.pyplot as plt
import xarray as xr
import cmocean
from dask.distributed import Client

from xmitgcm import llcreader
import ecco_v4_py
```

Get an xarray dataset with *all* ASTE_R1 variables, depth levels, and time steps

The function `get_dataset` by default grabs all available output, at all depth levels and all time steps, where each time step represents a monthly mean for that field. This may be suboptimal when operating on a machine with limited memory, for instance on a laptop. See the [llcreader documentation](#) for more examples on how to subset the data with the [get_dataset method](#), including how to grab specific variables, vertical slices, or time steps.

```
In [2]: aste = llcreader.CRIOSPportalASTE270Model()
```

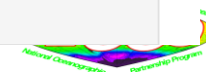
```
In [3]: ds = aste.get_dataset()
```

Grab a single monthly average

Here we subset the dataset to show the ASTE ocean state during a single month, September 2012. Alternatively, one can provide the corresponding iteration to the `iters` option in `get_dataset` to achieve the same behavior.

```
In [4]: ds = ds.sel(time='2012-09')
```

File display



Questions?

ASTE Extension / Devel

- Higher res: Ilc1080-based ASTE (exists in forward mode)
 - Multigrid
 - following ECCO efforts, also testing tape saving capability (Matt Mazloff's code)
- Adjusted parameters:
 - (initial velocities)
 - (open boundaries)
 - (Ice thickness & area ICs)
 - (time-mean spatially-varying ice drags?)

