

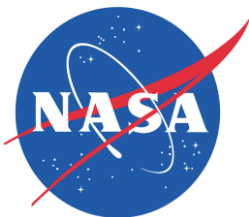


Ocean carbon and biogeochemistry

Galen A. McKinley

Professor of Earth and Environmental Sciences
Columbia University and Lamont-Doherty Earth Observatory

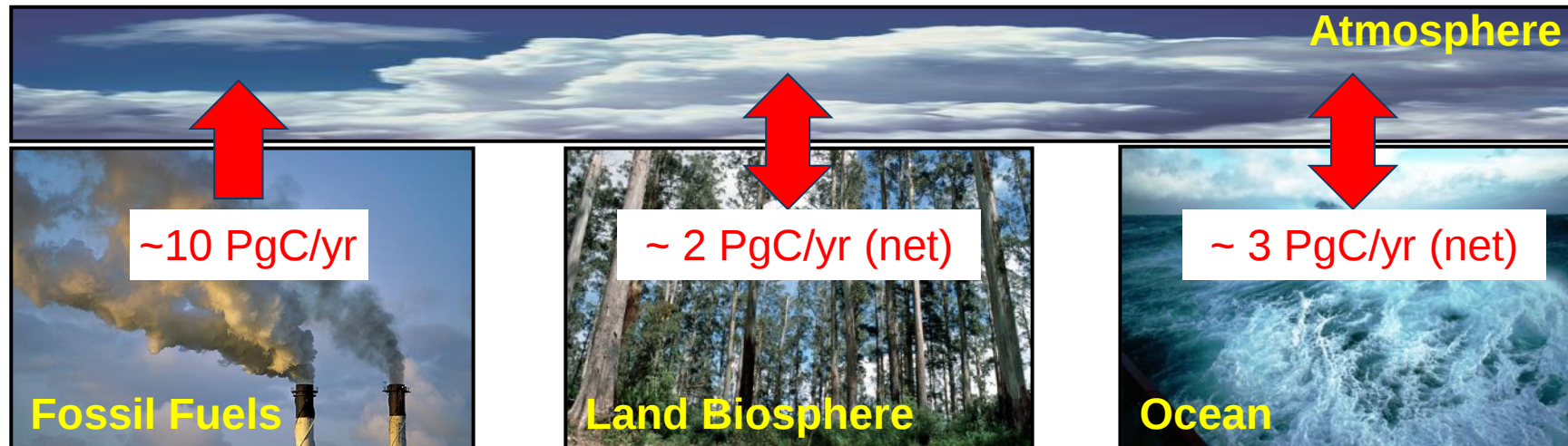
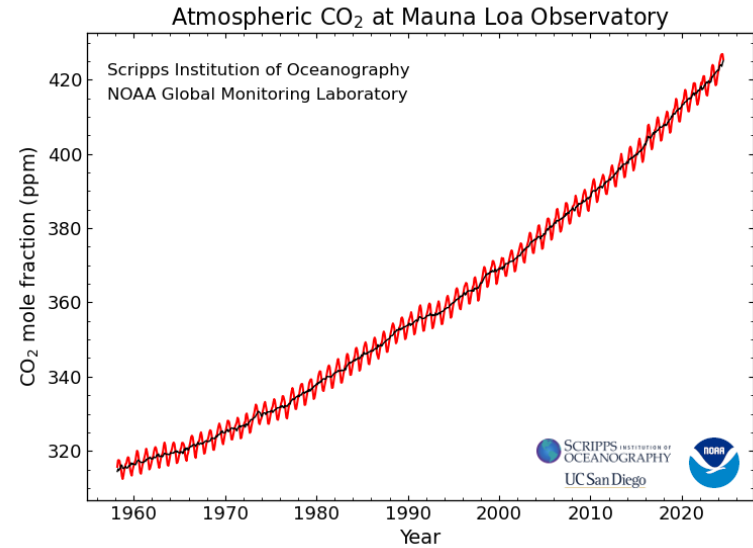
27 May 2025



LEAP

Atmospheric CO₂ is rising due to human emissions

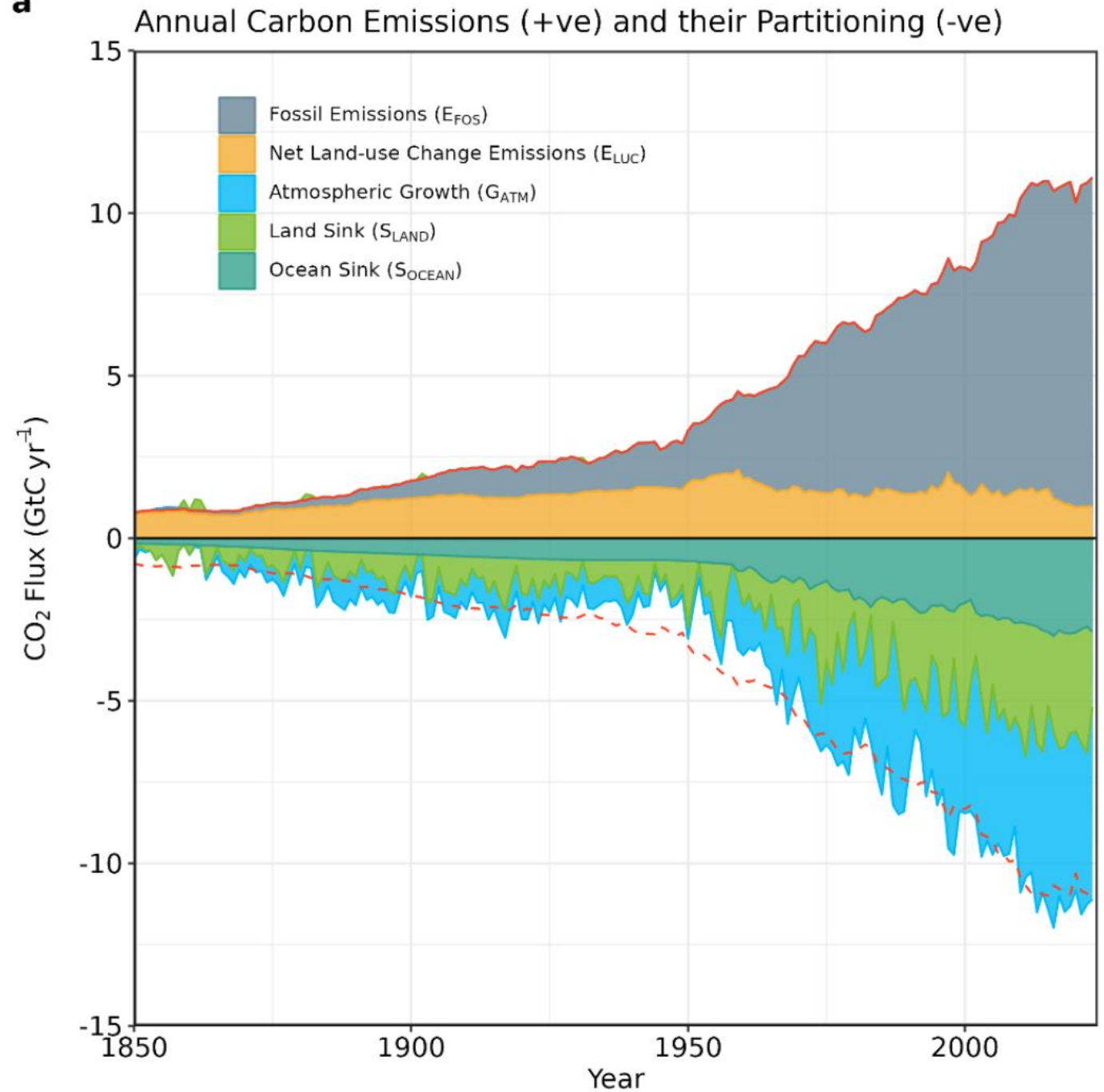
Ocean and terrestrial biosphere absorb ~50% of emissions



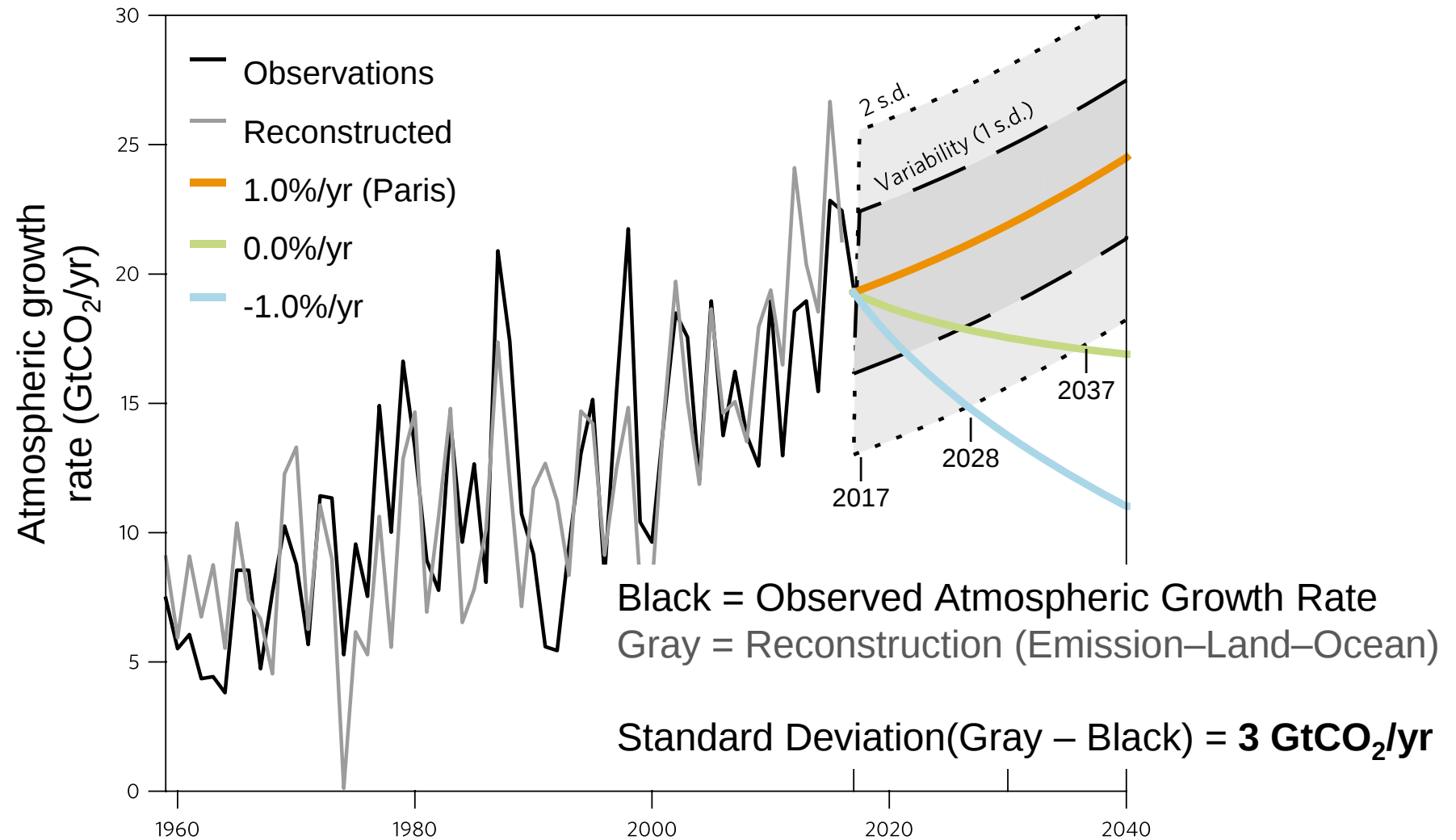
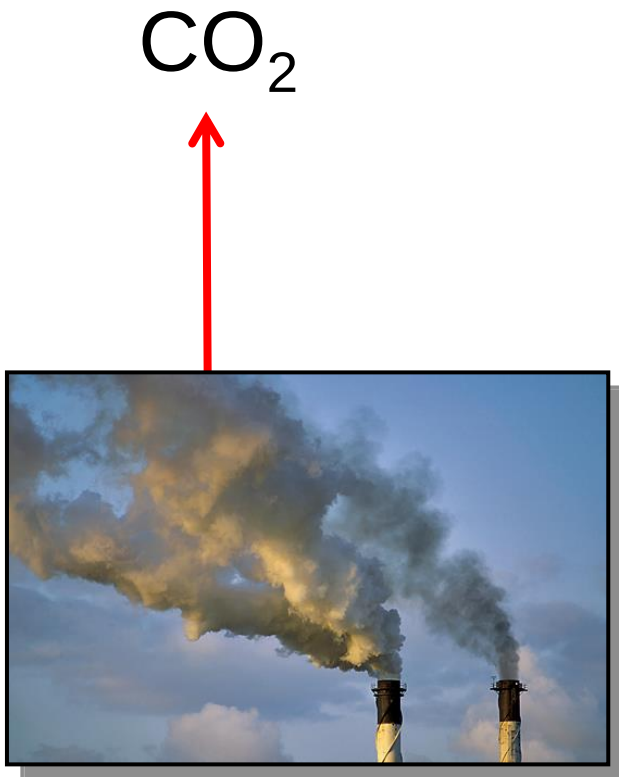
Global Carbon Budget

Friedlingstein et al. 2025, ESSD

a



Reduced uncertainty in ocean and land sinks required to usefully detect emissions mitigation



Peters et al. 2017 Nature Climate Change

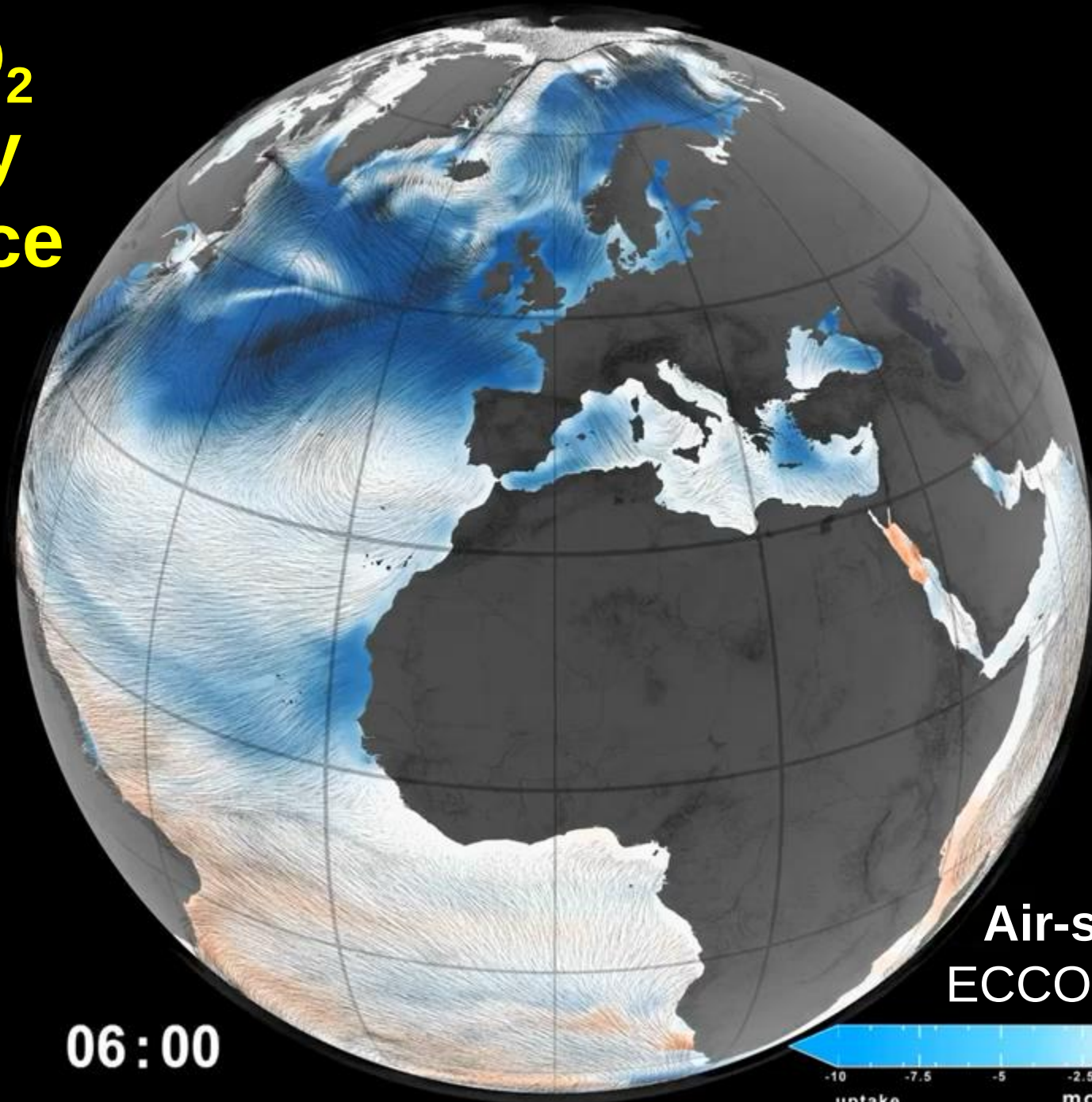
Outline

- Key processes of the ocean carbon cycle
- Improving quantification of the ocean carbon sink
 - *Models*: evaluate ML-based data product skill, given sampling
 - *pCO₂ data products*: apply to identify model mean-state biases
 - *Models and data*: combine using ML
 - *marine Carbon Dioxide Removal (mCDR)* – quantifying “additionality”?

Outline

- **Key processes of the ocean carbon cycle**
- Improving quantification of the ocean carbon sink
 - *Models*: evaluate ML-based data product skill, given sampling
 - *pCO₂ data products*: apply to identify model mean-state biases
 - *Models and data*: combine using ML
 - *marine Carbon Dioxide Removal (mCDR)* – quantifying “additionality”?

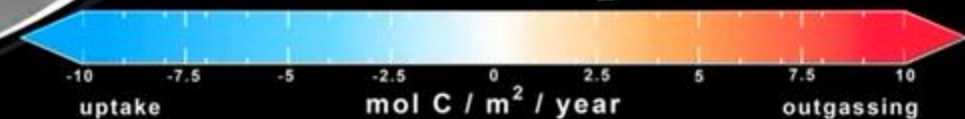
Air-sea CO₂ fluxes vary across space and time

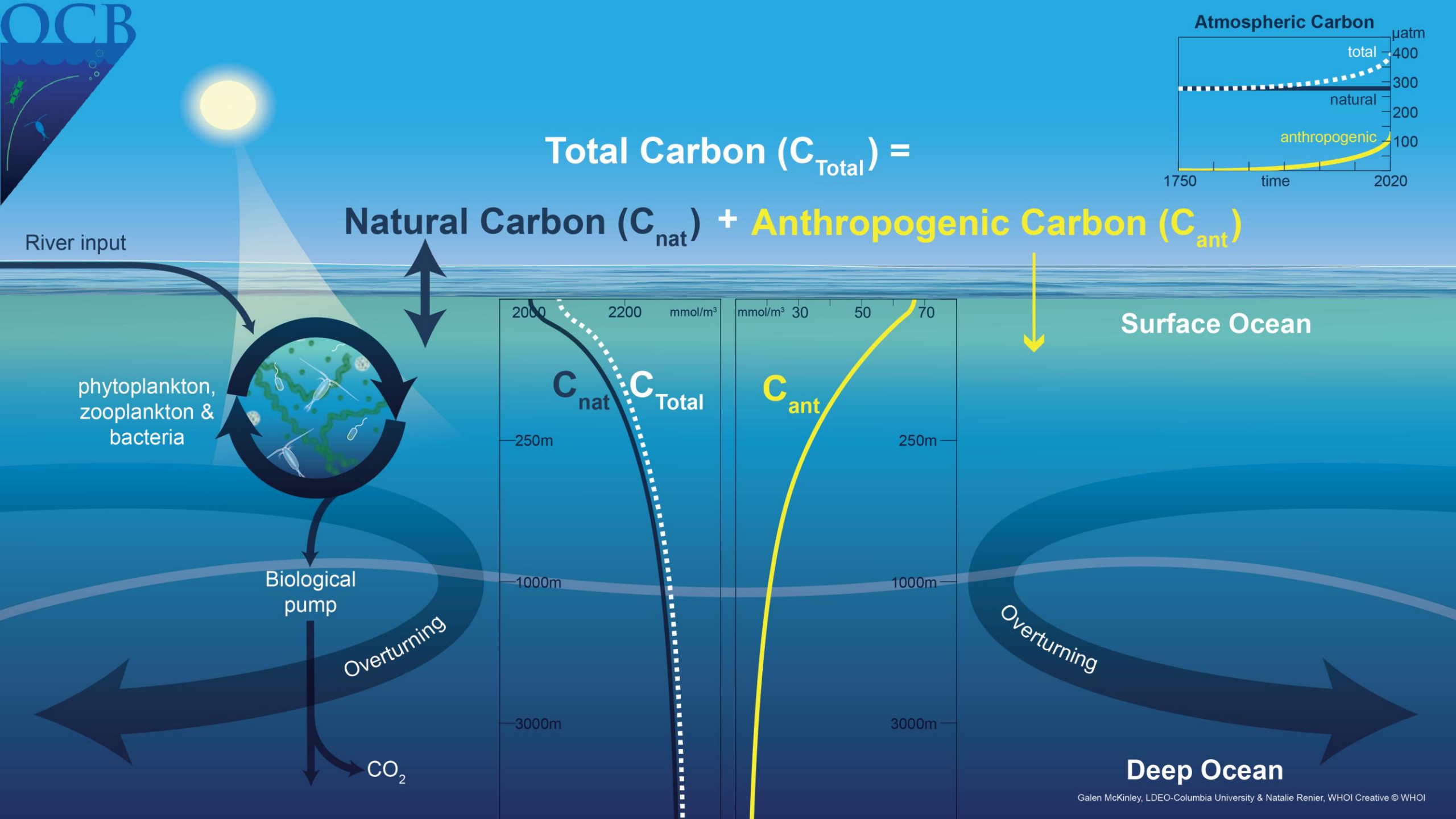


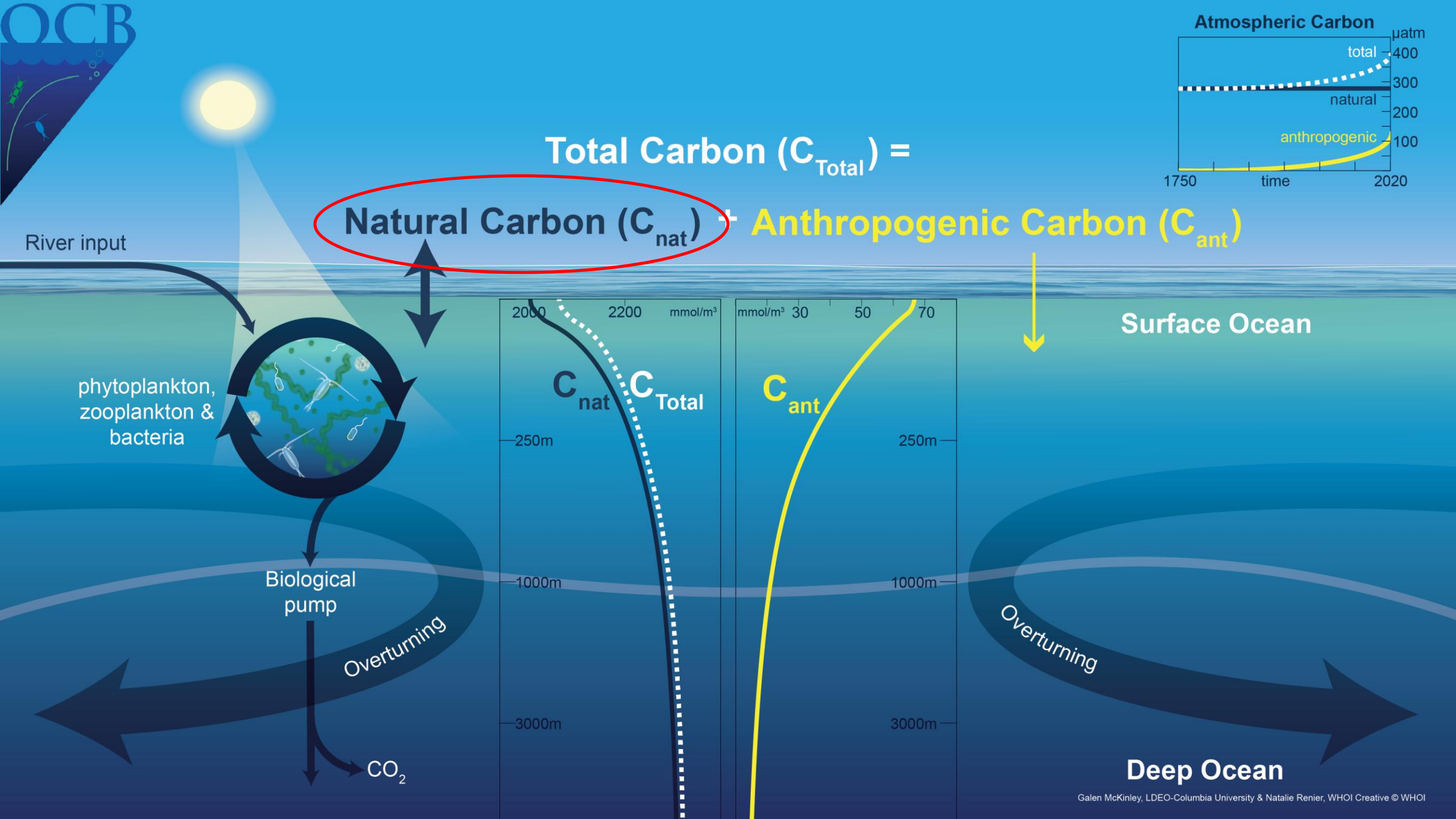
Carroll et al. 2020, 2022

Air-sea CO₂ Flux
ECCO-Darwin Model

03 Jan 2012 06:00

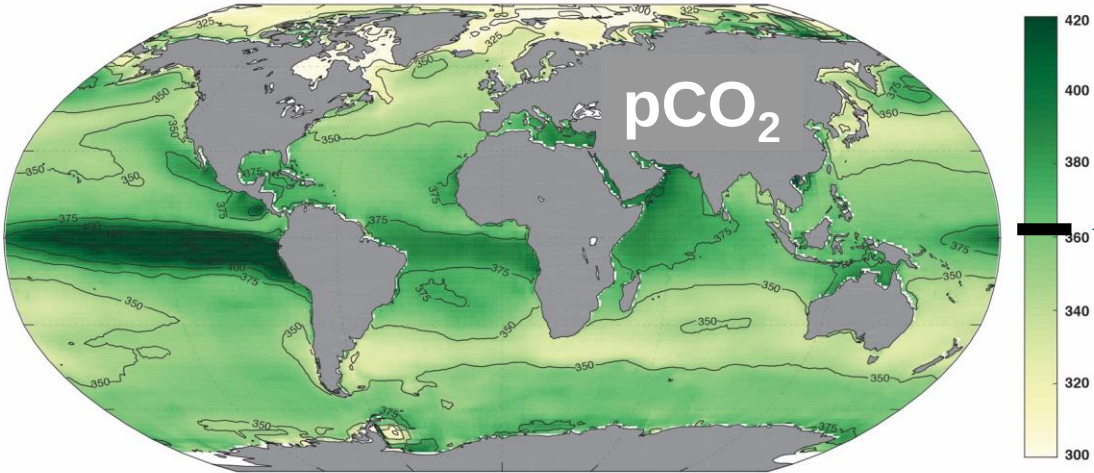




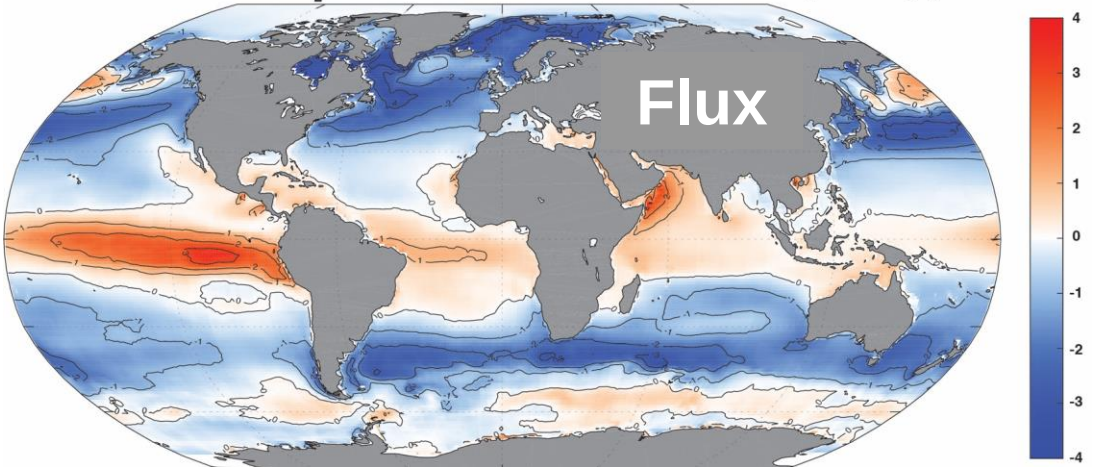


Surface ocean $p\text{CO}_2$ is primary control on air-sea CO_2 flux

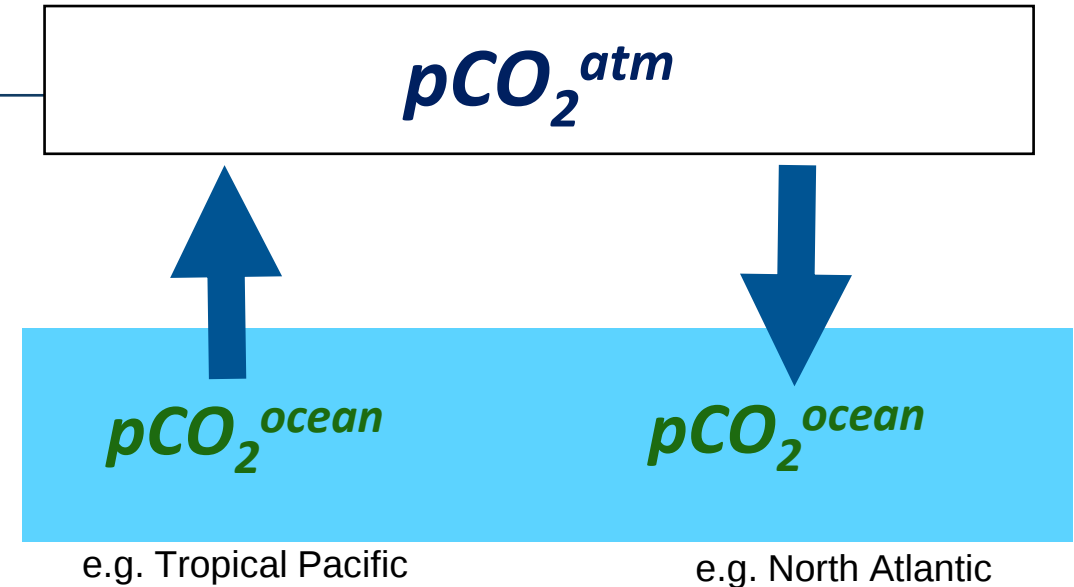
1988-2018 mean $p\text{CO}_2$ from observation-based products (μatm)



1988-2018 mean CO_2 flux from observation-based products ($\text{mol}/\text{m}^2/\text{yr}$)



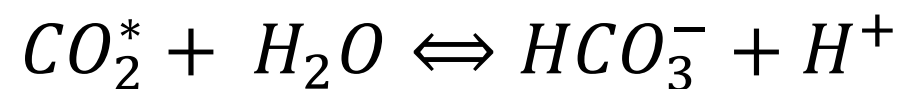
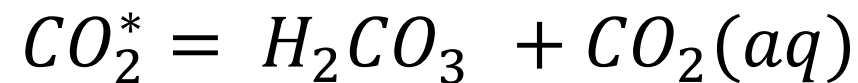
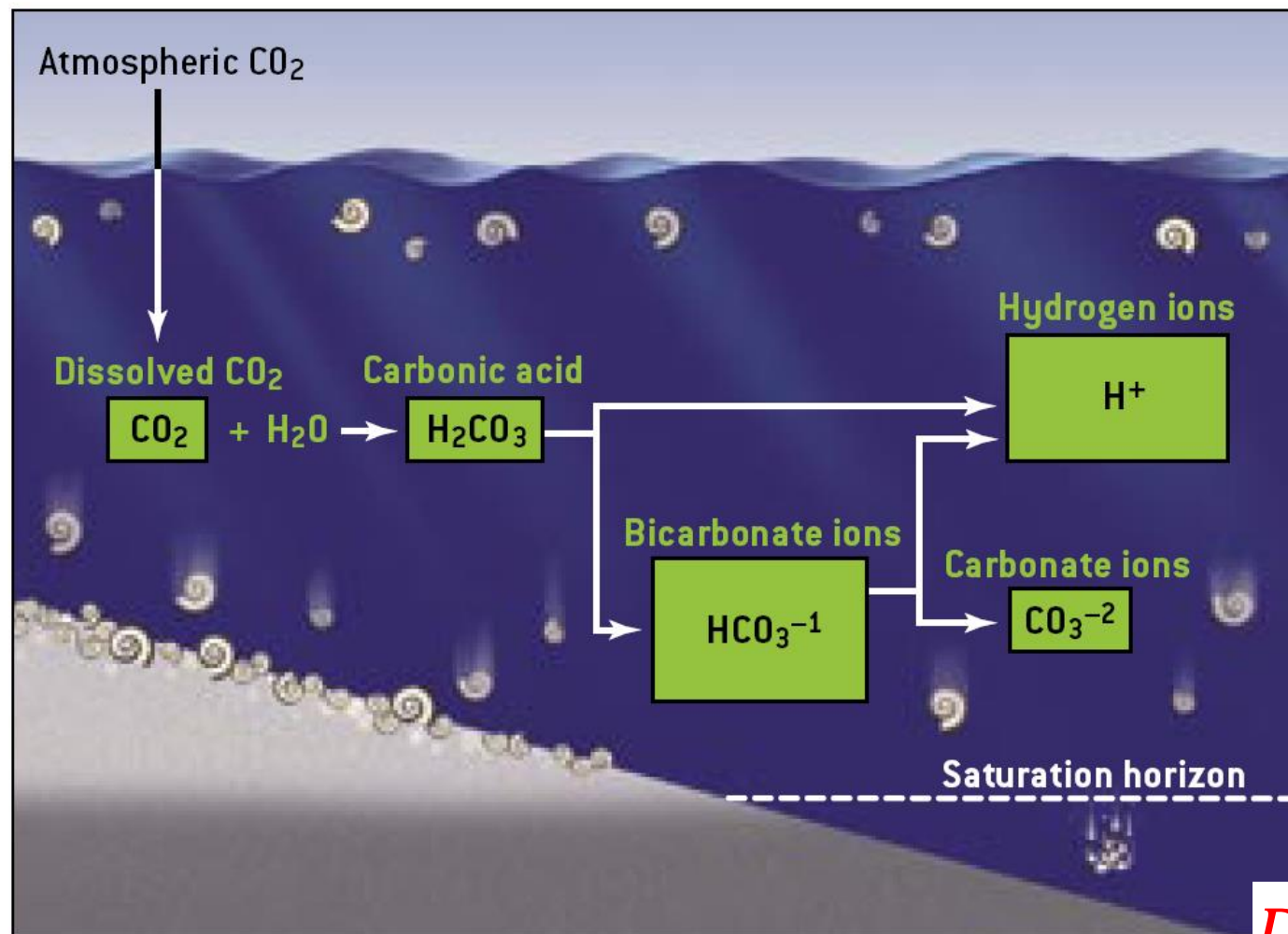
If $p\text{CO}_2^{\text{ocean}} > p\text{CO}_2^{\text{atm}}$ flux is out of ocean (positive)
 If $p\text{CO}_2^{\text{ocean}} < p\text{CO}_2^{\text{atm}}$ flux is into ocean (negative)



$$\text{Air sea } \text{CO}_2 \text{ flux} = k_w S_{\text{CO}_2} (1 - f_{\text{ice}}) (p\text{CO}_2^{\text{ocean}} - p\text{CO}_2^{\text{atm}})$$

Crisp et al., 2022, Review of Geophysics

Ocean Inorganic Carbon Cycle



$$K'_1 = \frac{[HCO_3^-][H^+]}{[CO_2^*]}$$



$$K'_2 = \frac{[CO_3^{2-}][H^+]}{[HCO_3^-]}$$

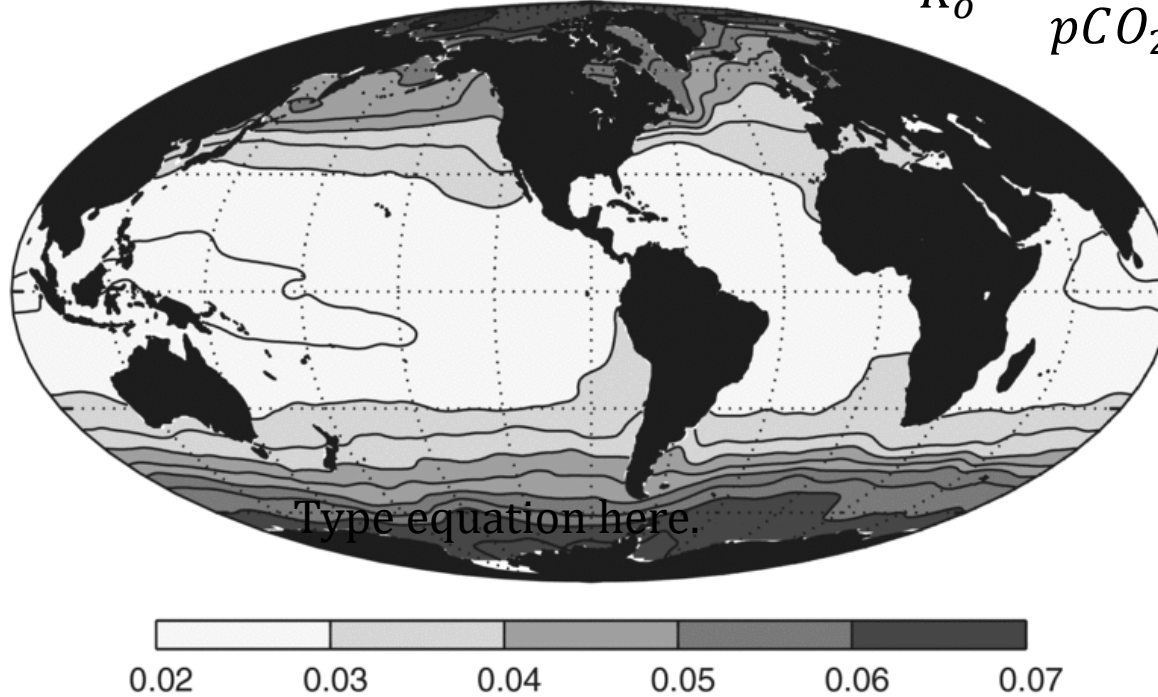


SOLUBILITY

$$pCO_2 = \frac{[CO_2^*]}{K_o}$$

(a) solubility ($\text{mol kg}^{-1} \text{ atm}^{-1}$)

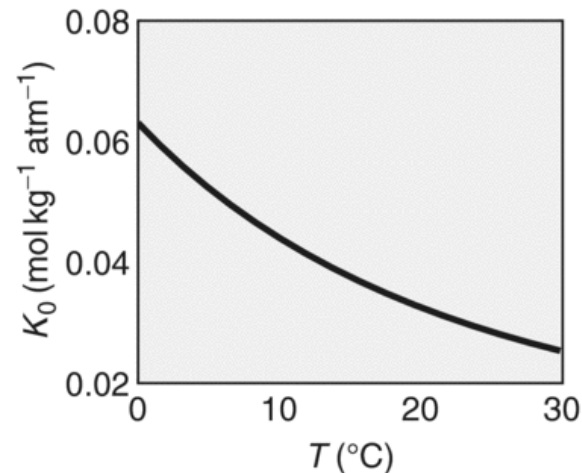
$$K_o = \frac{[CO_2^*]}{pCO_2}$$



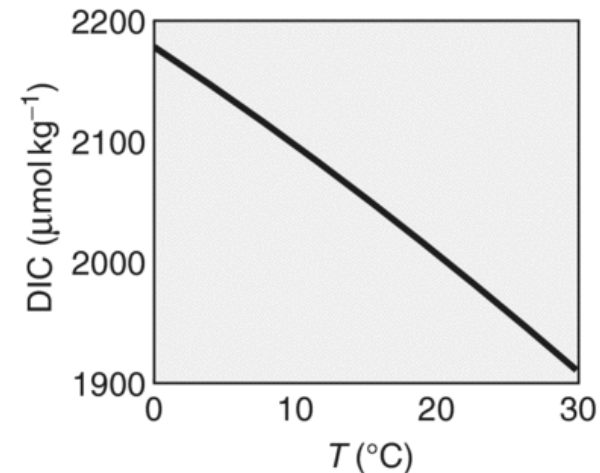
K_o is an inverse function of temperature

For same $[CO_2^*]$, pCO_2 is higher at high temperature, thus warm waters outgas and cold waters absorb CO_2

(b) solubility versus temperature

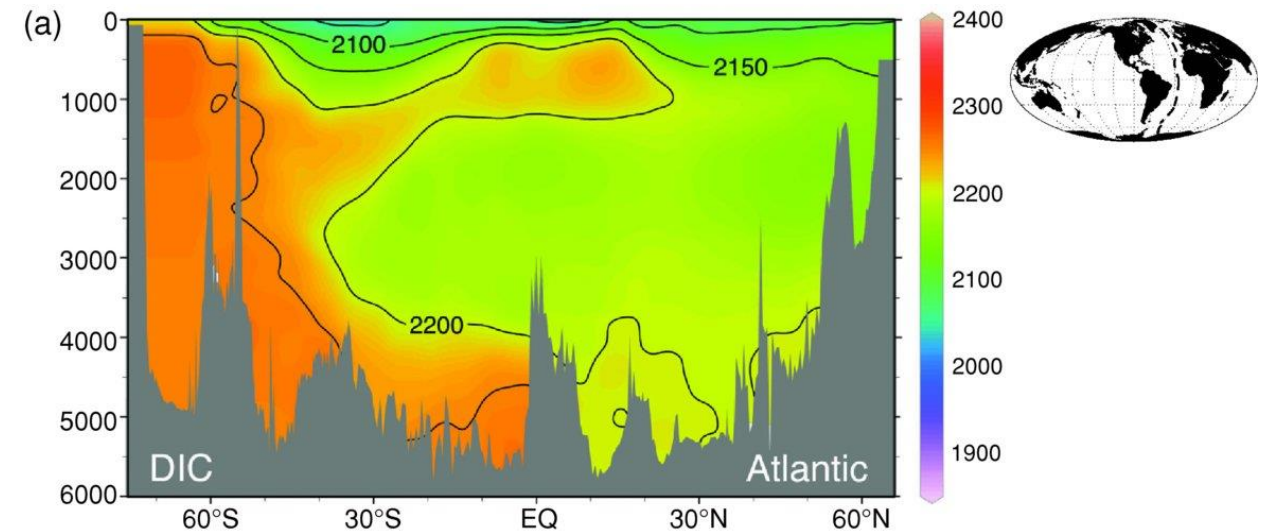
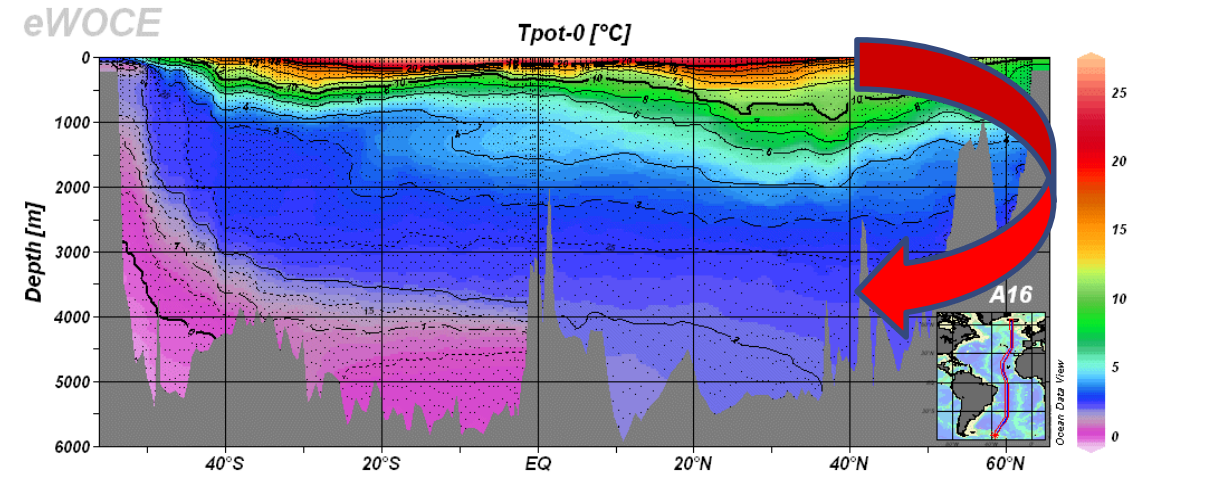
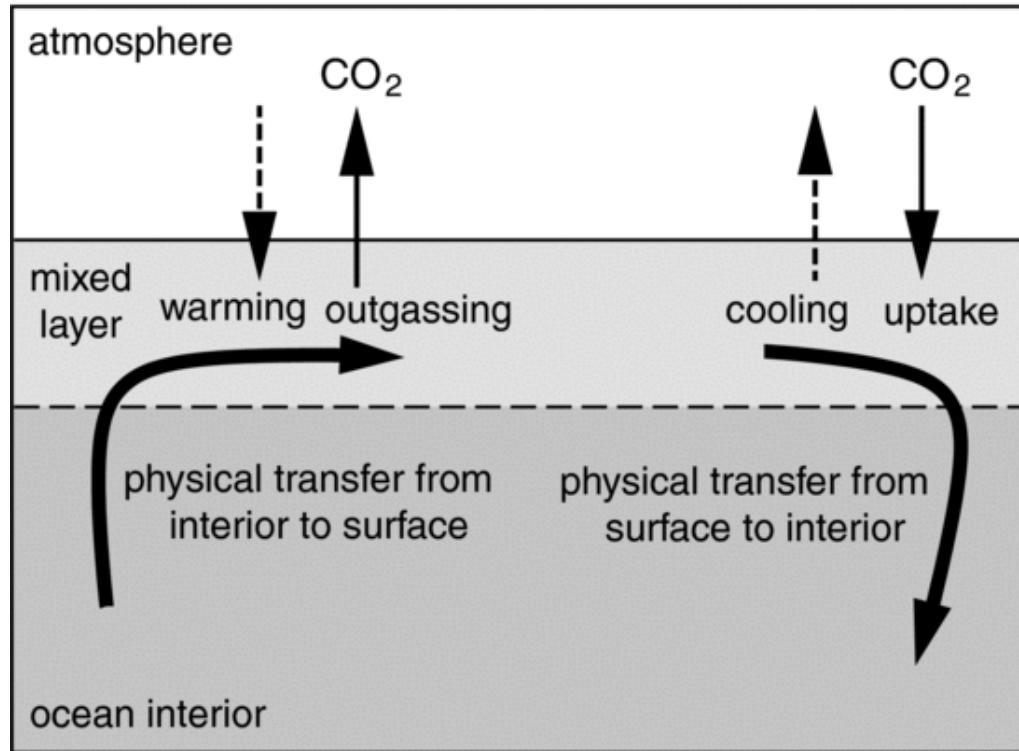


(c) DIC versus temperature



Solubility and physical transport

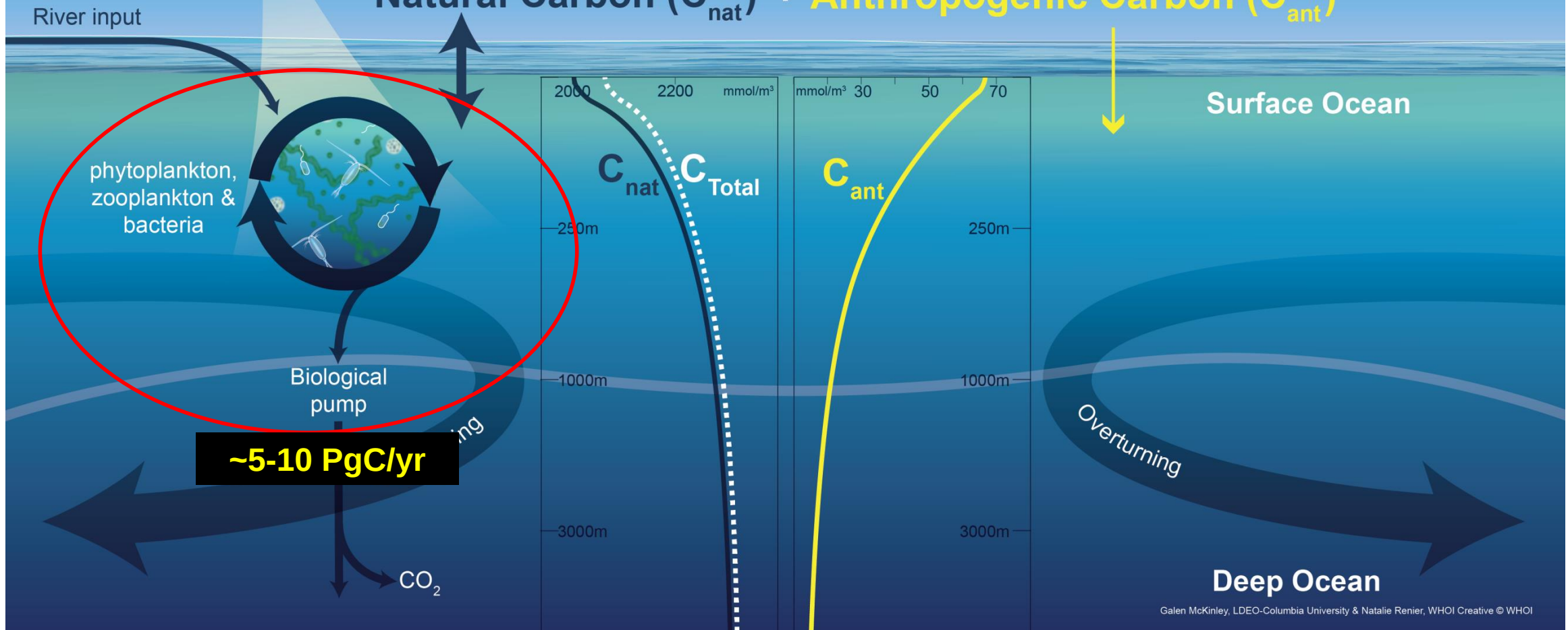
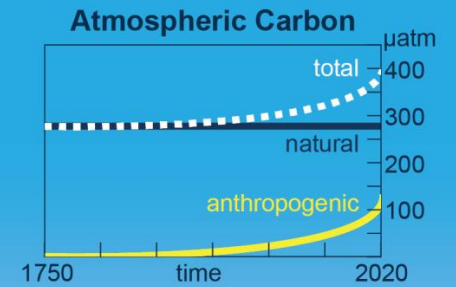
(a) physical transport and solubility



OCB

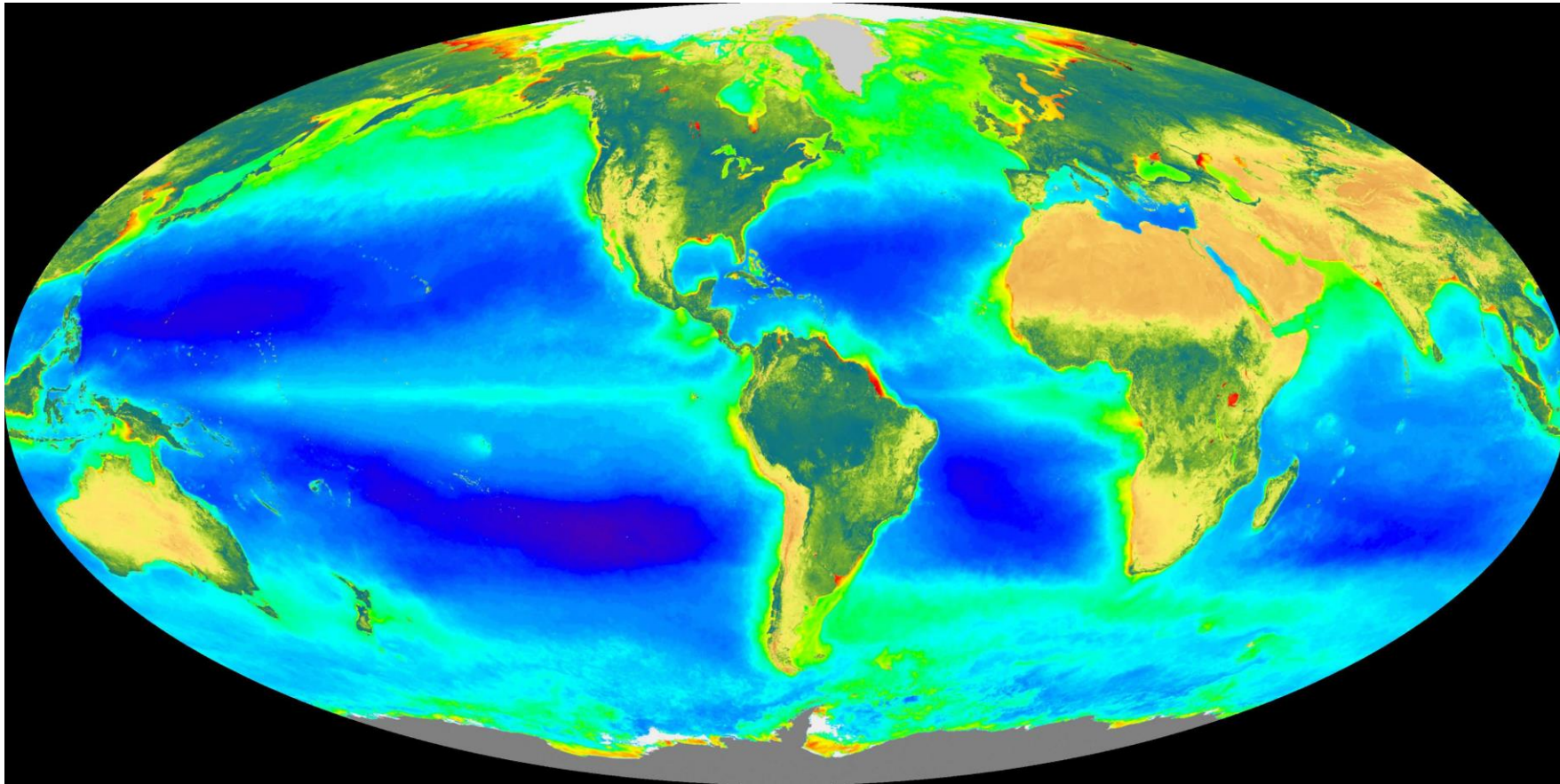
$$\text{Total Carbon (C}_{\text{Total}}) =$$

$$\text{Natural Carbon (C}_{\text{nat}}) + \text{Anthropogenic Carbon (C}_{\text{ant}})$$



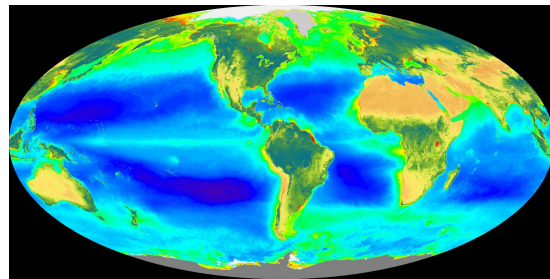
Galen McKinley, LDEO-Columbia University & Natalie Renier, WHOI Creative © WHOI