ASTE Extension (to end of 2023)

- Inputs update
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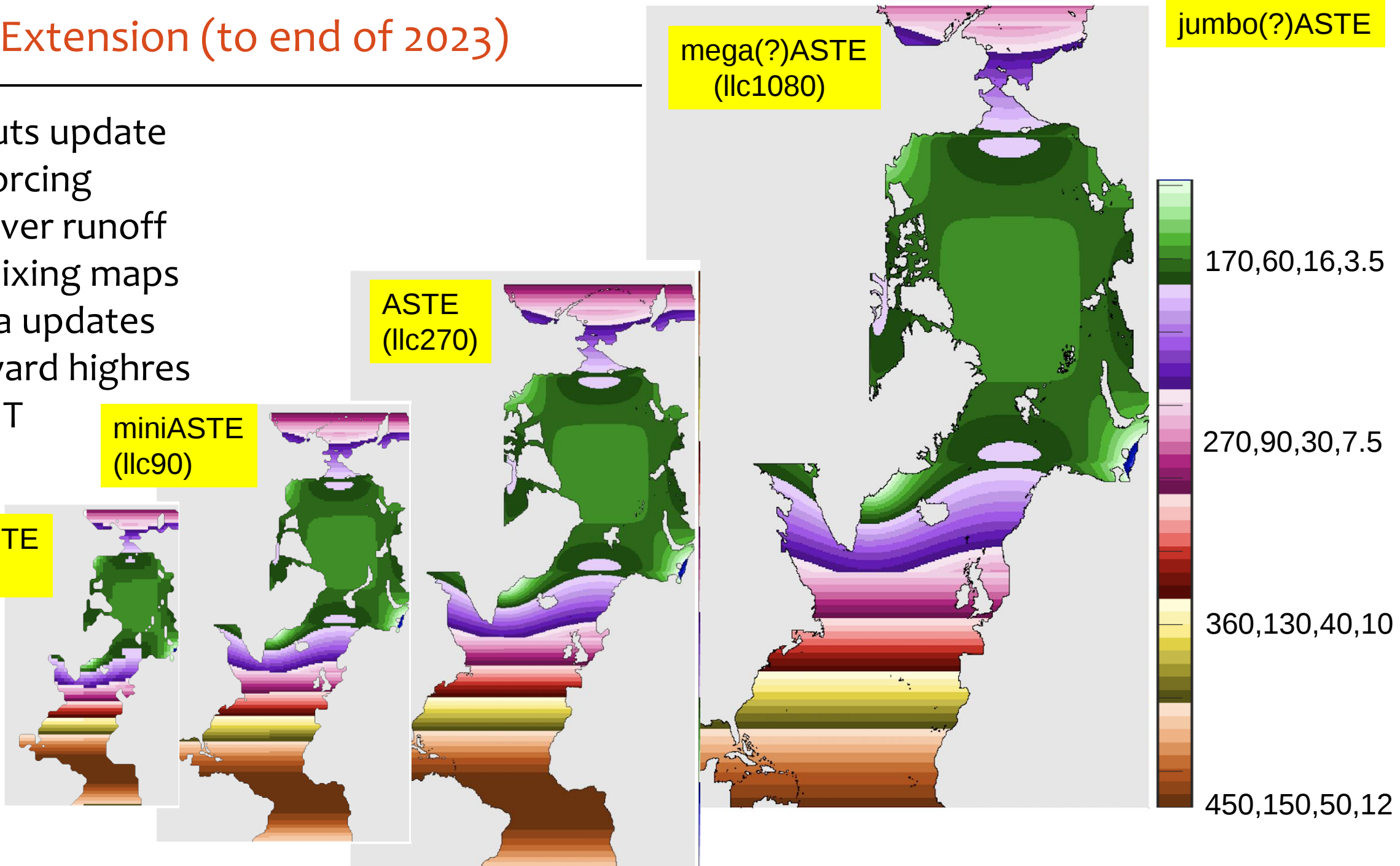
microASTE
(llc30)

miniASTE
(llc90)

ASTE
(llc270)

mega(?)ASTE
(llc1080)

jumbo(?)ASTE



From idealized / theoretical: constraint on circulation

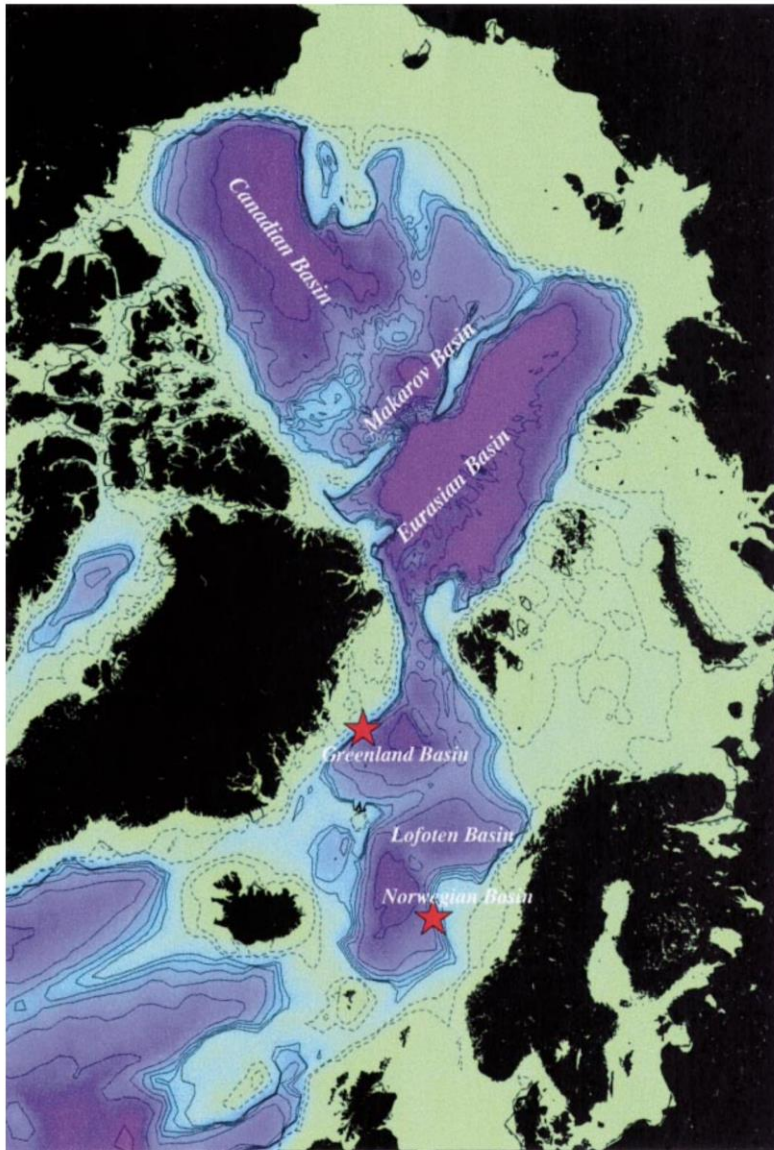


FIG. 3. The f/H field (lines) and bathymetric field (shades) for the Nordic seas and Arctic Ocean. The location of two moorings discussed later are shown with stars. Note the f/H contours approximately follow the isobaths.

- Viscous dissipation (circulation) -- uncertain
- 3-D internal turbulence mixing – uncertain
- Unresolve mesoscale eddies (Redi) -- uncertain

$$\frac{\partial}{\partial t} \mathbf{u} + \mathbf{k} \times f\mathbf{u} = -g\nabla\eta + \frac{\boldsymbol{\tau}}{\rho H} - \frac{R\mathbf{u}}{H}$$

$$\frac{\partial}{\partial t} \oint \mathbf{u}_0 \cdot d\mathbf{l} = \oint \frac{\boldsymbol{\tau}}{H} \cdot d\mathbf{l} - \oint \frac{\mathbf{u}_0}{H} \cdot d\mathbf{l}. \quad (7)$$

The terms on the rhs of (7) represent the transports in the surface and bottom Ekman layers; any imbalance between them alters the circulation in the bounded region.

western moorings (numbers 410–413), the principal components of the currents align closely with the local tangent of the f/H contours. This component explains more than 85% of the total variance when currents are

Isachsen et al., JPO 2003