Ocean primary productivity from NASA SeaWiFS satellite

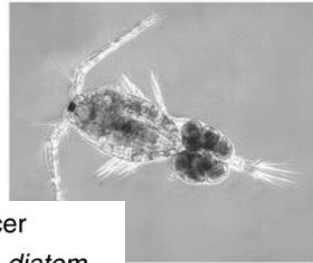


Dark blue = very low chlorophyll; green = moderate chlorophyll; red = very high

Biological carbon cycling in the ocean

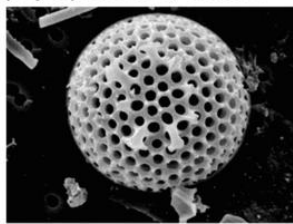


grazer
zooplankton, copepod



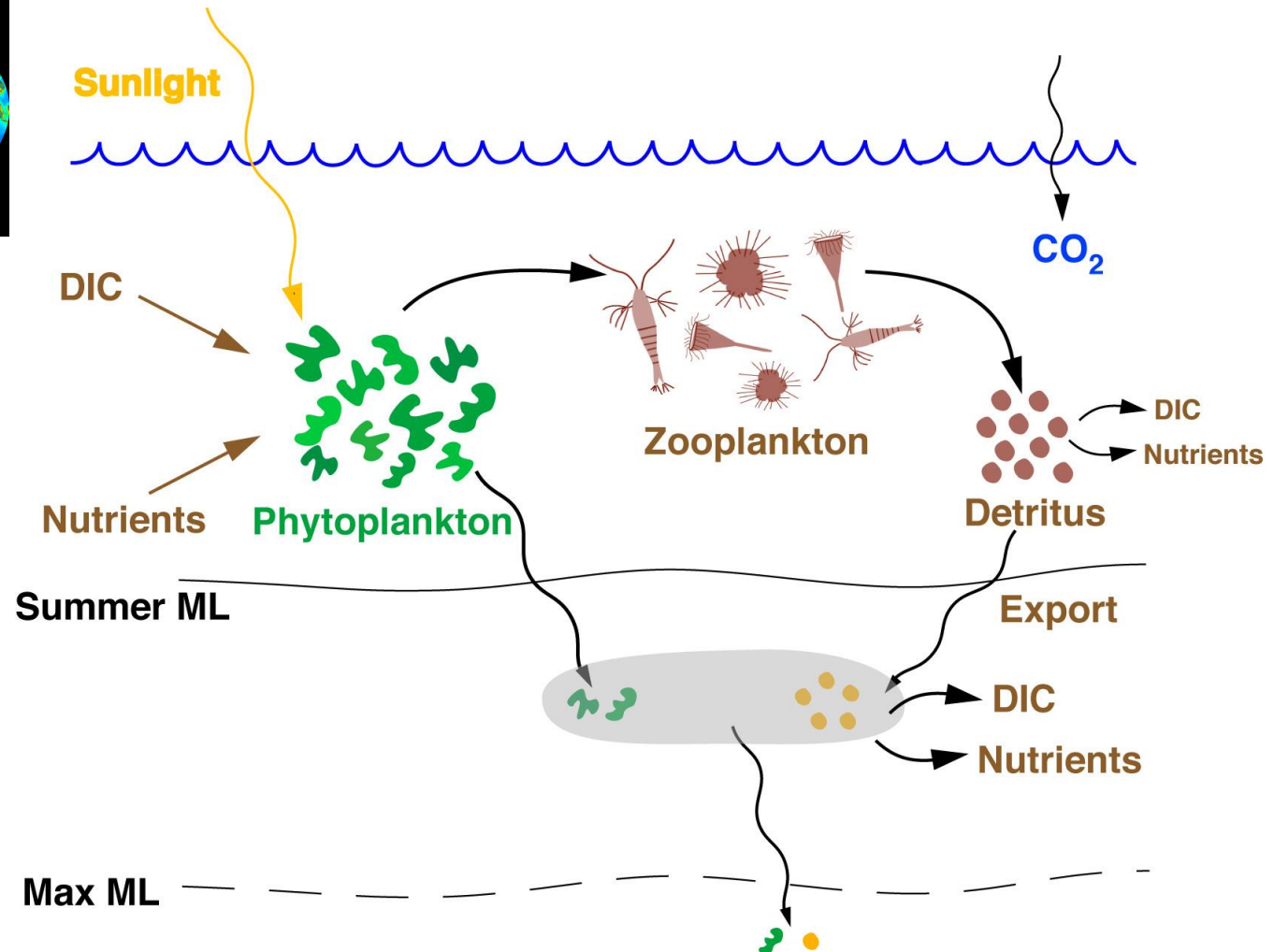
1 mm

primary producer
phytoplankton, diatom

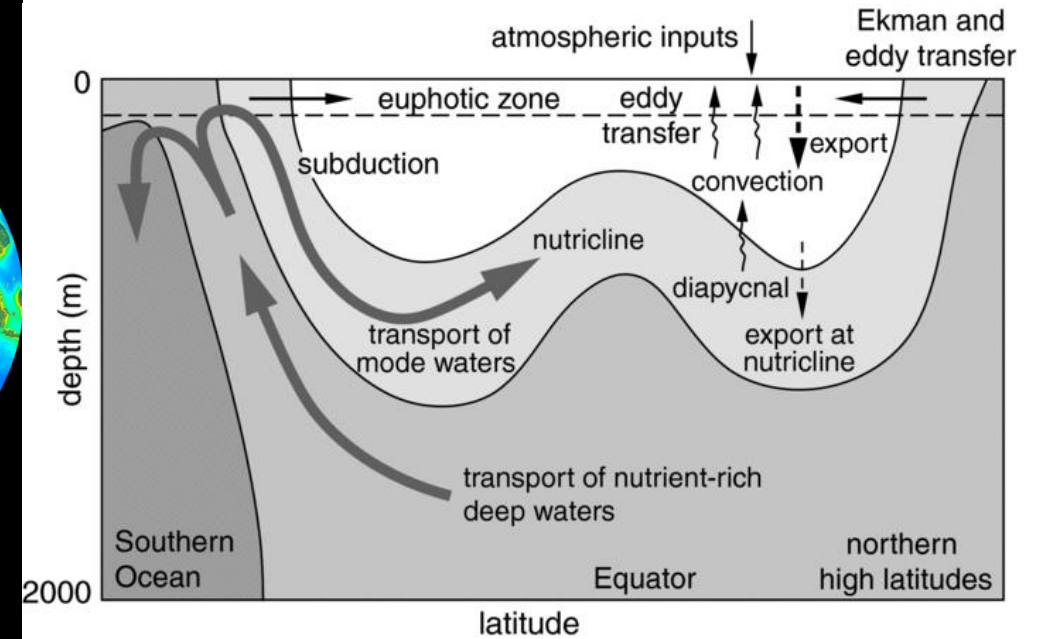
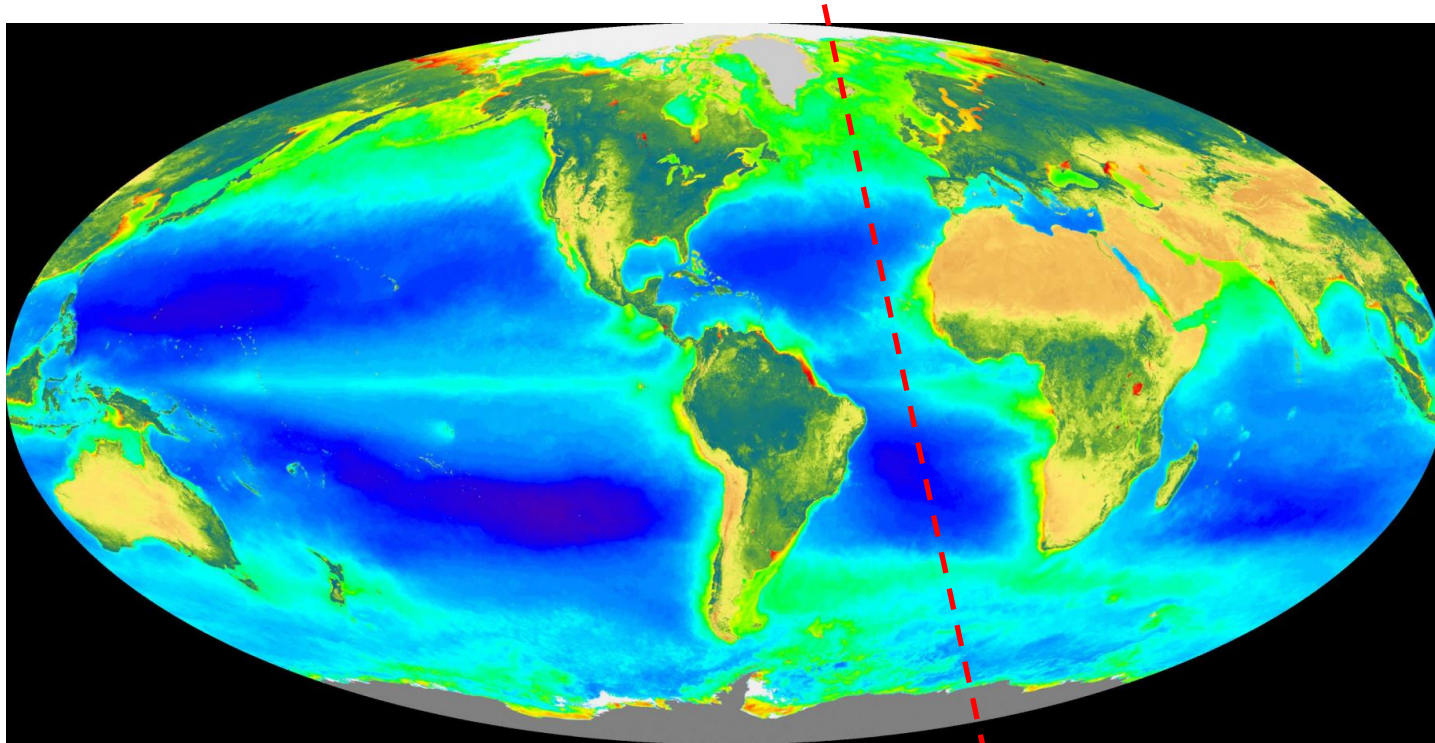


5 μ m

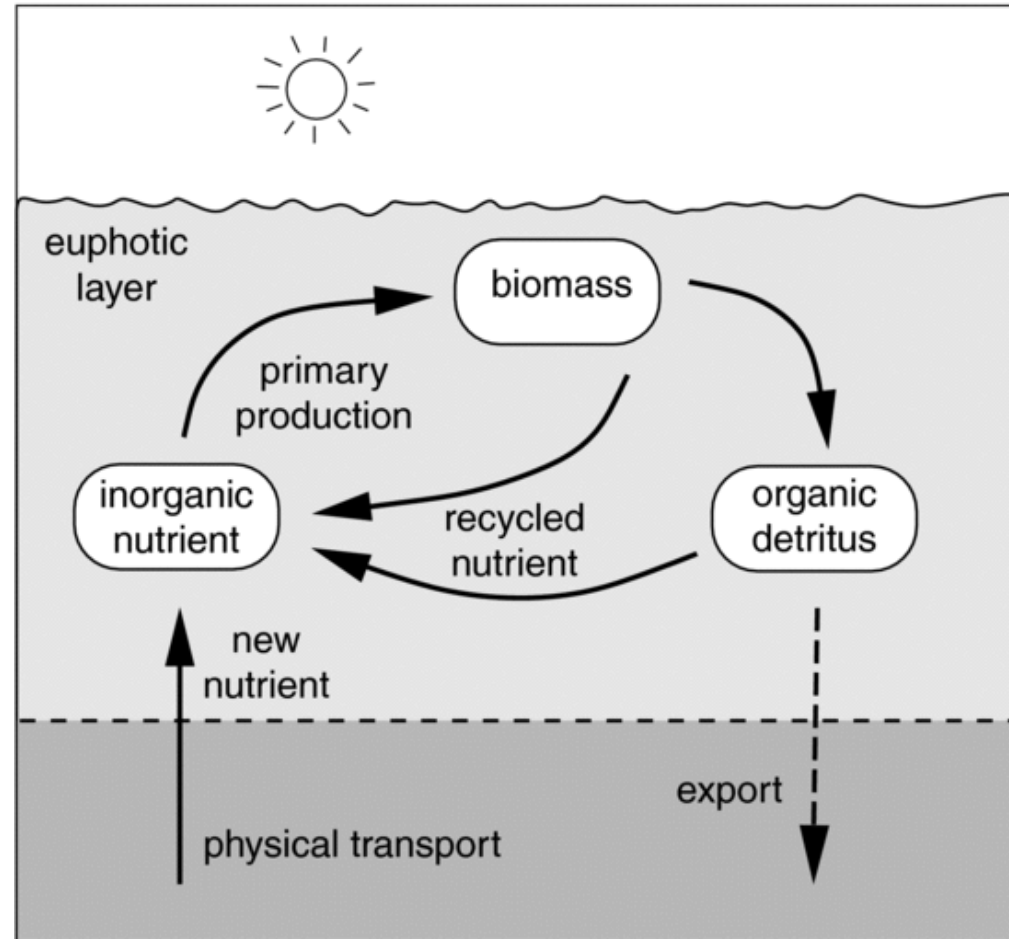
DIC = Dissolved
Inorganic Carbon



Productivity is structured by nutrients distributions, in turn set by the circulation

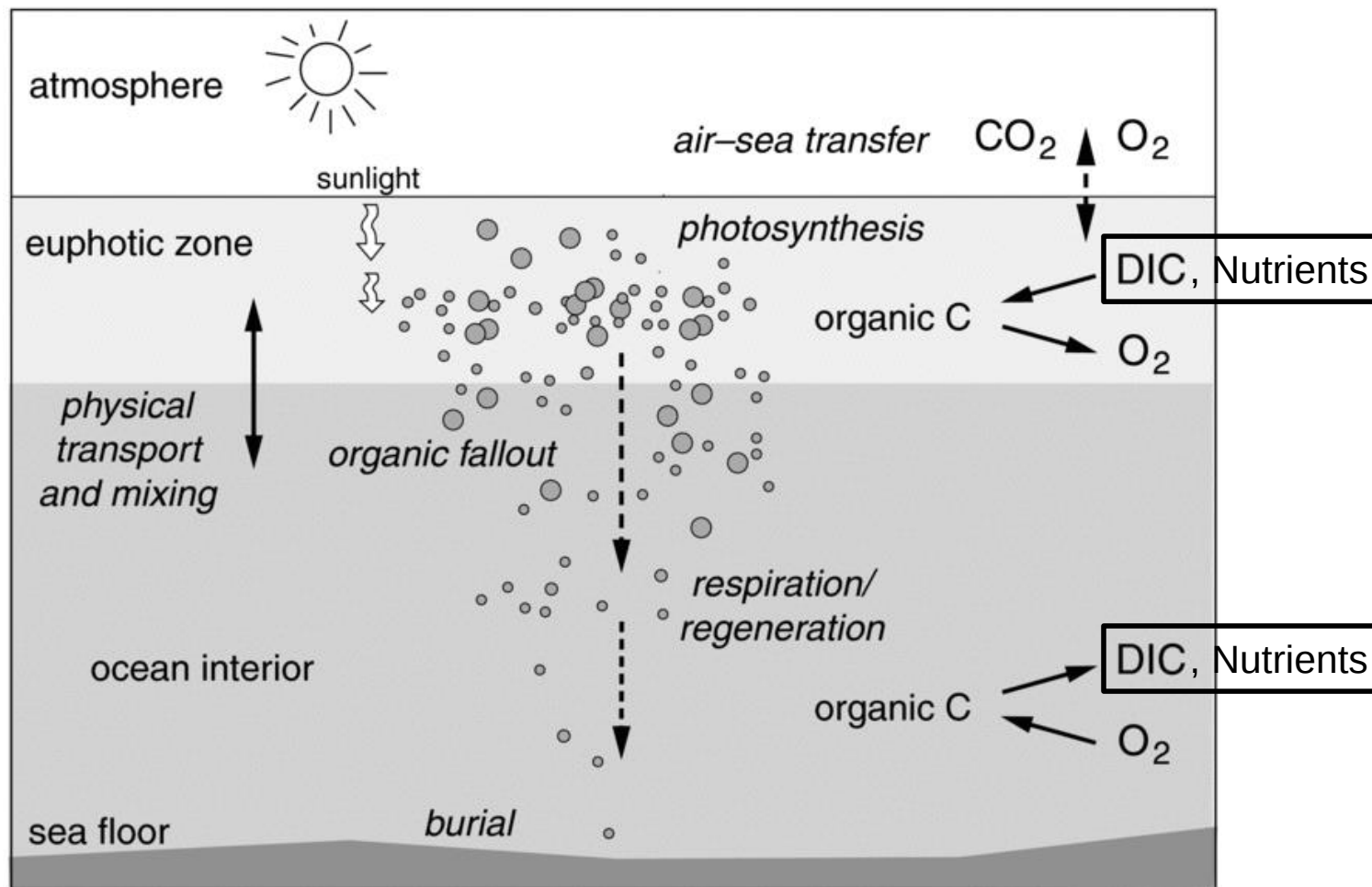


Nutrient cycling between biomass and inorganic nutrients



Biogeochemical Cycling = Coupled carbon and nutrient cycles

(b) carbon pathways in the ecosystem

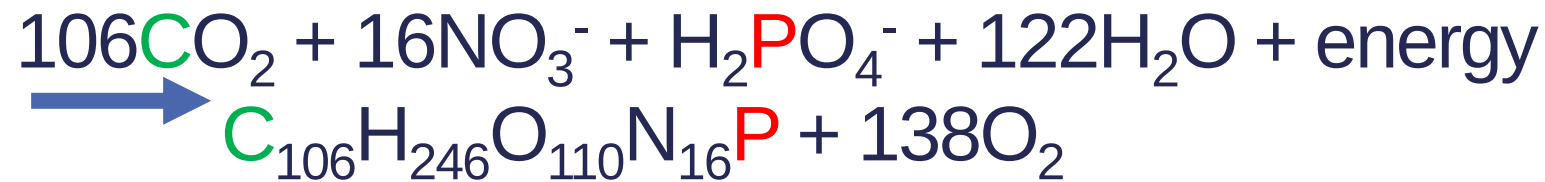


DIC = Dissolved Inorganic Carbon

Simplified photosynthesis

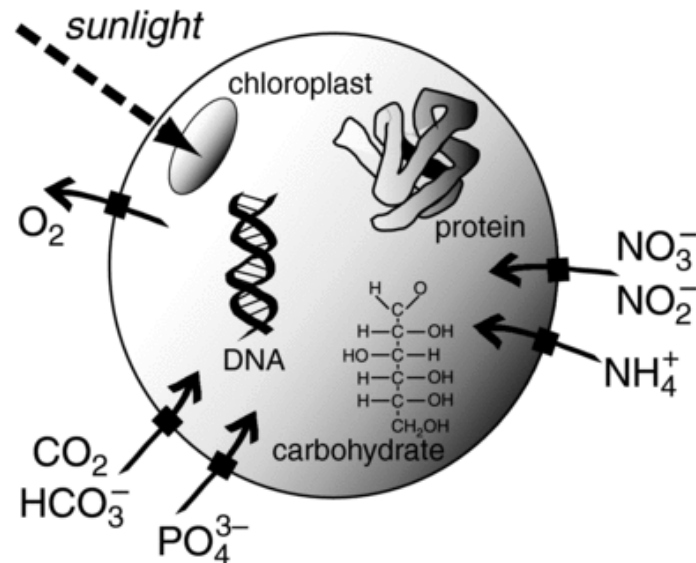


More realistic

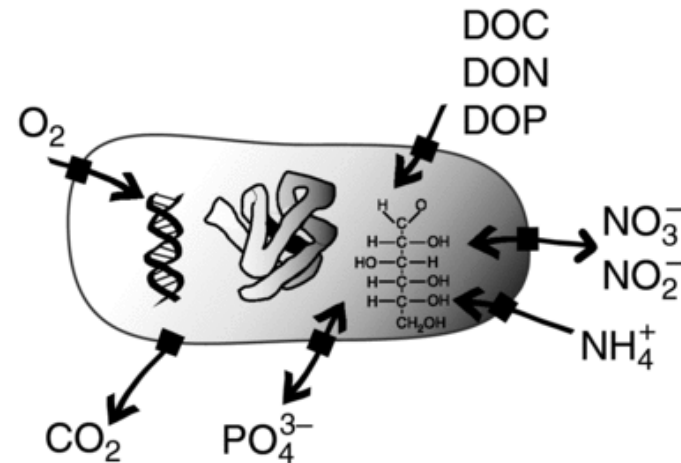


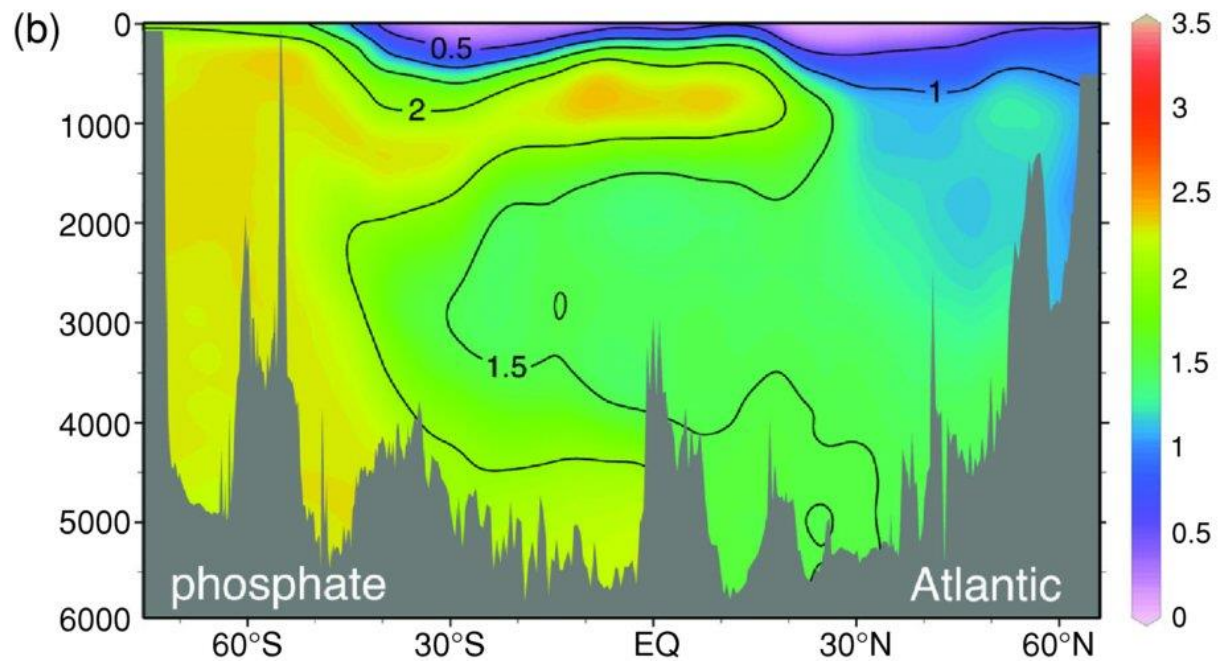
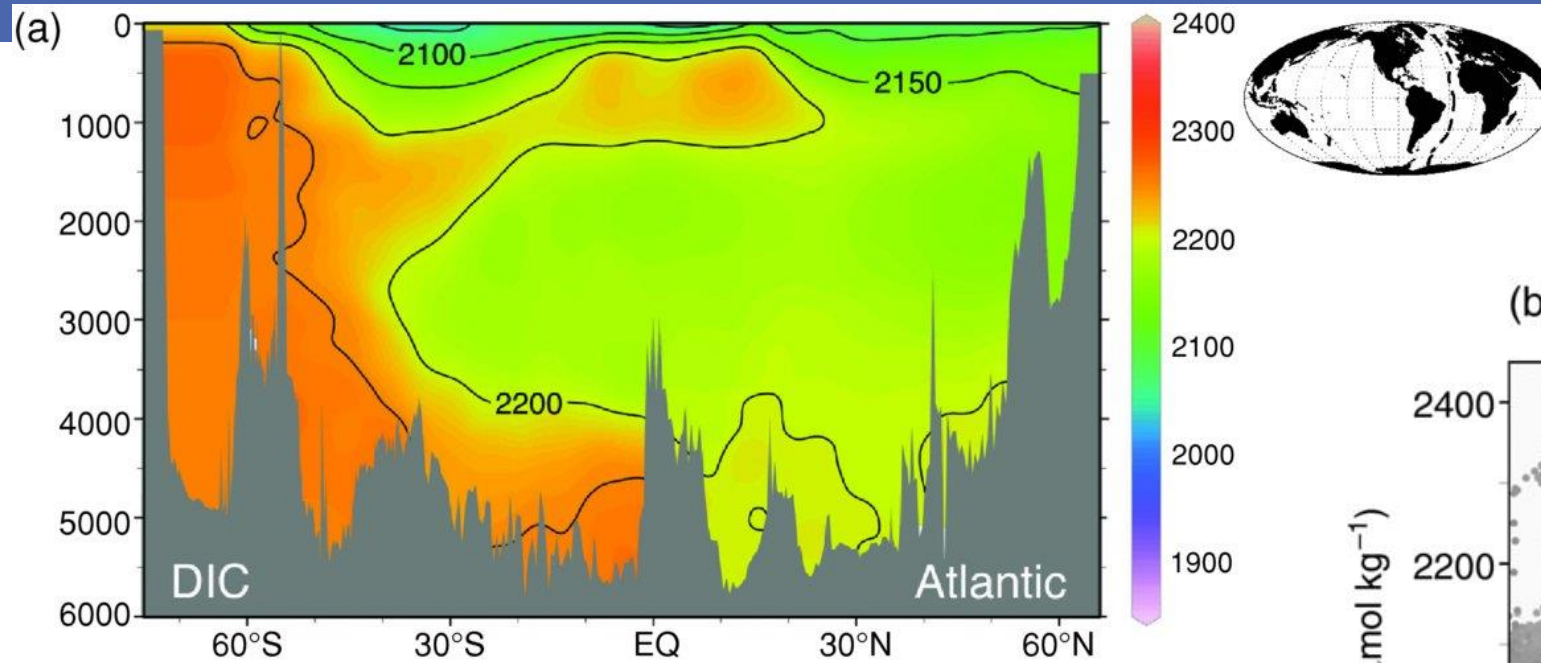
*Quantitative links
between **Carbon**,
Nutrients (N,P),
Oxygen*

(a) phytoplankton



(b) bacteria





Williams_Plate-6

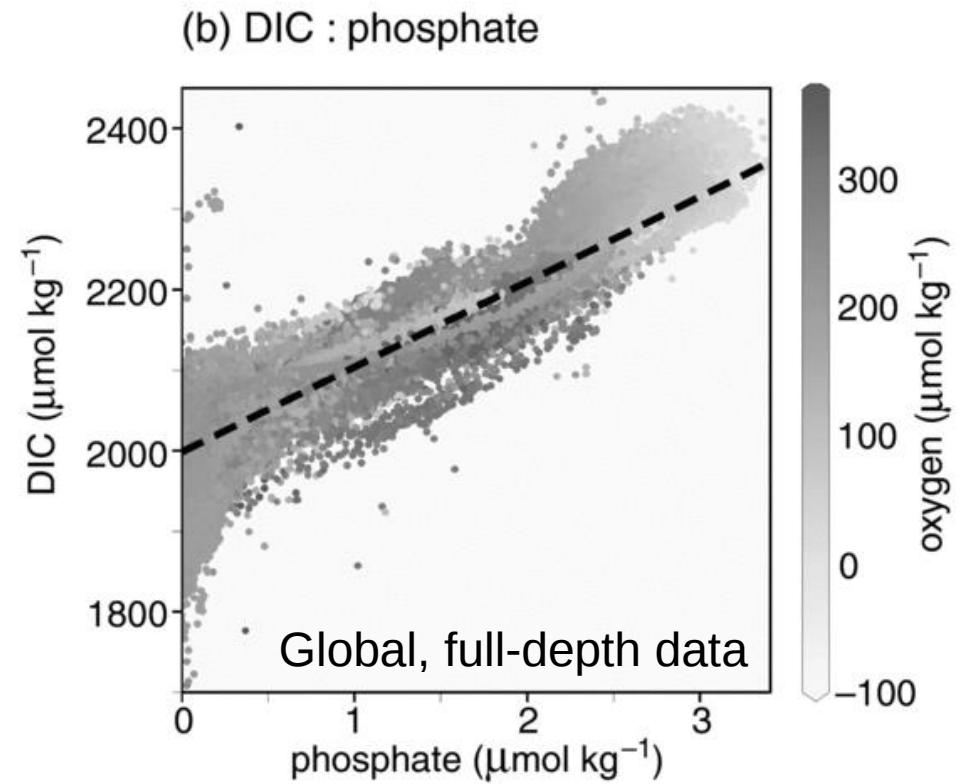


Table 5.2 | The ratios of key elements in dissolved inorganic form in the water column and in marine phytoplankton and zooplankton (^aRedfield *et al.*, 1963). The cultured phytoplankton data are the average of 15 species of estuarine, coastal and open ocean phytoplankton grown under controlled laboratory conditions (^bHo *et al.*, 2003). *Prochlorococcus* MED4 is a strain of tiny phytoplankton (about a micron in scale). Elemental ratios were measured in *Prochlorococcus* cultures grown under phosphorus replete and phosphorus limited conditions (^cBertilsson *et al.*, 2003). Ranges of values are also shown for isolated marine bacteria grown under various conditions of nutrient limitation (^dVrede *et al.*, 2002).

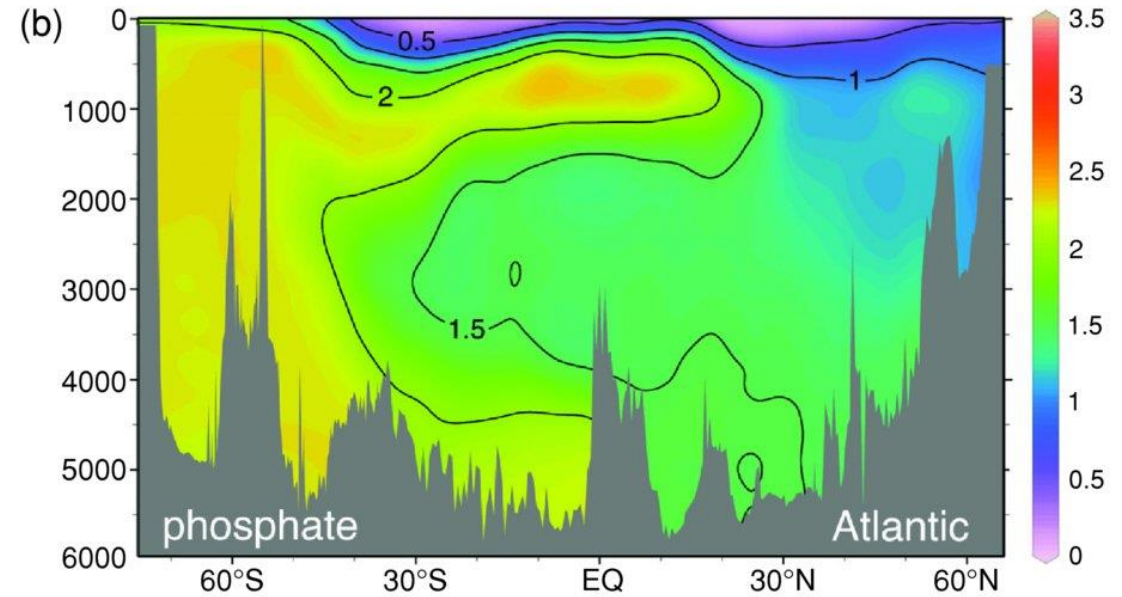
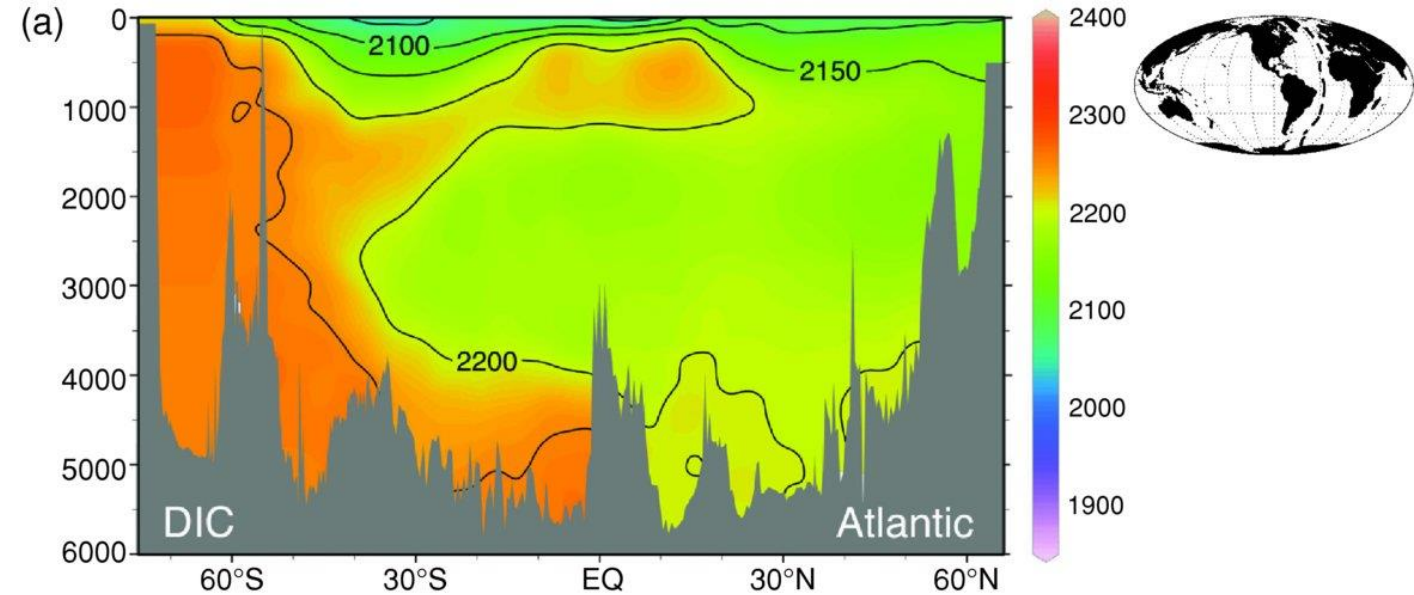
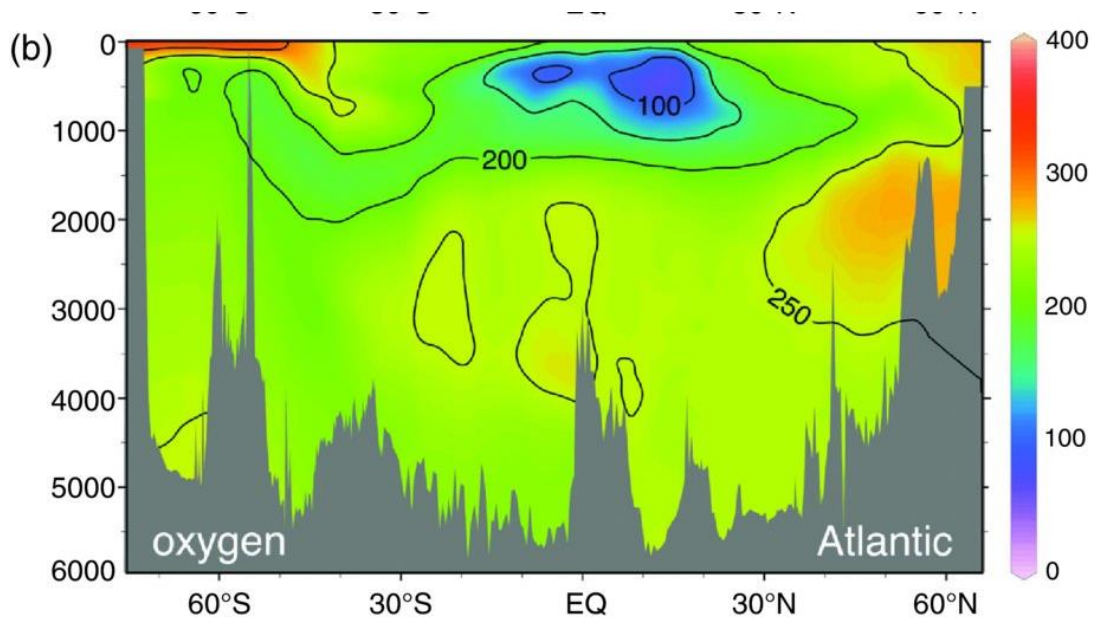
Reservoir	C	N	P
^a Marine inorganic nutrients		15	1
^a Bulk marine organic matter	106	16	1
^b Cultured phytoplankton	147 ± 19	16 ± 2	1 ± 0.2
^c <i>Prochlorococcus</i> MED4 (P-replete)	121 ± 17	21.2 ± 4.5	1
^c <i>Prochlorococcus</i> MED4 (P-limited)	464 ± 28	62.3 ± 14.1	1
^d Marine bacteria	35–178	7–18	1

Updated Redfield

$$\text{C:N:P:O}_2 = 117(\pm 14):16(\pm 1):1:-170(\pm 10)$$

Anderson and Sarmiento, 2004

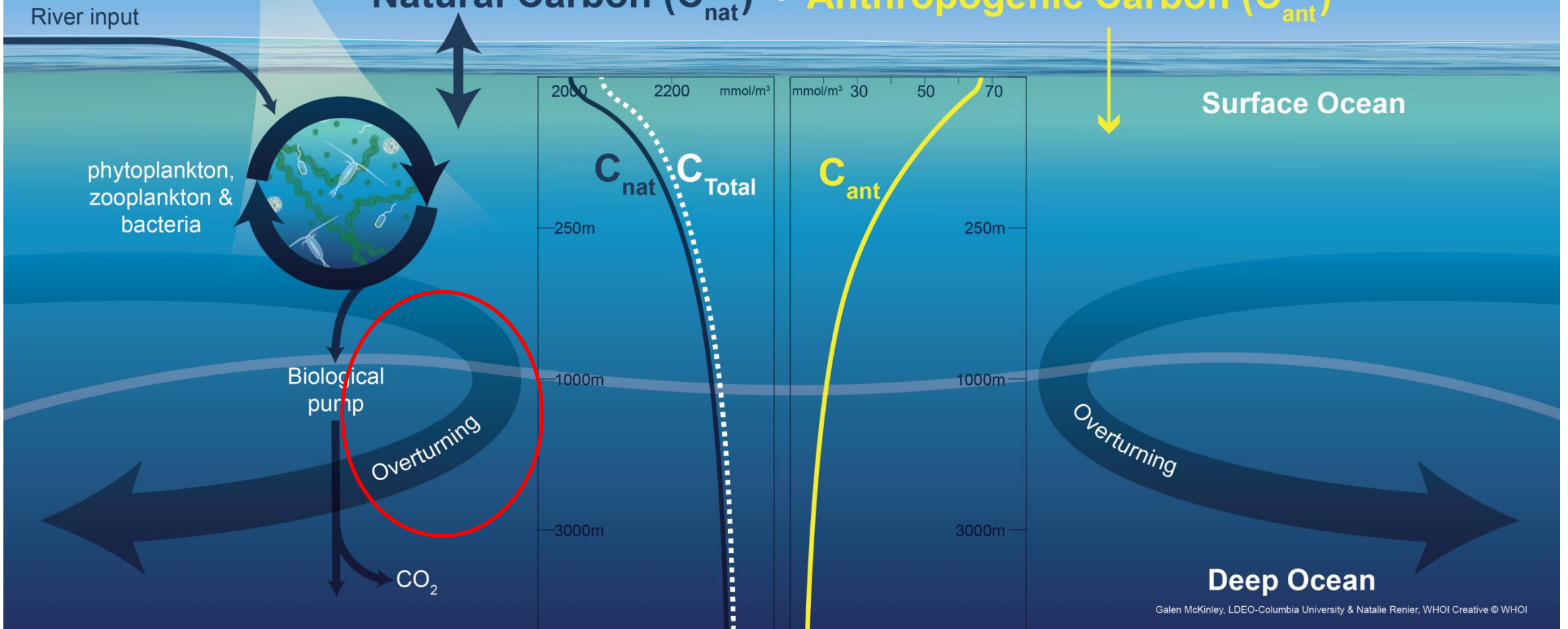
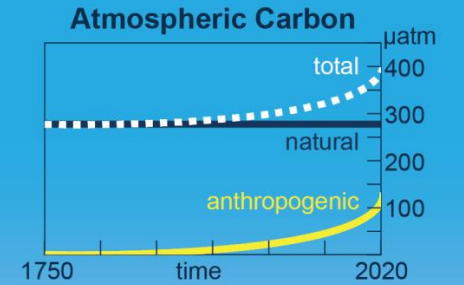
Oxygen distribution is inverse
of nutrients and DIC



OCB

$$\text{Total Carbon (C}_{\text{Total}}) =$$

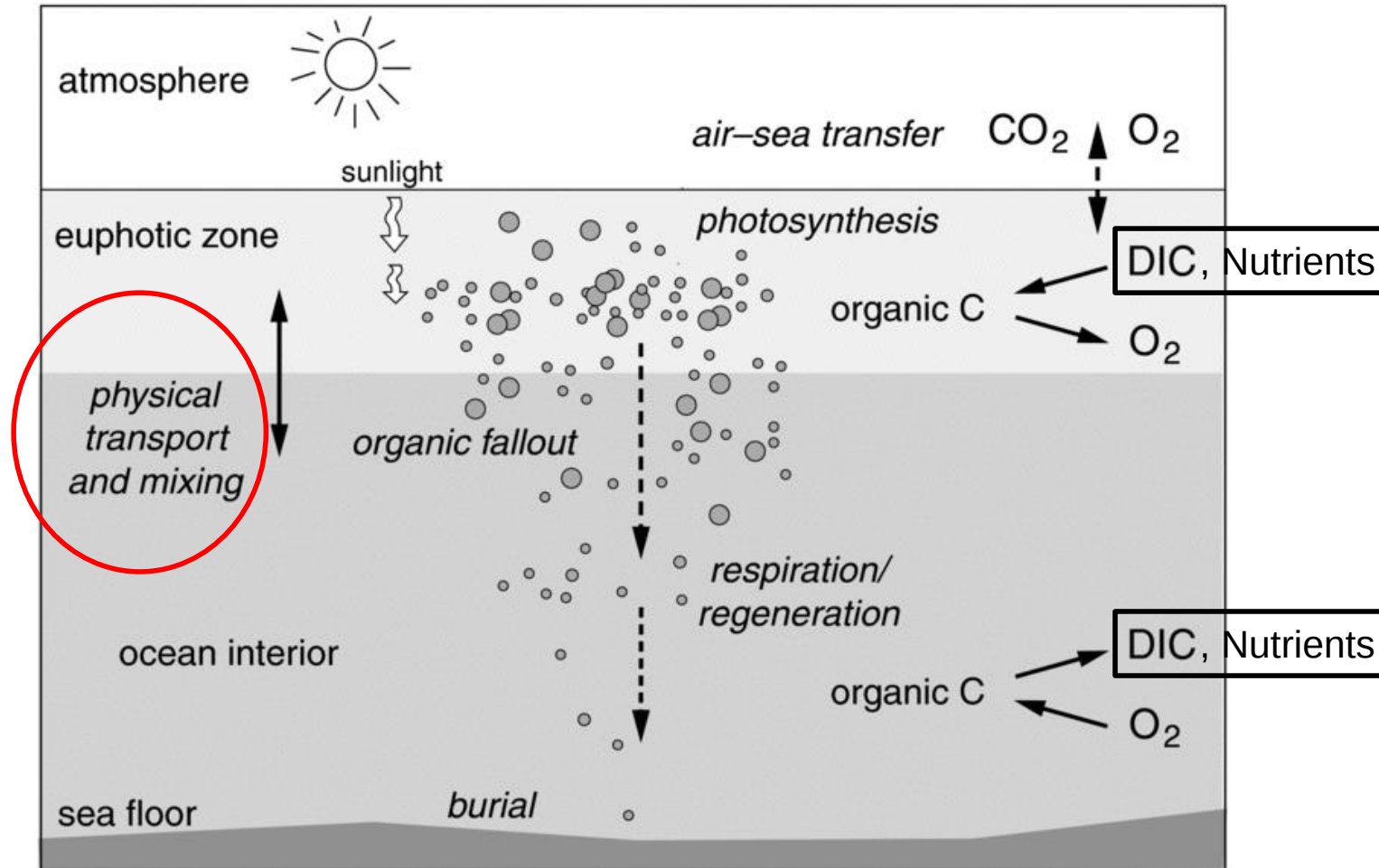
$$\text{Natural Carbon (C}_{\text{nat}}) + \text{Anthropogenic Carbon (C}_{\text{ant}})$$



Galen McKinley, LDEO-Columbia University & Natalie Renier, WHOI Creative © WHOI

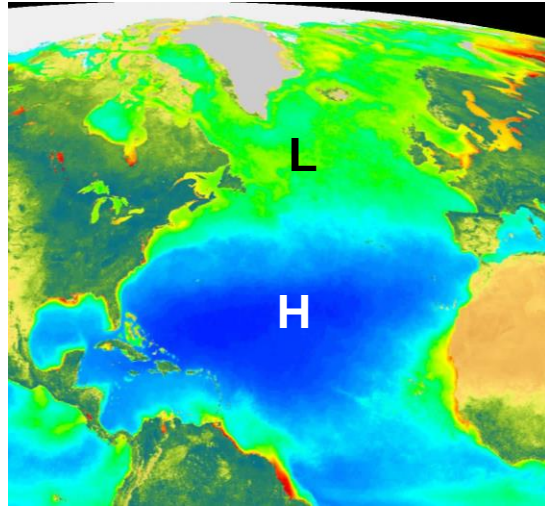
Biogeochemical Cycling = Coupled carbon and nutrient cycles

(b) carbon pathways in the ecosystem

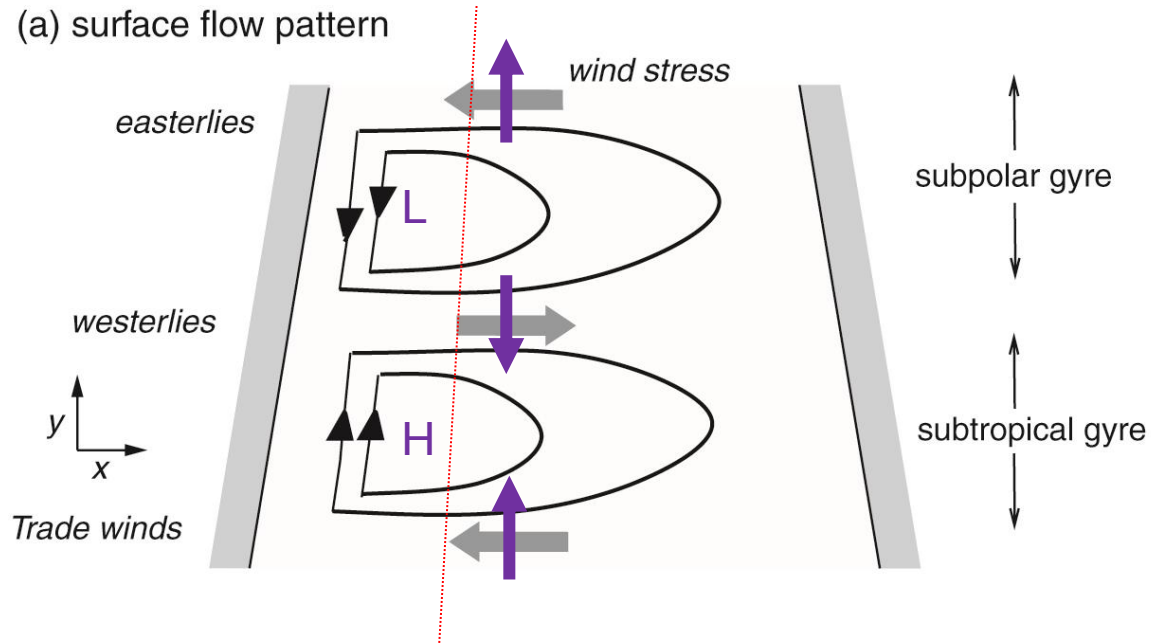


DIC = Dissolved Inorganic Carbon

Circulation redistributes carbon and nutrients

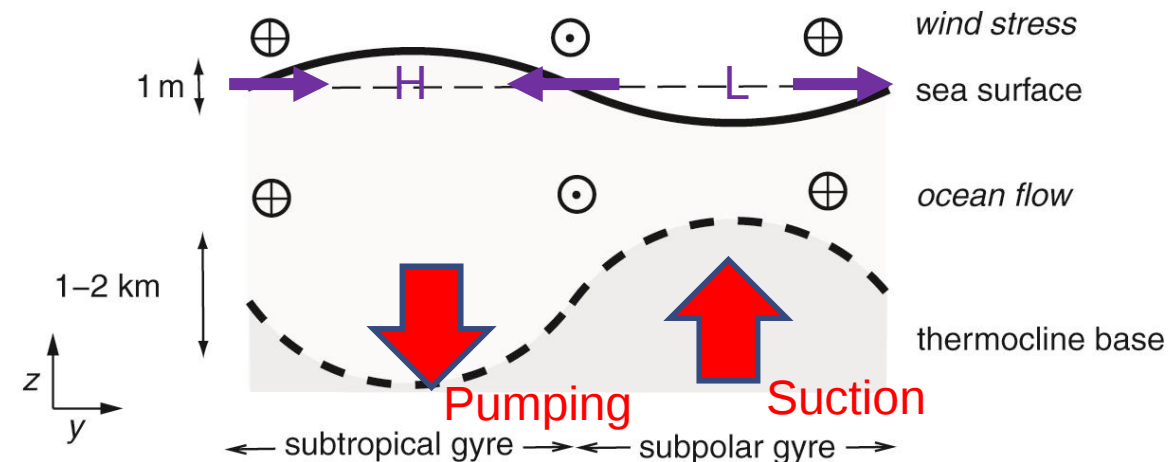


(a) surface flow pattern



Winds + Coriolis force leads to mass convergence, divergence at surface; causing Highs and Lows in surface height

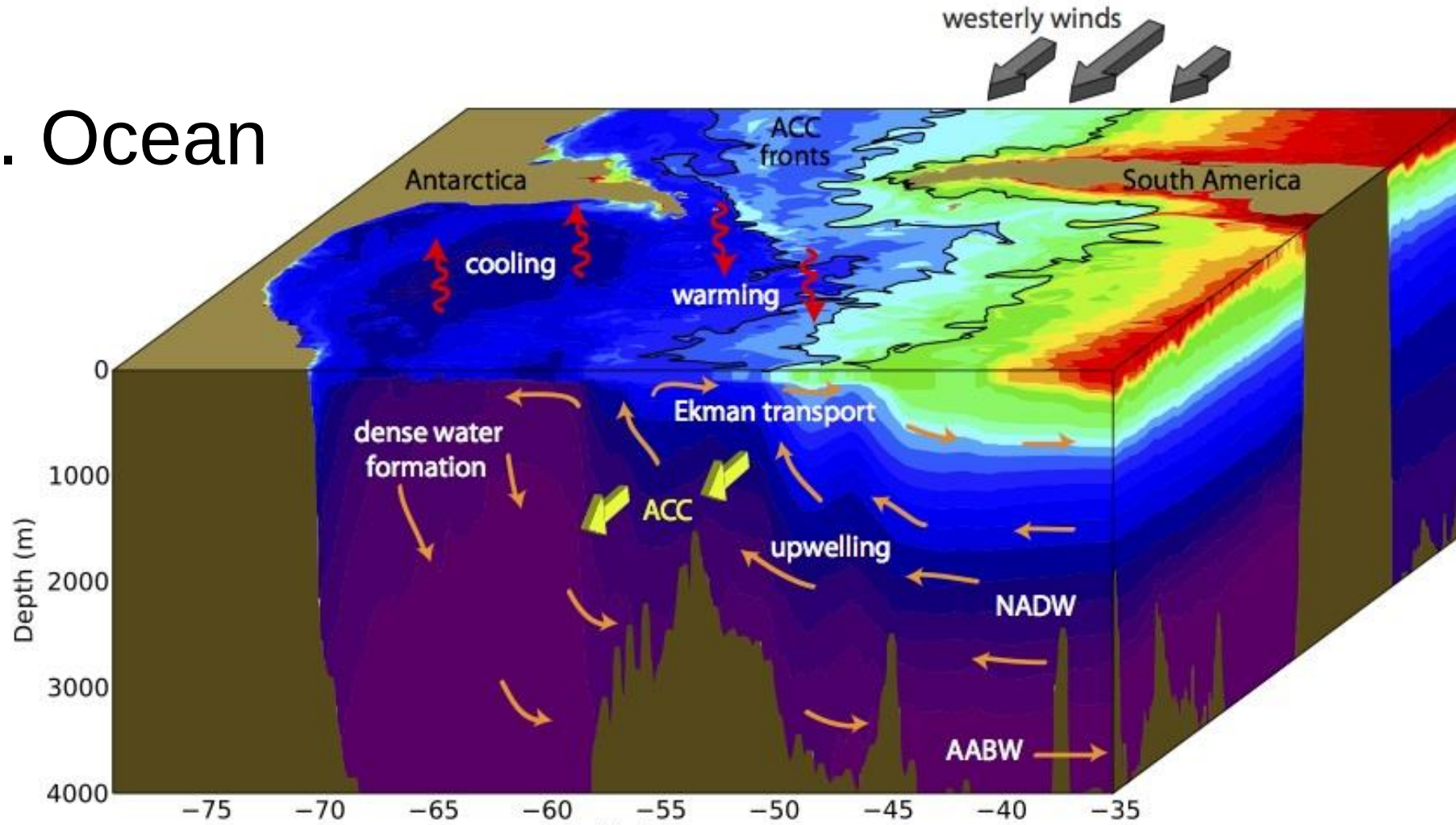
(b) meridional section (along dotted red line above)



Horizontal flows lead to vertical motions; advecting carbon and nutrients

Circulation redistributes carbon and nutrients

S. Ocean

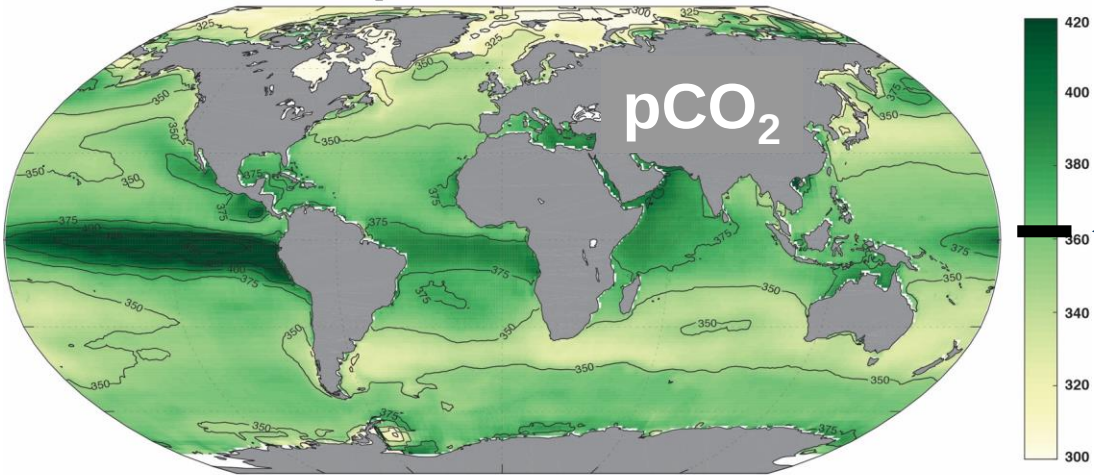


<http://rses.anu.edu.au/research/projects/southern-ocean-circulation>

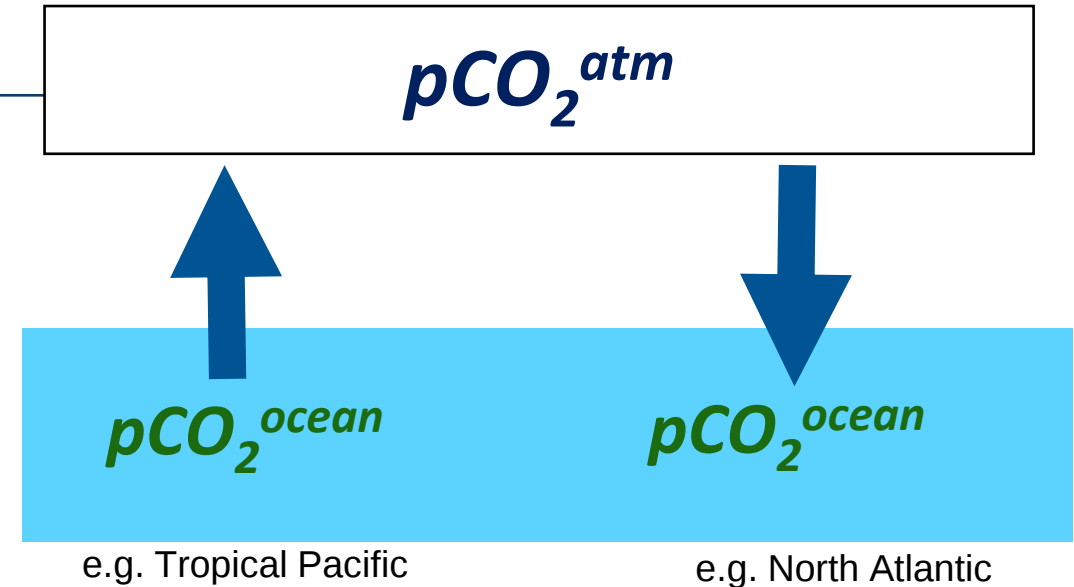
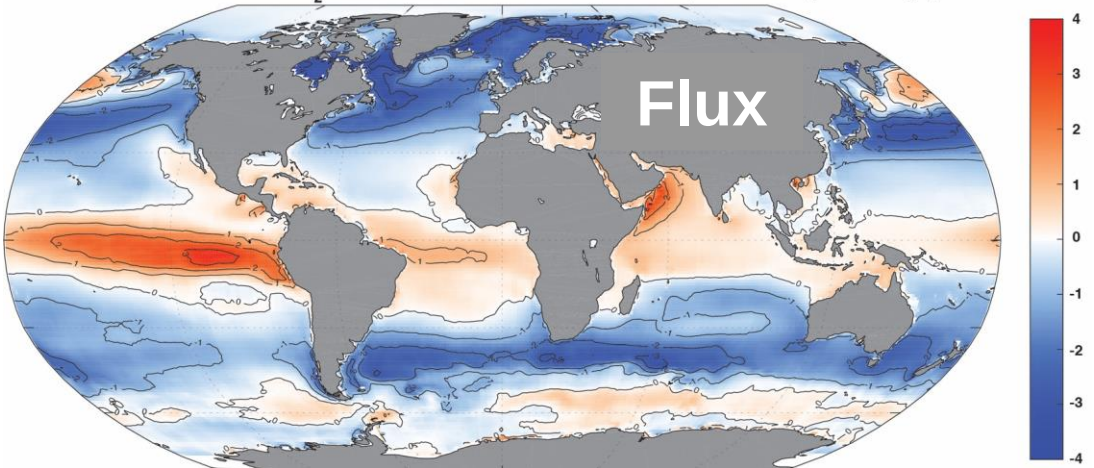
Surface ocean $p\text{CO}_2$ is primary control on air-sea CO_2 flux

If $p\text{CO}_2^{\text{ocean}} > p\text{CO}_2^{\text{atm}}$ flux is out of ocean (positive)
 If $p\text{CO}_2^{\text{ocean}} < p\text{CO}_2^{\text{atm}}$ flux is into ocean (negative)

1988-2018 mean $p\text{CO}_2$ from observation-based products (μatm)

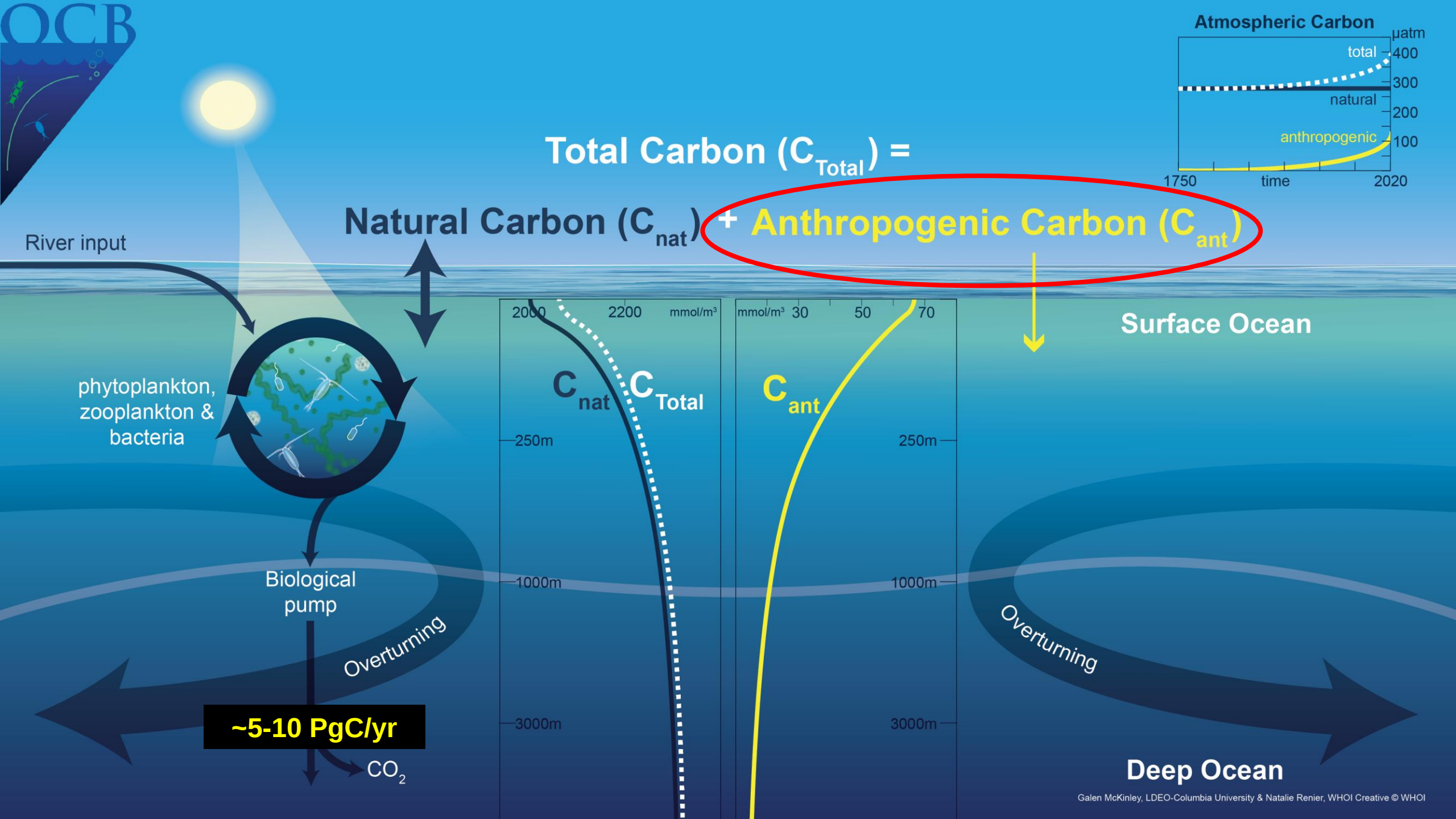


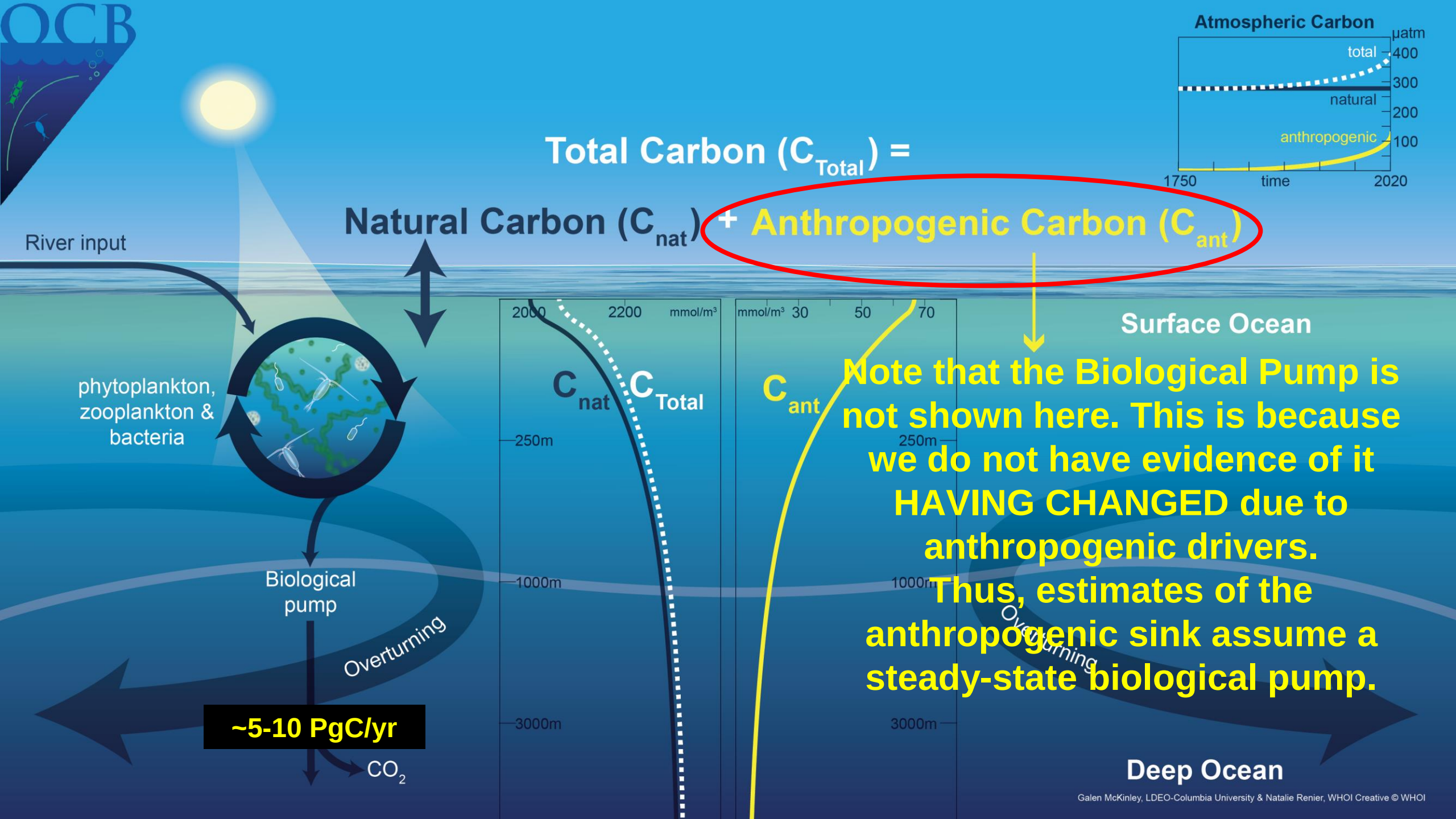
1988-2018 mean CO_2 flux from observation-based products ($\text{mol}/\text{m}^2/\text{yr}$)



$$\text{Air sea } \text{CO}_2 \text{ flux} = k_w S_{\text{CO}_2} (1 - f_{\text{ice}}) (p\text{CO}_2^{\text{ocean}} - p\text{CO}_2^{\text{atm}})$$

Crisp et al., 2022, Review of Geophysics





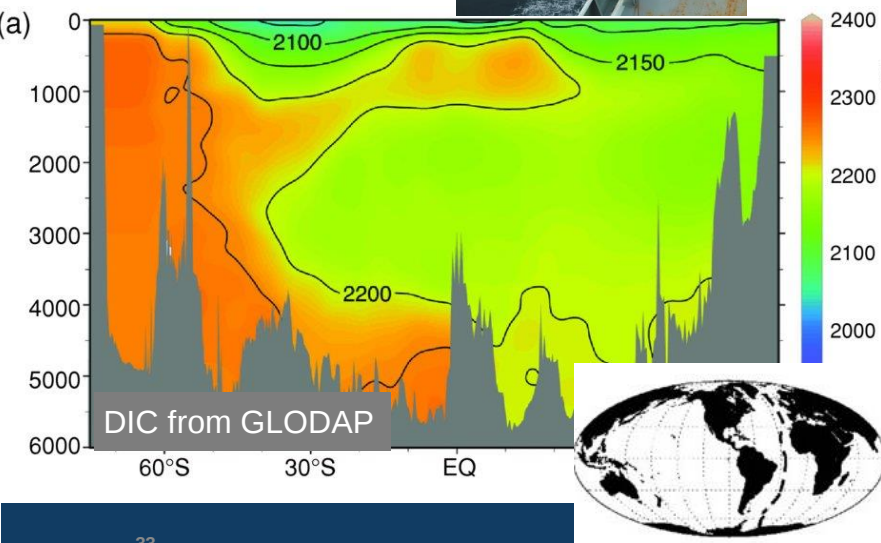
Three independent approaches constrain the ocean carbon sink

1. Interior observations / products

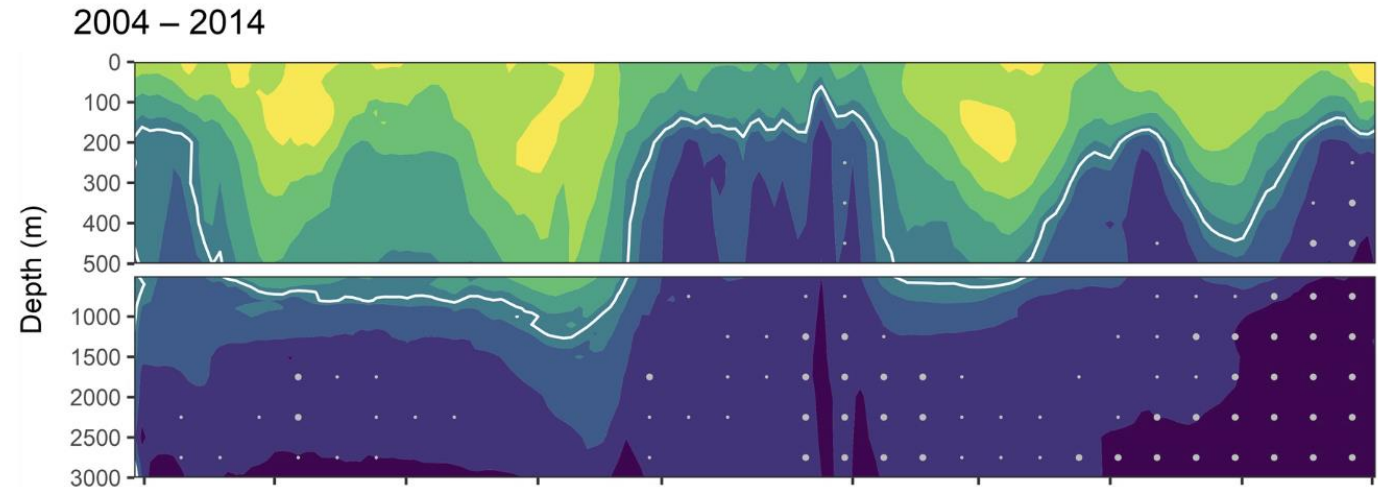
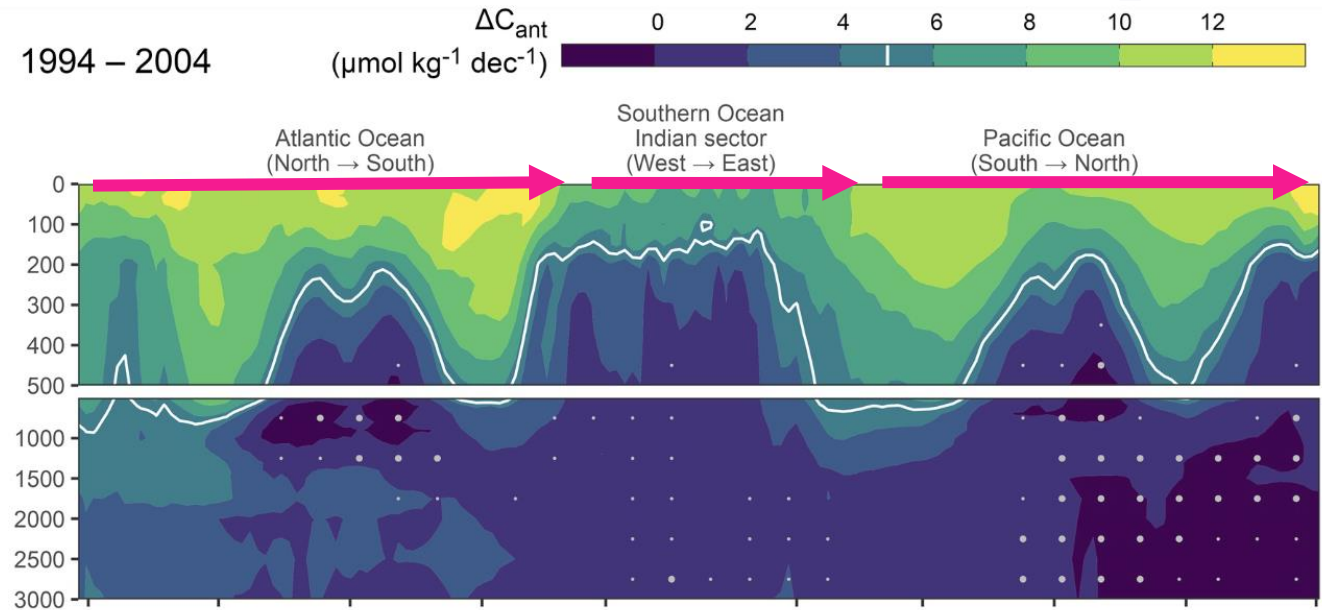
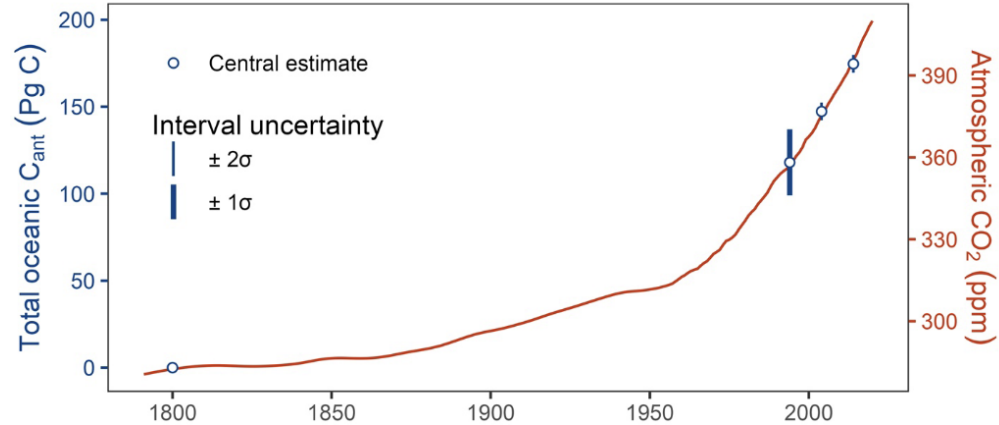
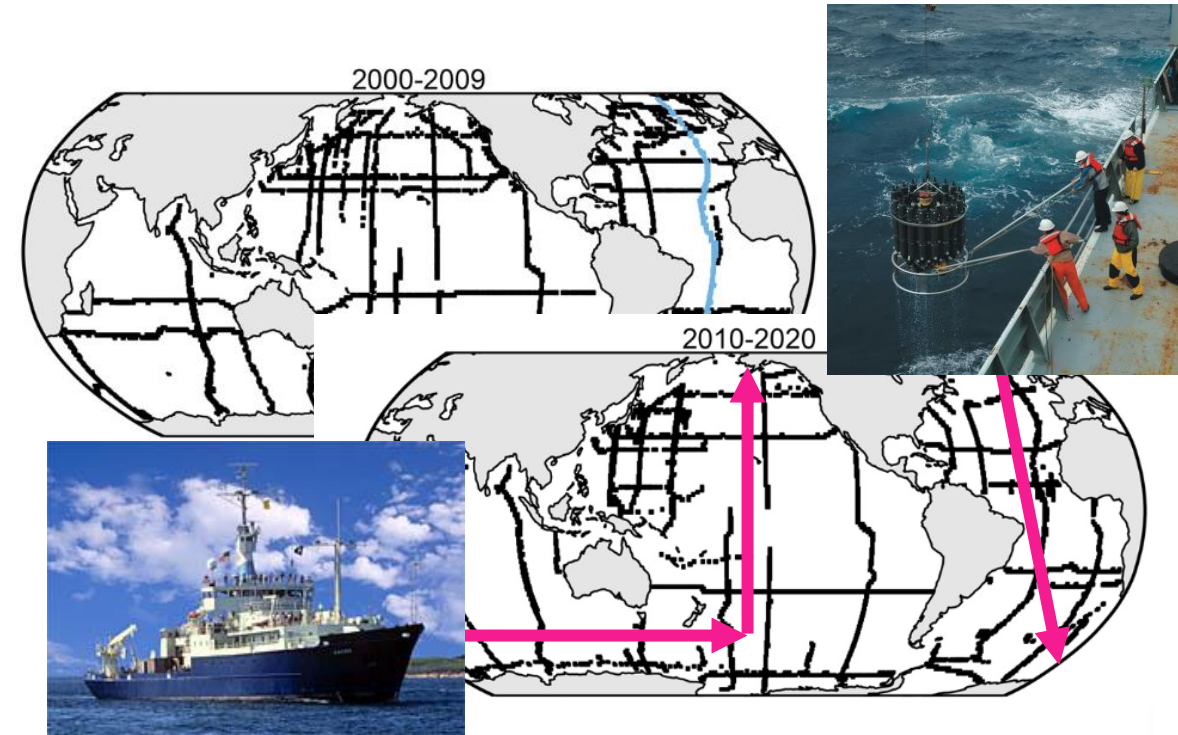
Interior carbon storage

Decadal closure of global budget

Model validation

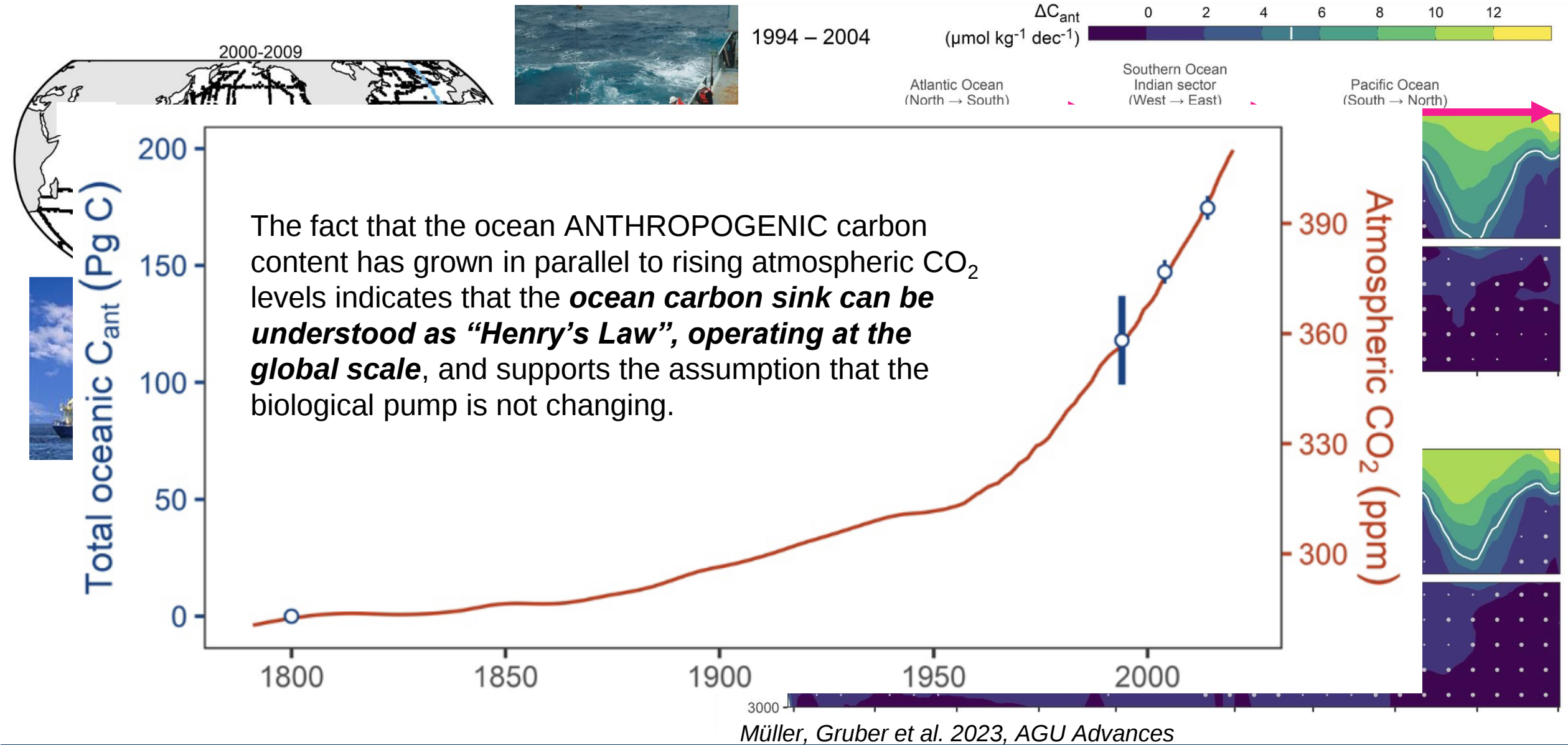


Decadal Interior Data: Shows ocean following atmospheric $p\text{CO}_2$



Müller, Gruber et al. 2023, AGU Advances

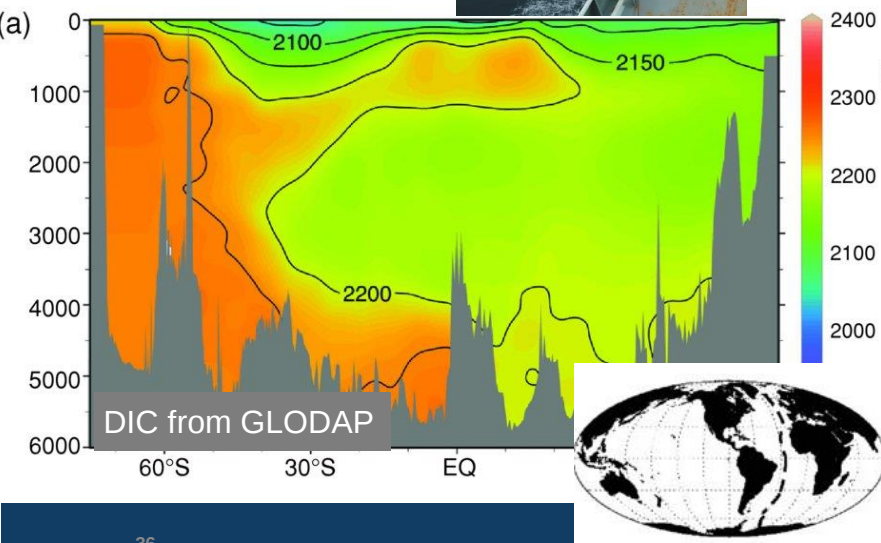
Decadal Interior Data: Shows ocean following atmospheric $p\text{CO}_2$



Three independent approaches constrain the ocean carbon sink

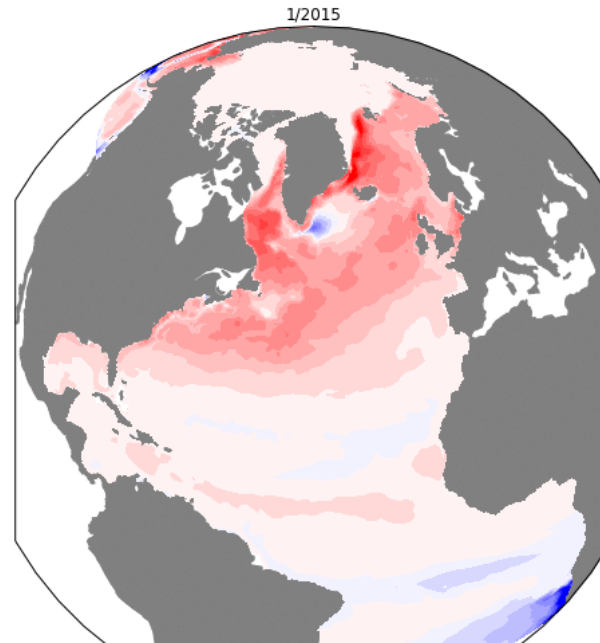
1. Interior observations / products

Interior carbon storage
Decadal closure of global budget
Model validation



2. Modeling

Air-sea fluxes
Mechanisms
Projections



ASTE-BGC
air-sea CO₂ flux

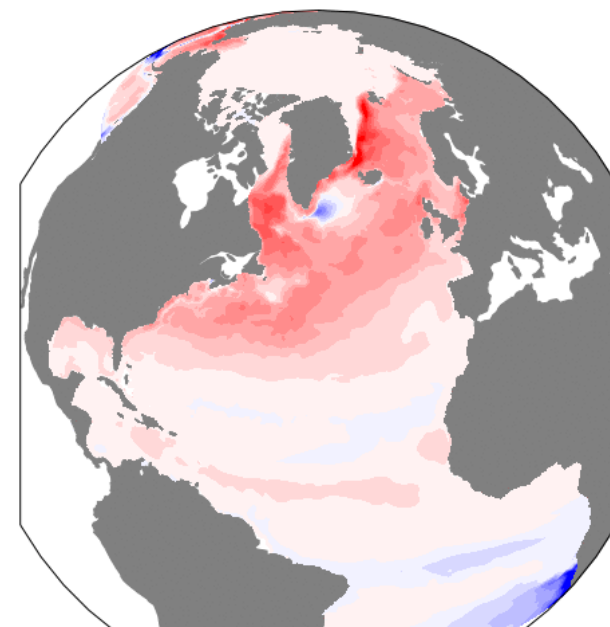
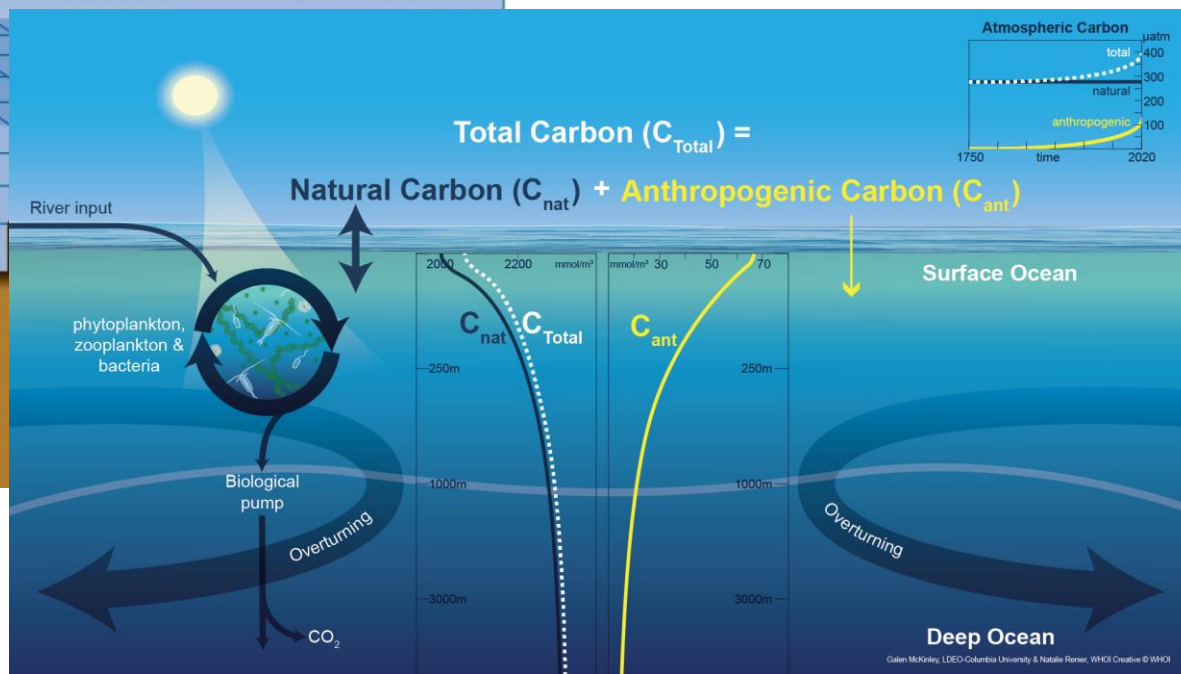
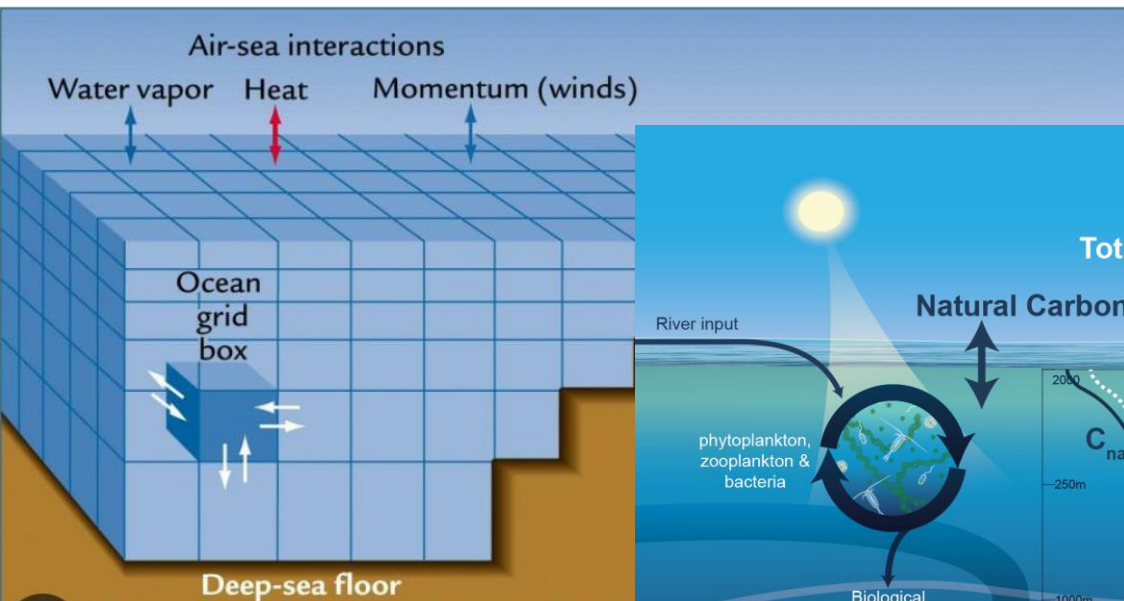
Moseley et al, in review

Ocean biogeochemical models

Circulation

+ Carbon
Processes

= Air-sea CO₂ fluxes



ASTE-BGC
air-sea CO₂ flux

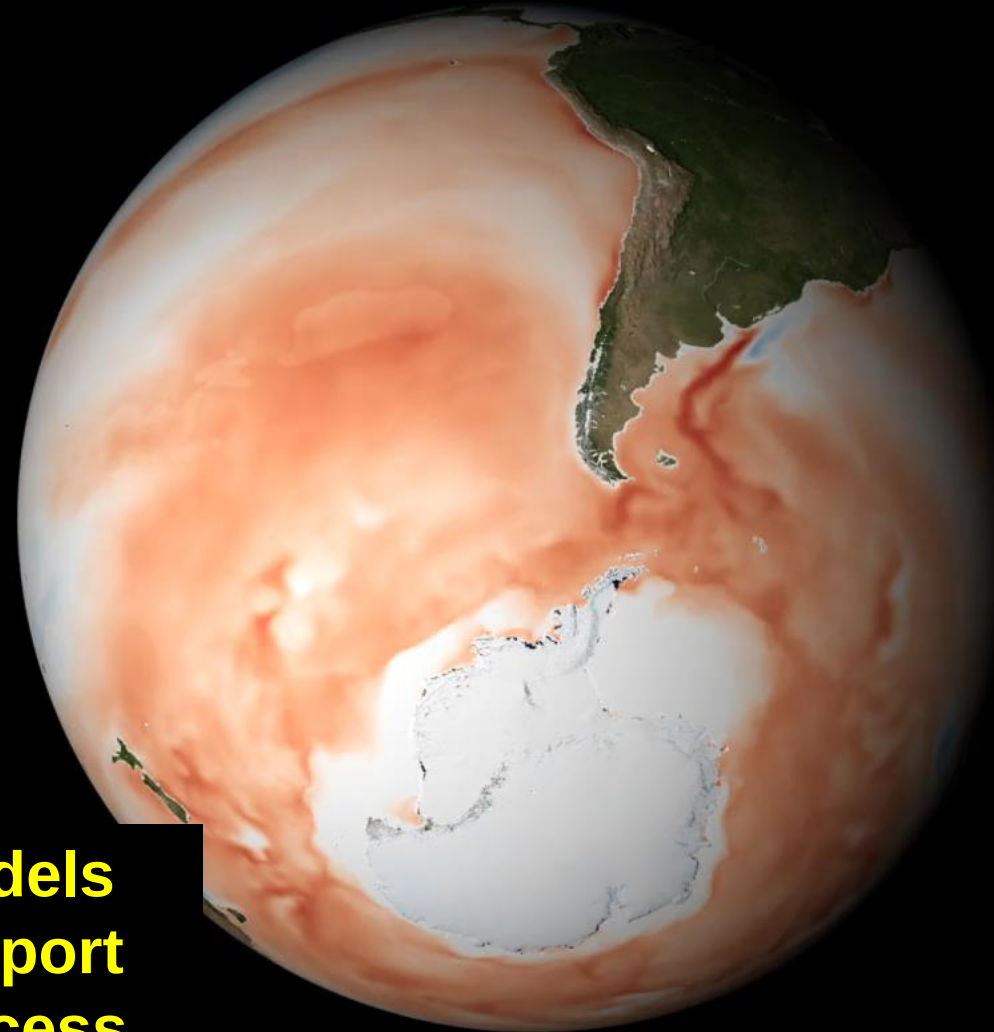
On a 3D grid, calculate circulation and carbon fluxes by integrating over time the equations of motion, conservation, biology and carbon chemistry

Moseley et al, in review

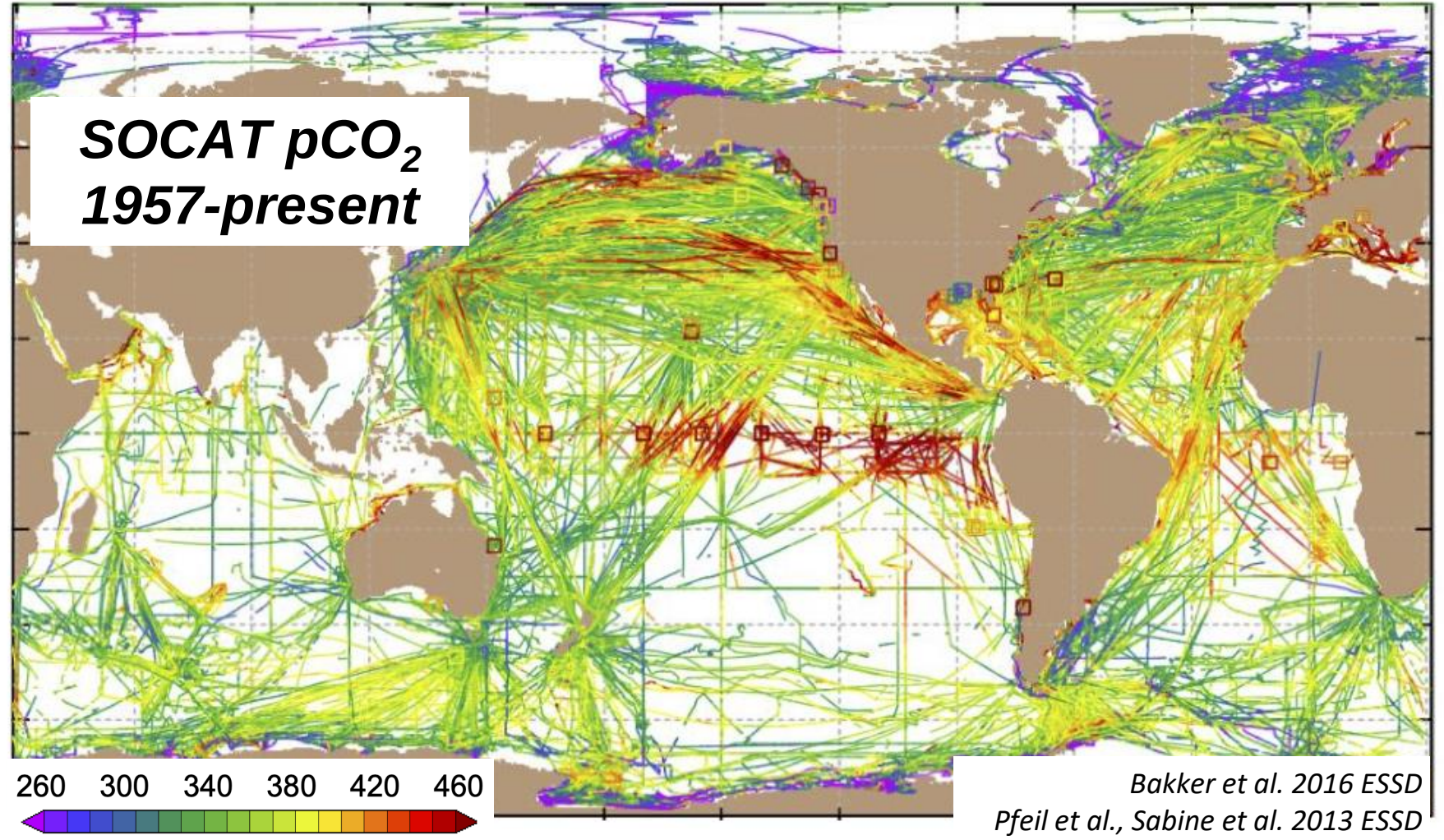
Heat CO₂ Flux

**Models
support
process
knowledge**

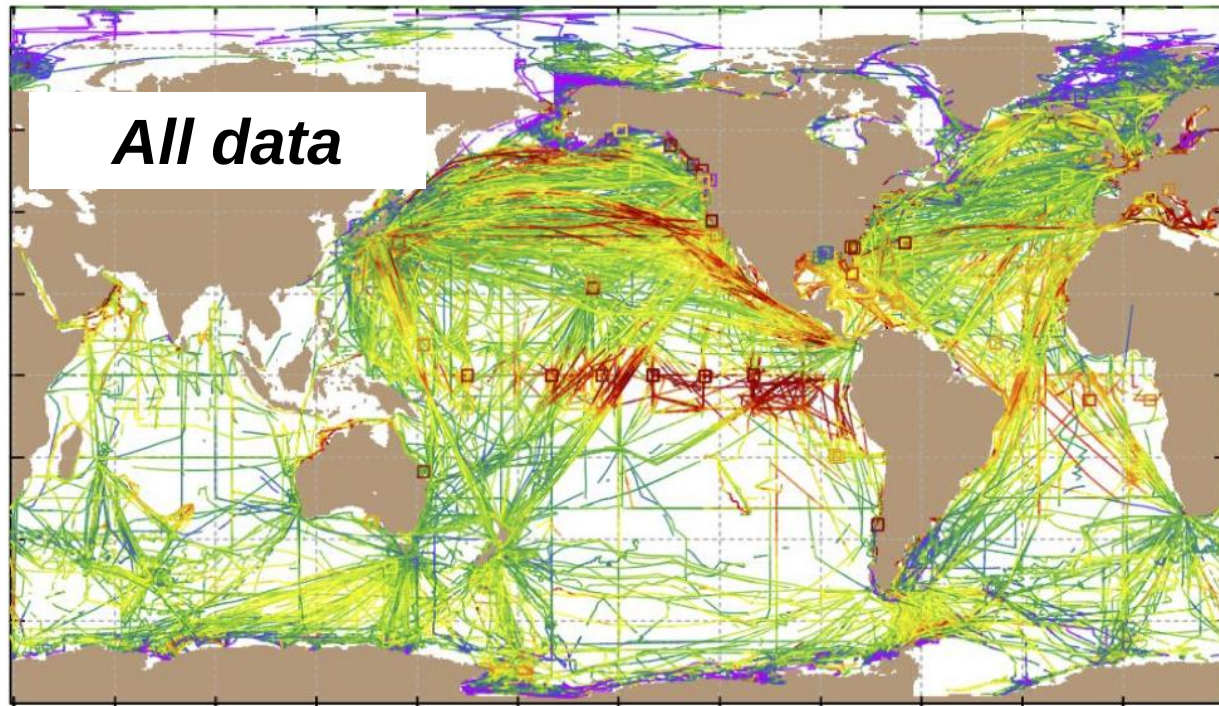
Lauderdale et al. 2016



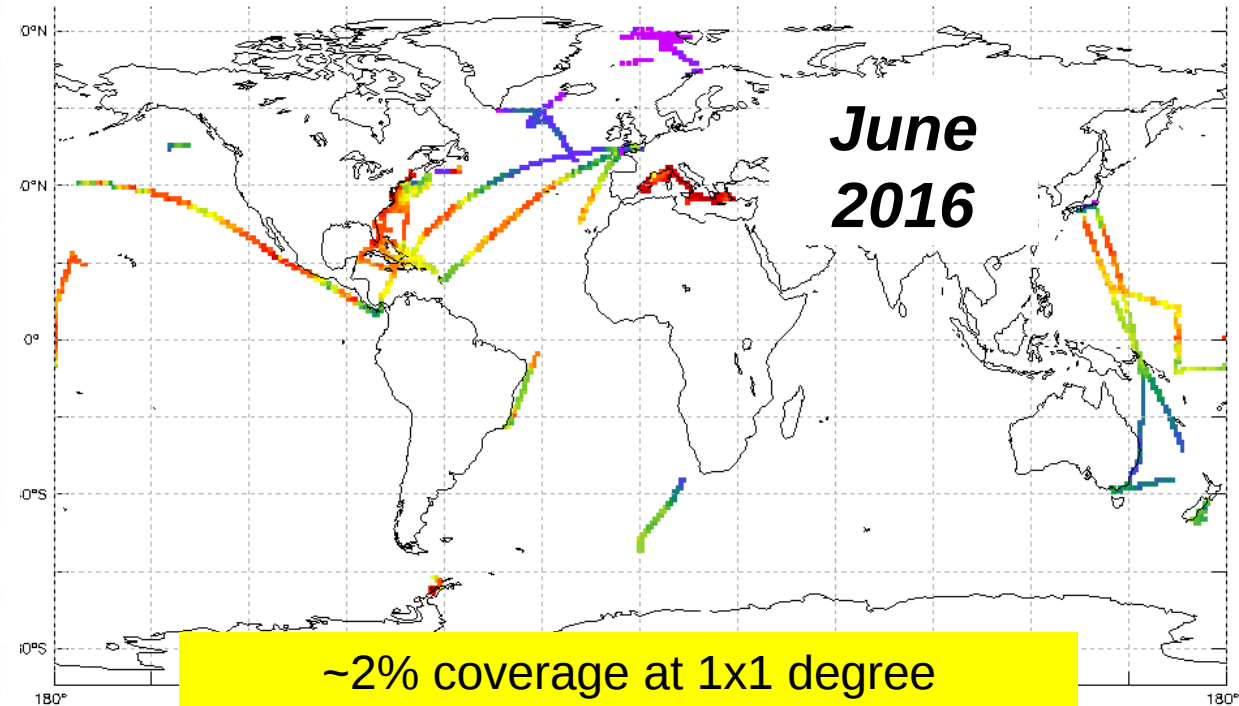
3. Surface ocean $p\text{CO}_2$ data used in reconstructions



3. Monthly $p\text{CO}_2$ data are very sparse ...

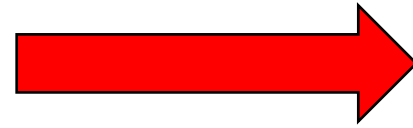
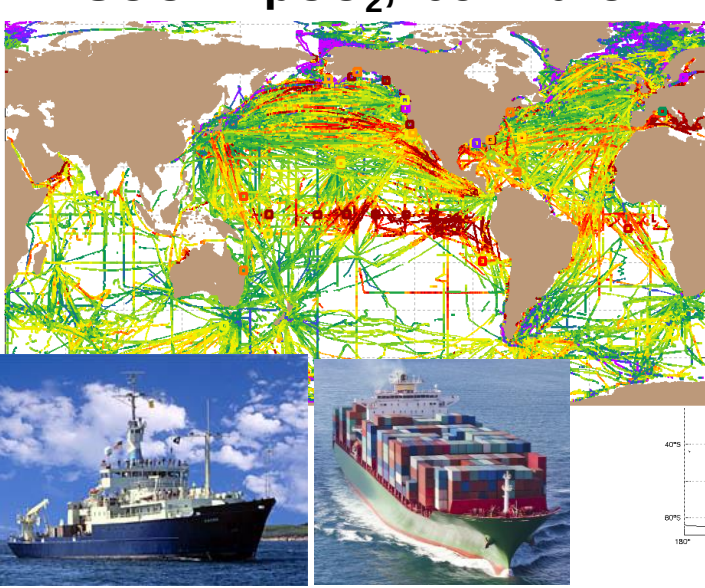


260 300 340 380 420 460

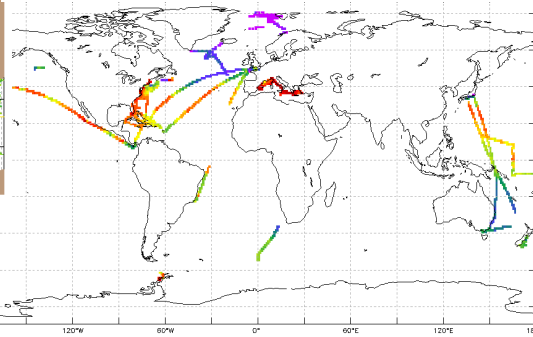


3. Data-based “Products” built with v. sparse $p\text{CO}_2$ data and ML

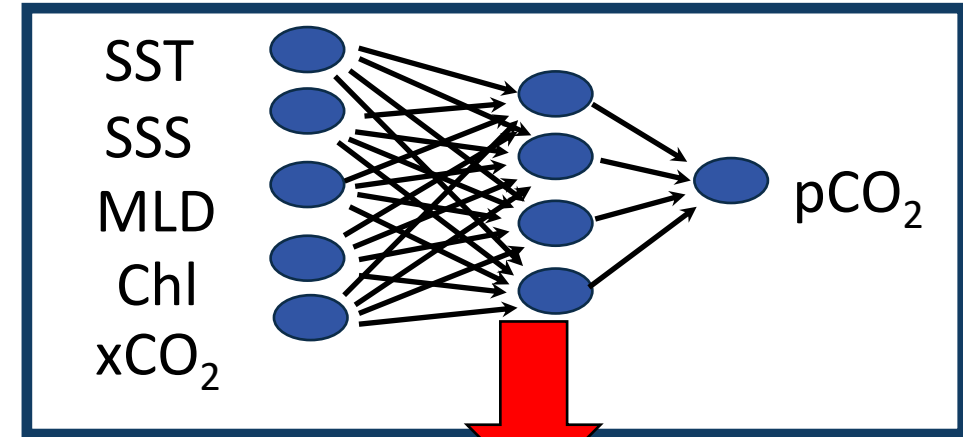
SOCAT $p\text{CO}_2$, 1957-2023



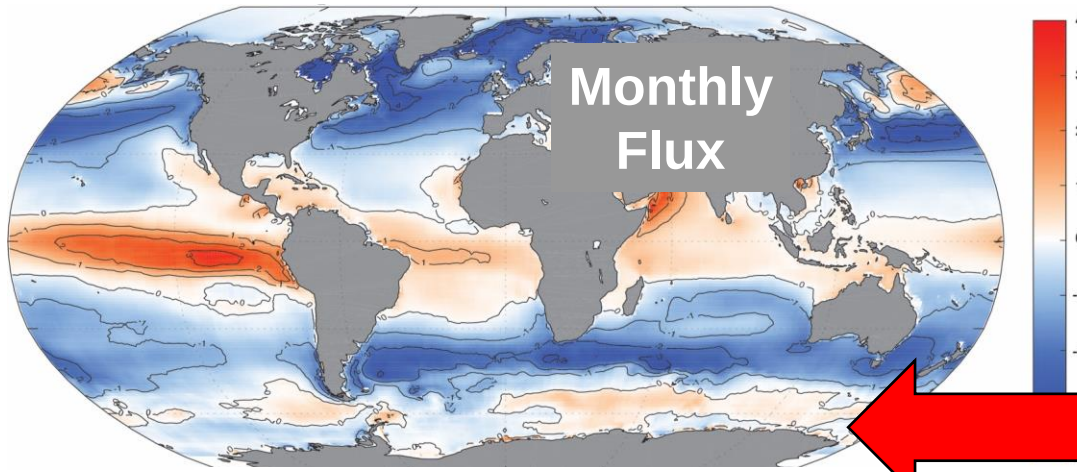
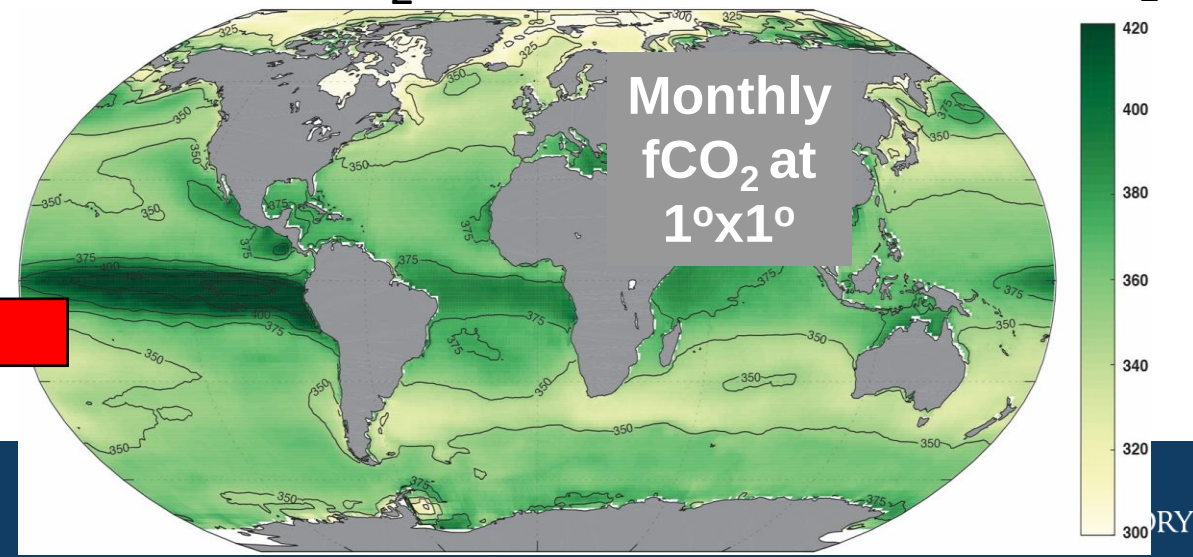
1 month ~ 2% coverage



1. Train Machine Learning on sparse data

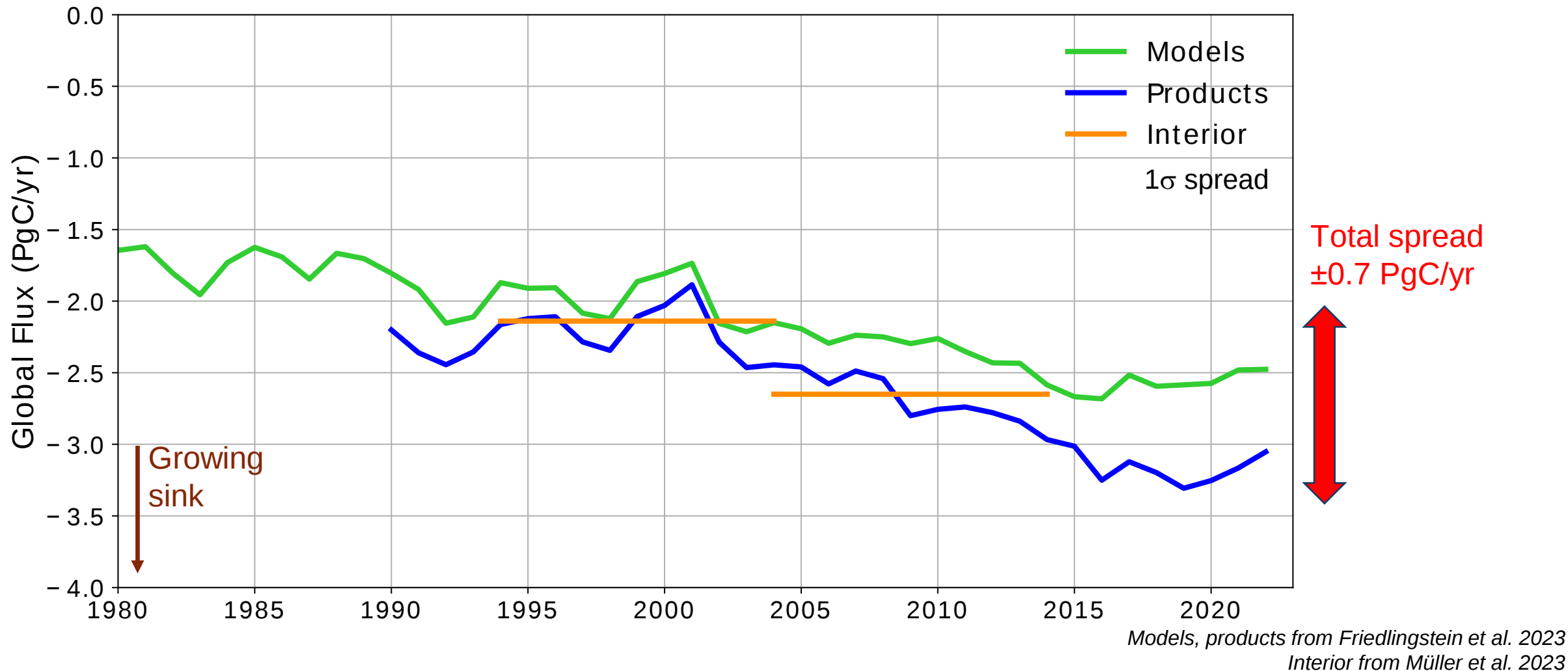


2. Predict $p\text{CO}_2 = f(\text{SST}, \text{SSS}, \text{MLD}, \text{Chl}, x\text{CO}_2)$



3. Calculate Flux from $p\text{CO}_2$

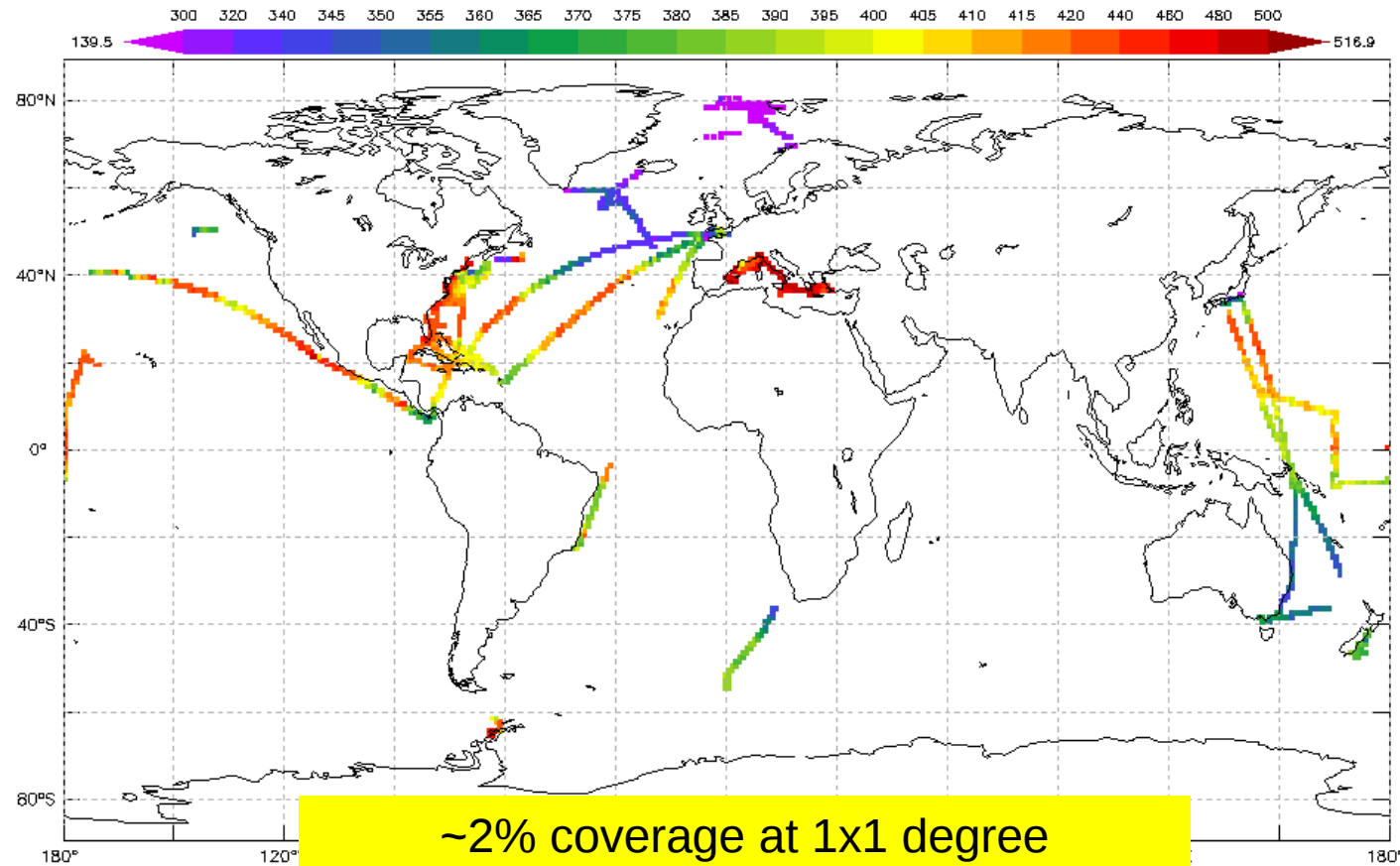
Globally integrated estimates agree, with ~30% uncertainty



Outline

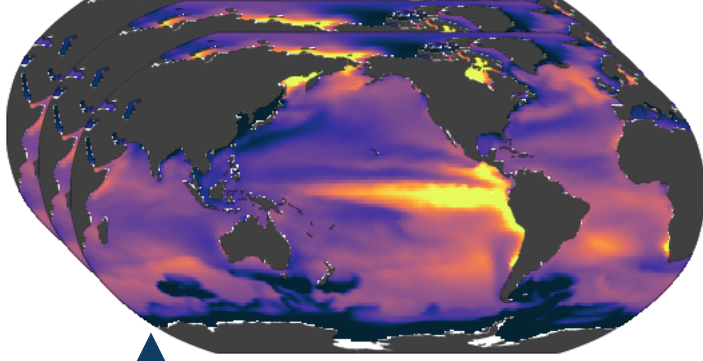
- Key processes of the ocean carbon cycle
- Improving quantification of the ocean carbon sink
 - **Models: evaluate ML-based data product skill, given sampling**
 - *pCO₂ data products*: apply to identify model mean-state biases
 - *Models and data*: combine using ML
 - *marine Carbon Dioxide Removal (mCDR)* – quantifying “additionality”?

pCO₂ data are very sparse: Are reconstructions robust?



ESM Large Ensembles as testbed for reconstruction skill

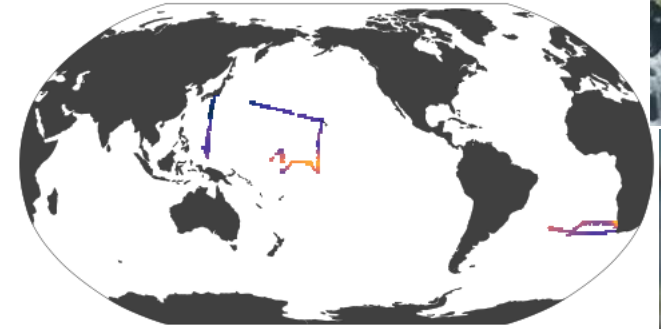
Earth System Models provide $p\text{CO}_2$ and drivers (SST, Chl, etc)



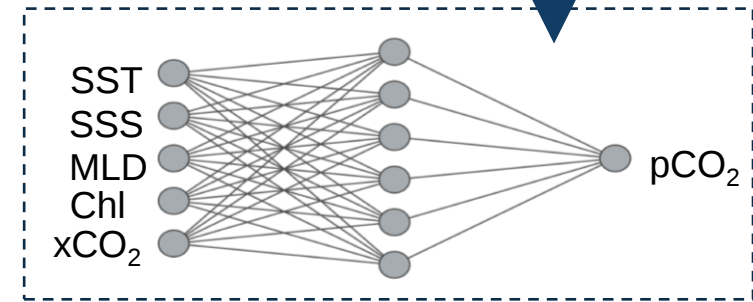
50 to 100 ensembles
from 4 or 9 ESMs



1. Sample model member as
SOCAT monthly $p\text{CO}_2$ product



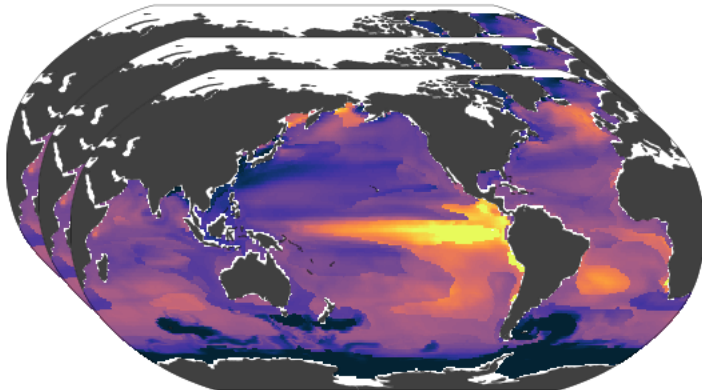
2. Train, evaluate, test



With various ML reconstructions

- MPI-SOMFFN (Gloege et al. 2021)
- NN, RF, XGB (Stamell et al. 2020)
- $p\text{CO}_2$ -Residual (Bennington et al. 2022, Heimdal et al, 2024a,b; Heimdal et al. in press)

4. Statistically compare reconstructed
 CO_2 flux to model truth. Each spatial
point is temporally decomposed



reconstruction

3. Estimate monthly
varying $p\text{CO}_2$ on global
scale using trained
model, calculate flux



X ensemble
members



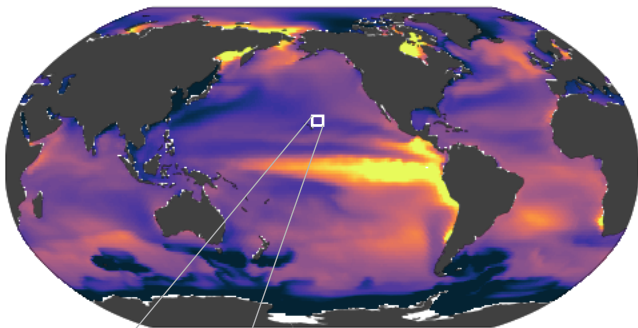
Luke Gloege



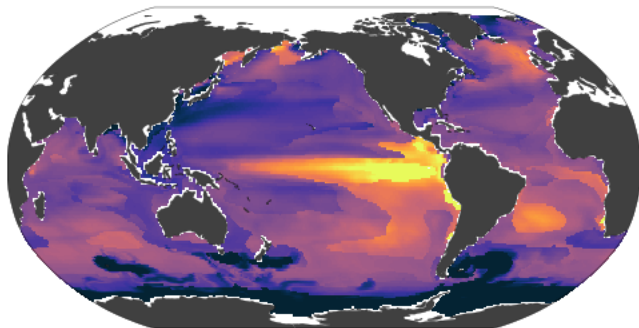
Thea Heimdal

Skill evaluation: temporal decomposition

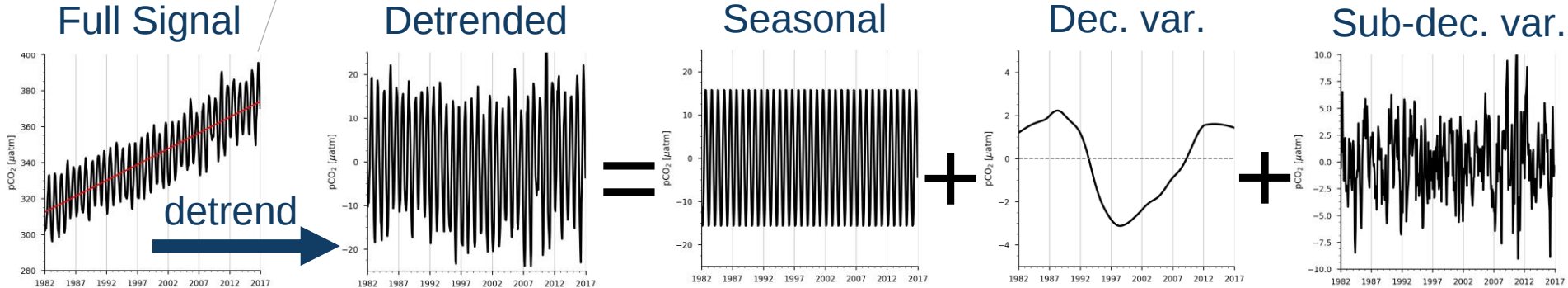
Ensemble member



Reconstruction



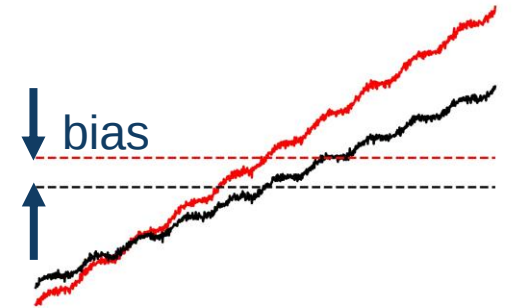
Each location in each member and reconstruction decomposed into seasonal, decadal, and sub-decadal variability



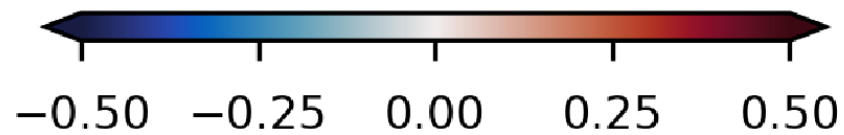
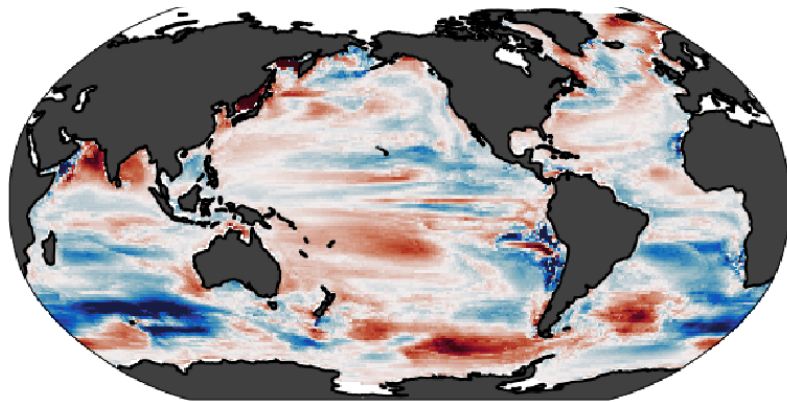
Gloeger et al. 2021, GBC

Bias is small globally; larger where data sparse

Model Reconstruction



Global avg = $-0.01 \text{ mol C m}^{-2} \text{ yr}^{-1}$

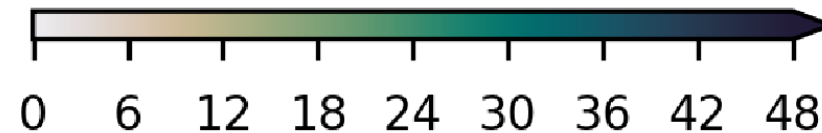
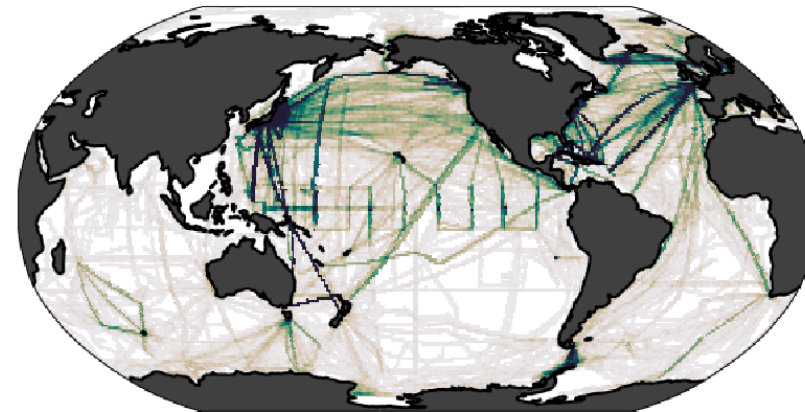


bias [$\text{mol C m}^{-2} \text{ yr}^{-1}$]

100-member mean

1985-2016

MPI-SOMFFN



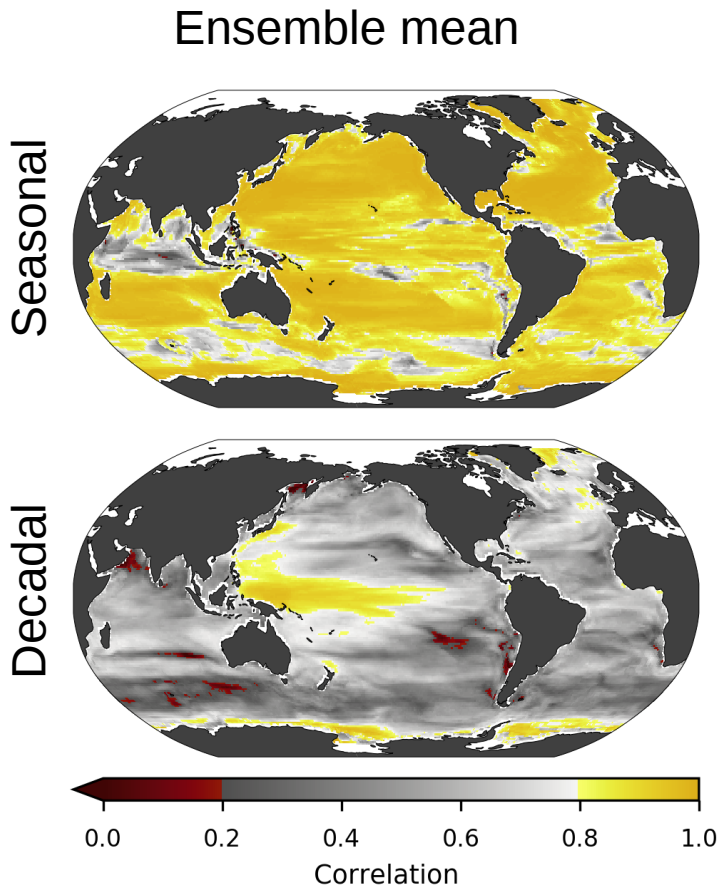
number of months with data

Gloeger et al. 2021, GBC

Gloeger et al. 2021, GBC

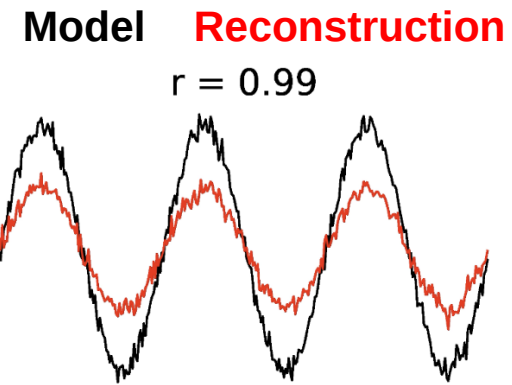
Correlation

Is the reconstruction in phase with the original data?



- Captures seasonal cycle well everywhere

- Decadal phase is more challenging to reconstruct



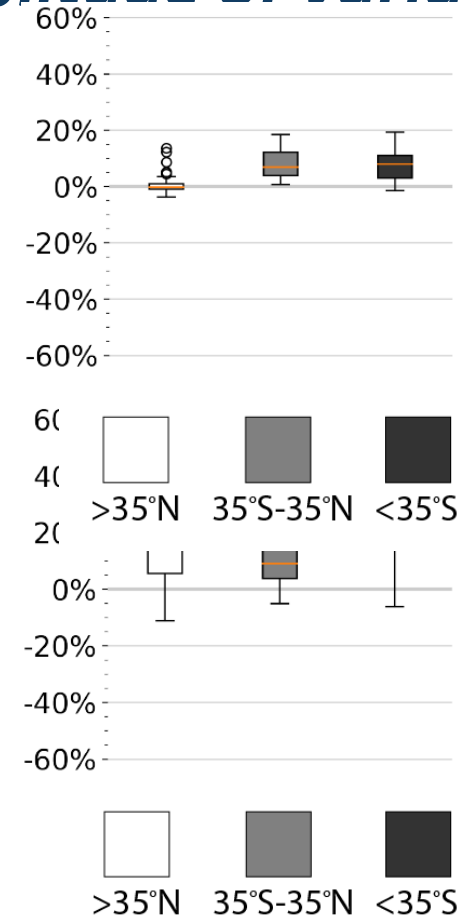
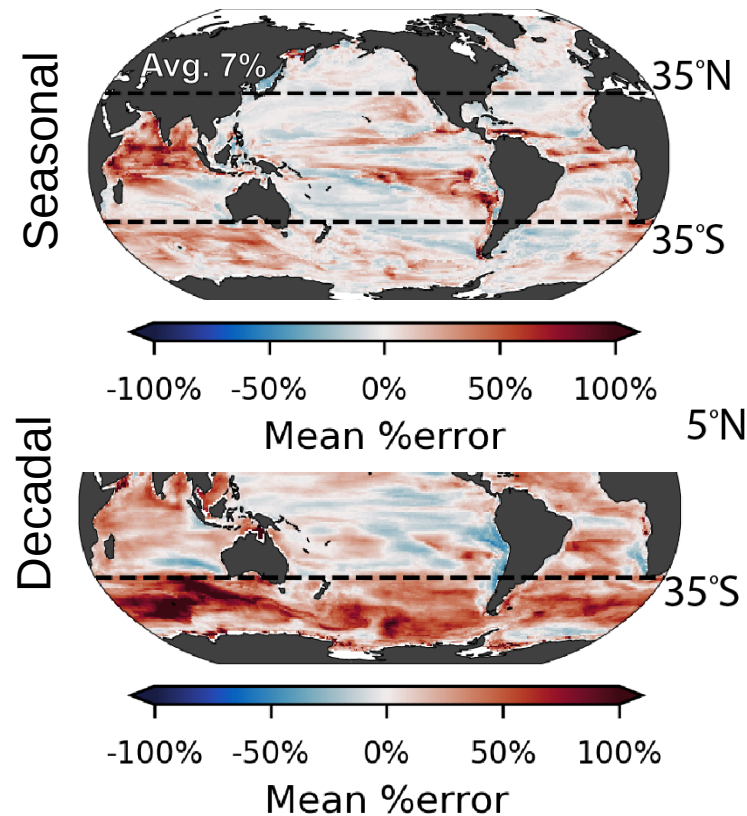
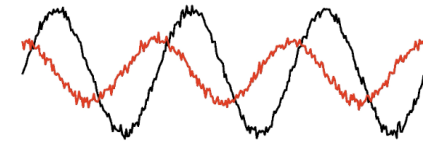
100-member mean
1985-2016
MPI-SOMFFN

Gloeger et al. 2021, GBC

Standard deviation of % error

Does the reconstruction capture the amplitude of variability?

$$\sigma^* = -50.1\%$$

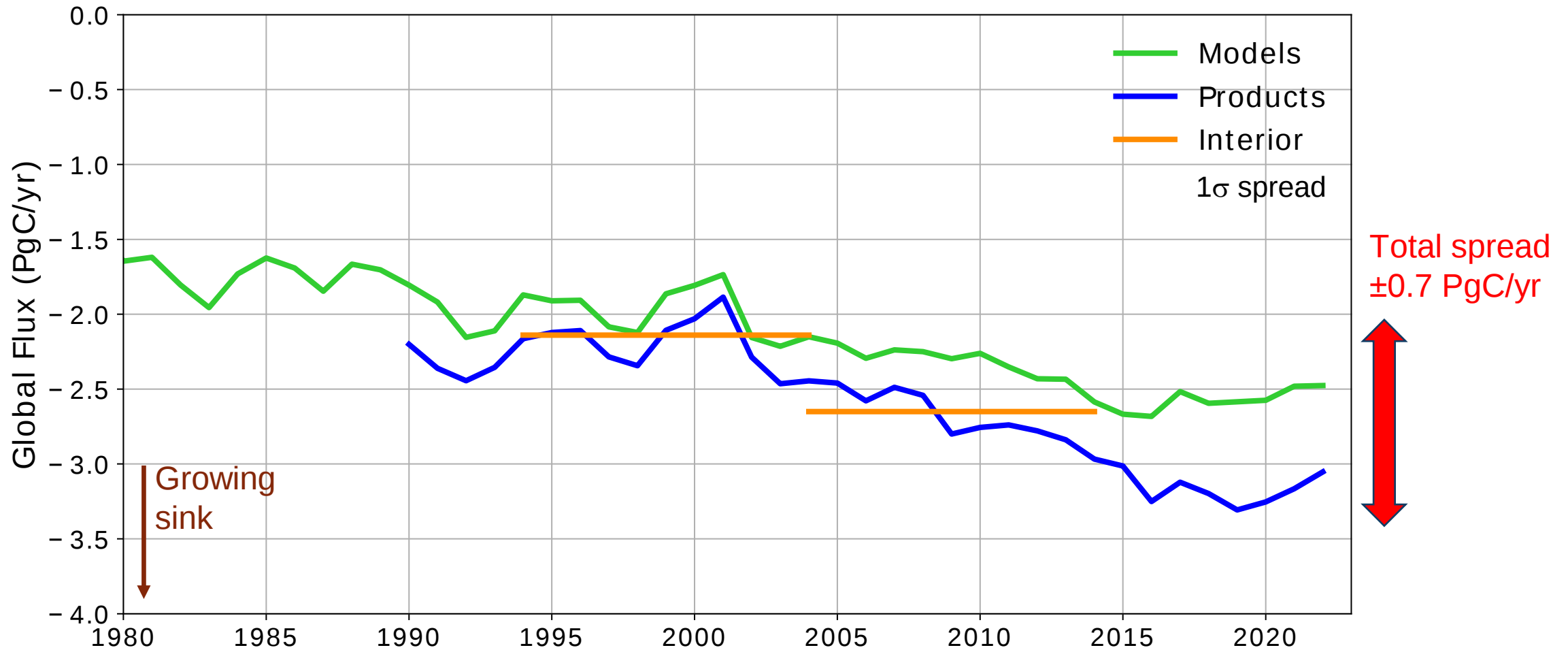


100-member mean
1985-2016
MPI-SOMFFN

- Globally the seasonal cycle is overestimated by ~7%
- Overestimates decadal variability in Southern Ocean by ~39%

Gloege et al. GBC 2021

Reduced decadal amplitude in products should increase agreement between products and models



Models, products from Friedlingstein et al. 2023
Interior from Müller et al. 2023

Take home messages

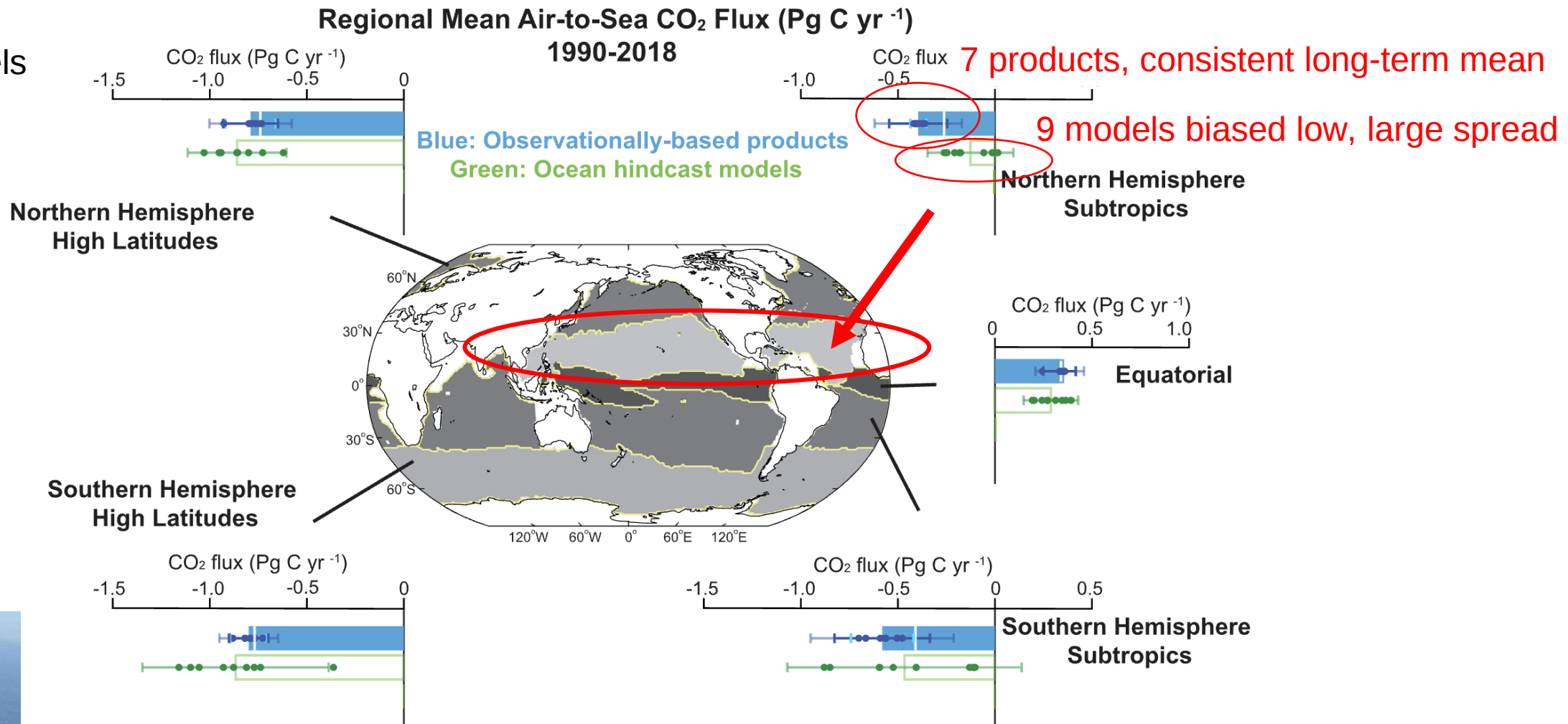
- Air-sea CO₂ flux mean and seasonality can be reconstructed from sparse pCO₂ data, but longer-timescale variability remains uncertain

Outline

- Key processes of the ocean carbon cycle
- Improving quantification of the ocean carbon sink
 - *Models*: evaluate ML-based data product skill, given sampling
 - ***pCO₂ data products: apply to identify model mean-state biases***
 - *Models and data*: combine using ML
 - *marine Carbon Dioxide Removal (mCDR)* – quantifying “additionality”?

Ocean models have significant mean-state errors

Products and hindcast models for real fluxes 1990-2018



Fay and McKinley 2021

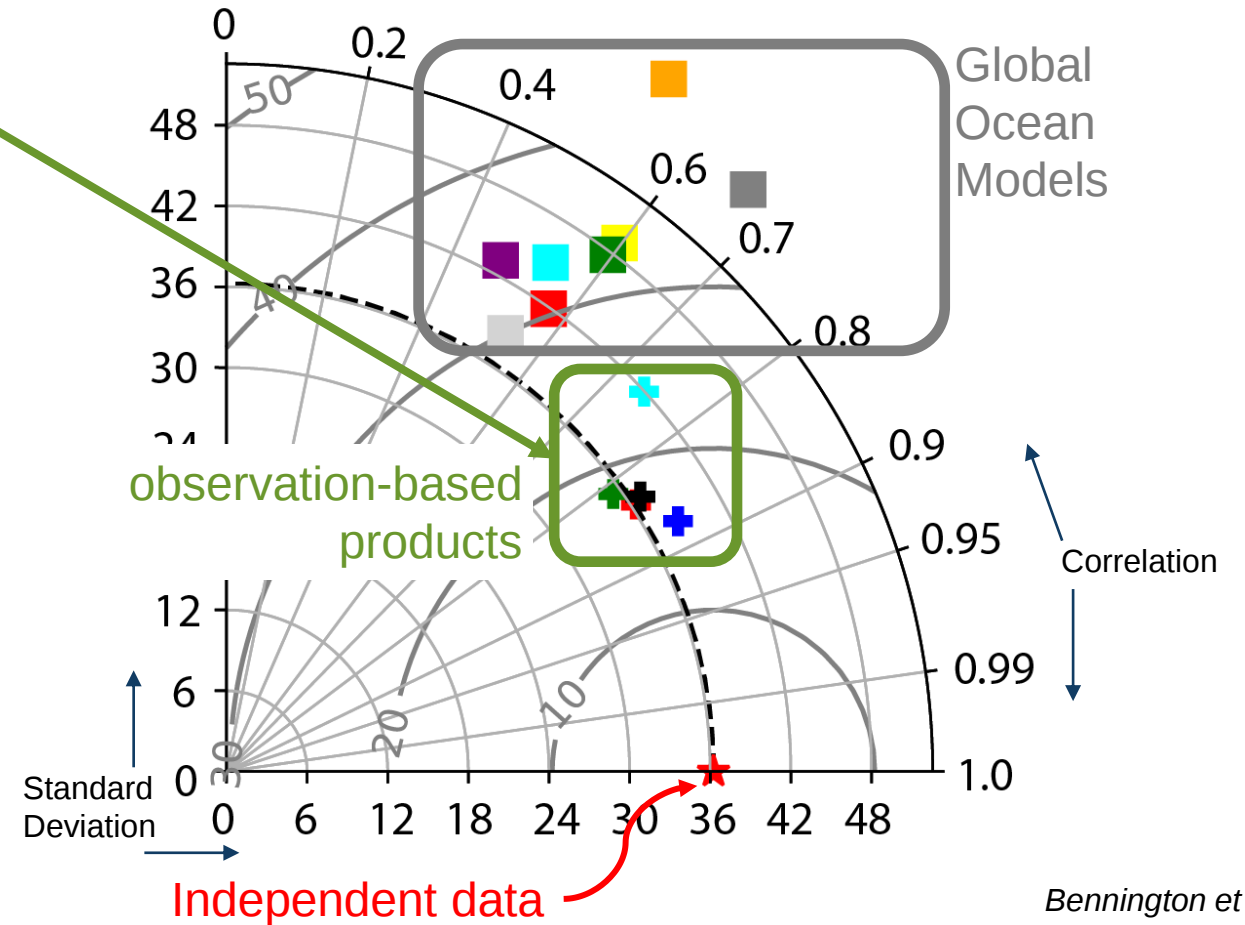
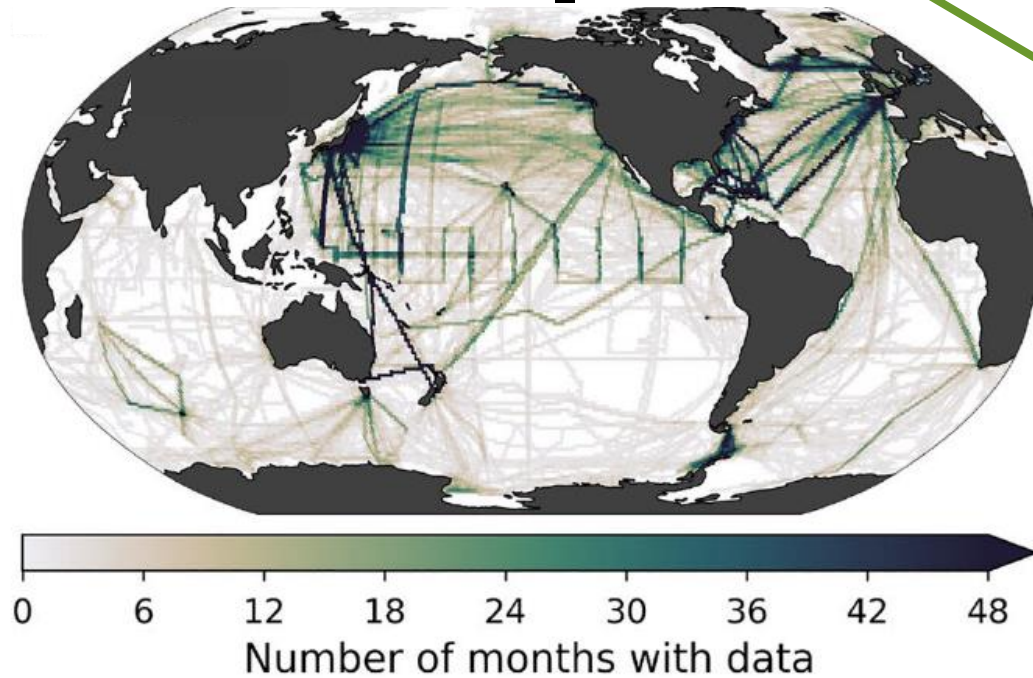


Amanda Fay

Ocean models have significant mean-state errors

Model + product pCO₂ validation with GLODAP

Surface ocean pCO₂ data (SOCAT)

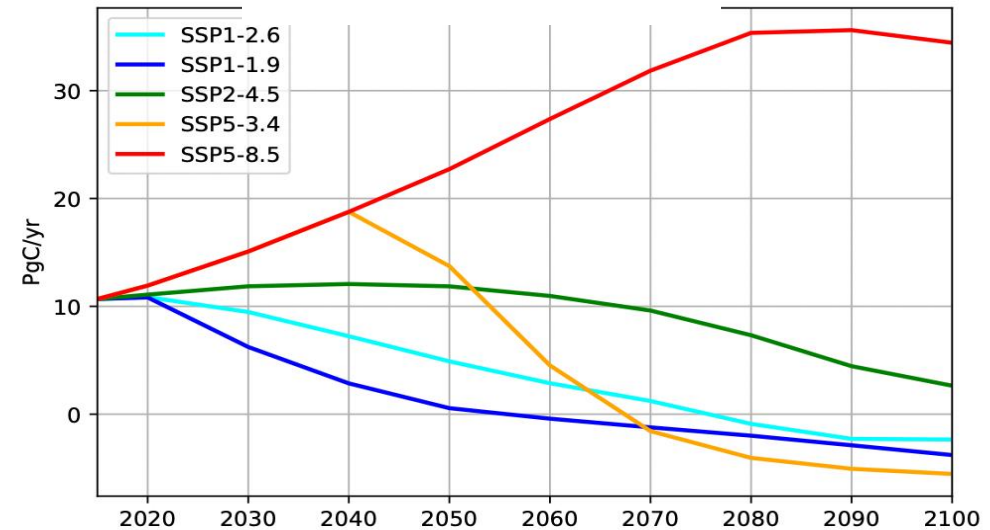


Bennington et al., 2022
Gloege et al. 2021, Fay and McKinley 2021

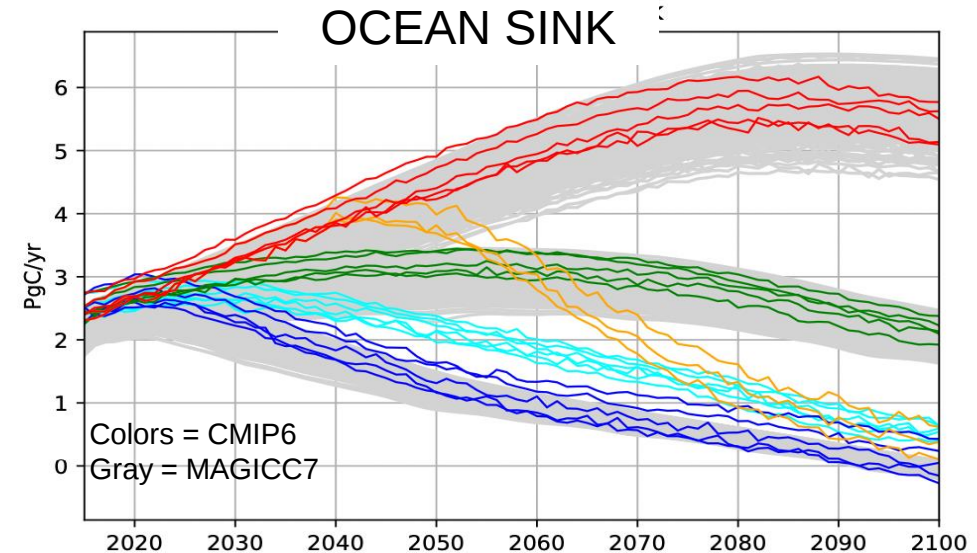
Mean-state biases are retained in CMIP6 projections

Future Emissions and Ocean Sink

EMISSIONS



OCEAN SINK



CMIP6

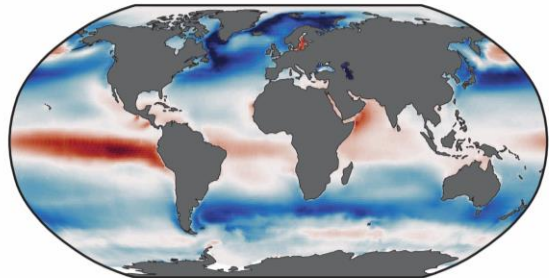
3-6 ensemble means
for each scenario

McKinley et al. 2023, ERL

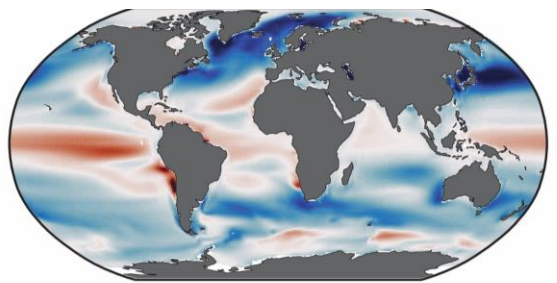
Mean-state biases are retained in CMIP6 projections

Modern = 2010-2020 Flux

(a) Observation-based
products 2010s

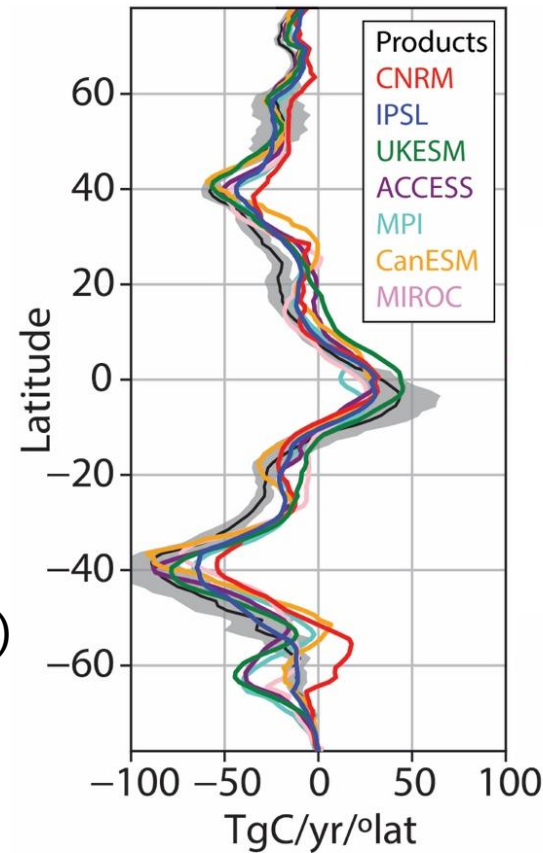


Example CMIP6 (IPSL)



→
Zonal average of this
comparison:
7 models (colors)
8 products (black/gray)

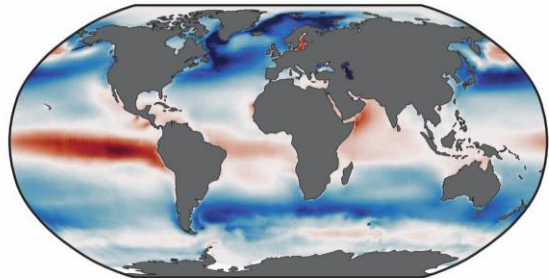
Modern Zonal-mean



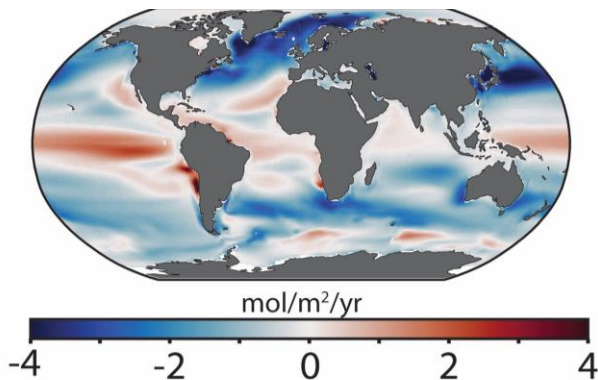
Mean-state biases are retained in CMIP6 projections

Modern = 2010-2020 Flux

(a) Observation-based products 2010s

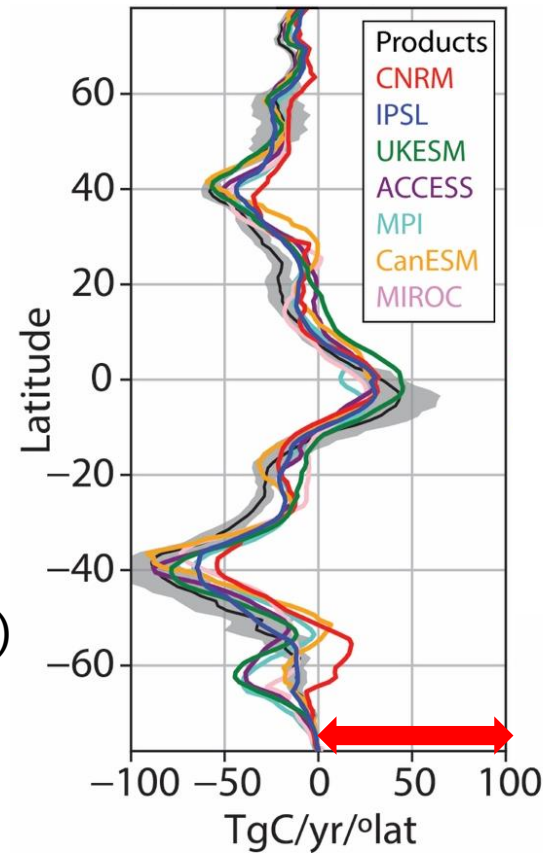


Example CMIP6 (IPSL)



Zonal average of this comparison:
7 models (colors)
8 products (black/gray)

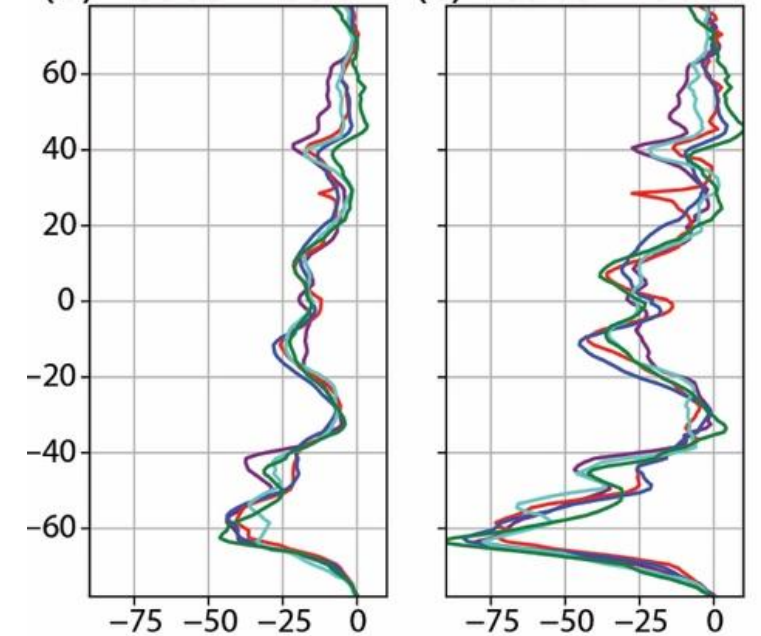
Modern Zonal-mean



Future Flux change

SSP5-8.5

(b) 2050s-Modern (e) 2090s-Modern

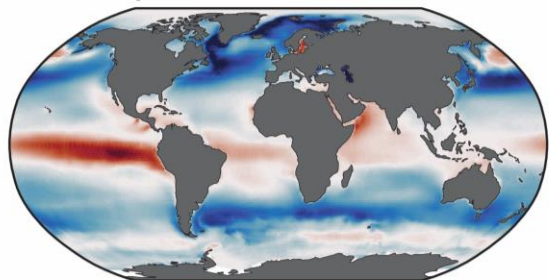


Note: expanded scale

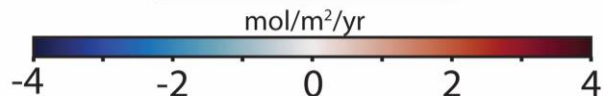
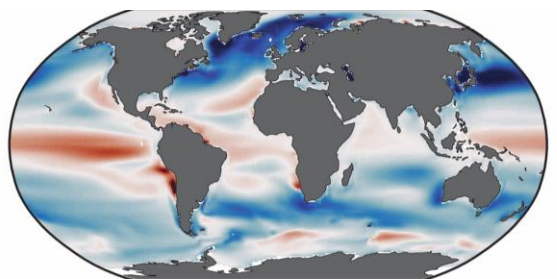
Mean-state biases are retained in CMIP6 projections

Modern = 2010-2020 Flux

(a) Observation-based products 2010s

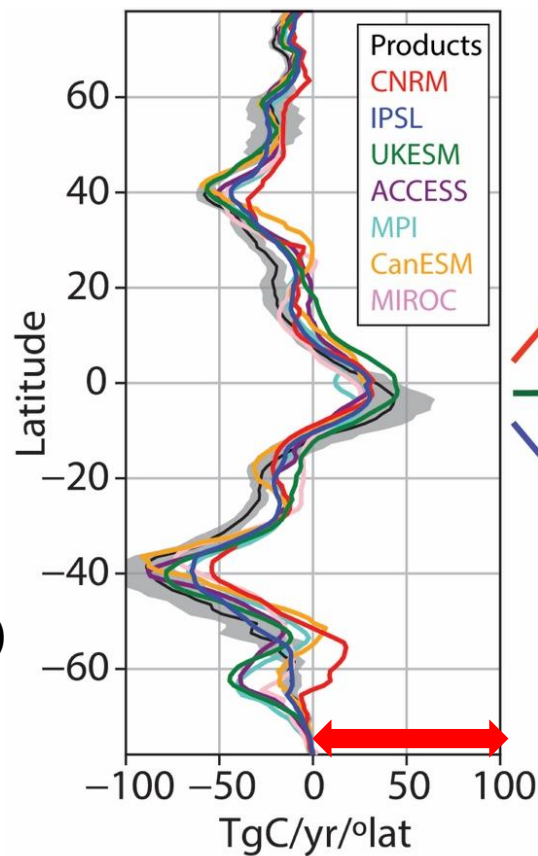


Example CMIP6 (IPSL)



Zonal average of this comparison:
7 models (colors)
8 products (black/gray)

Modern Zonal-mean

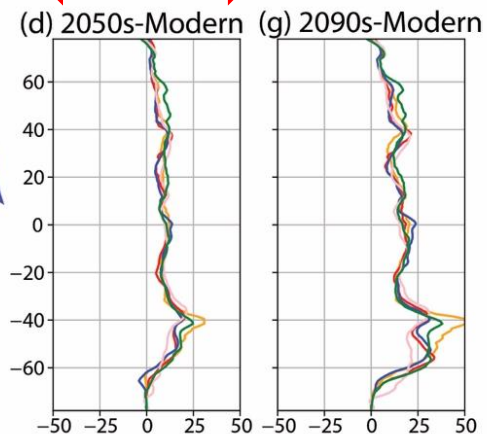
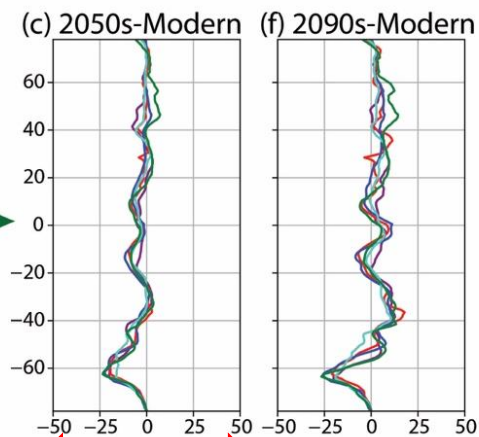
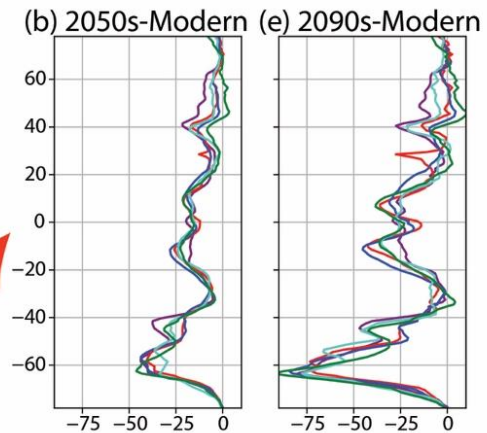


SSP5-8.5

SSP2-4.5

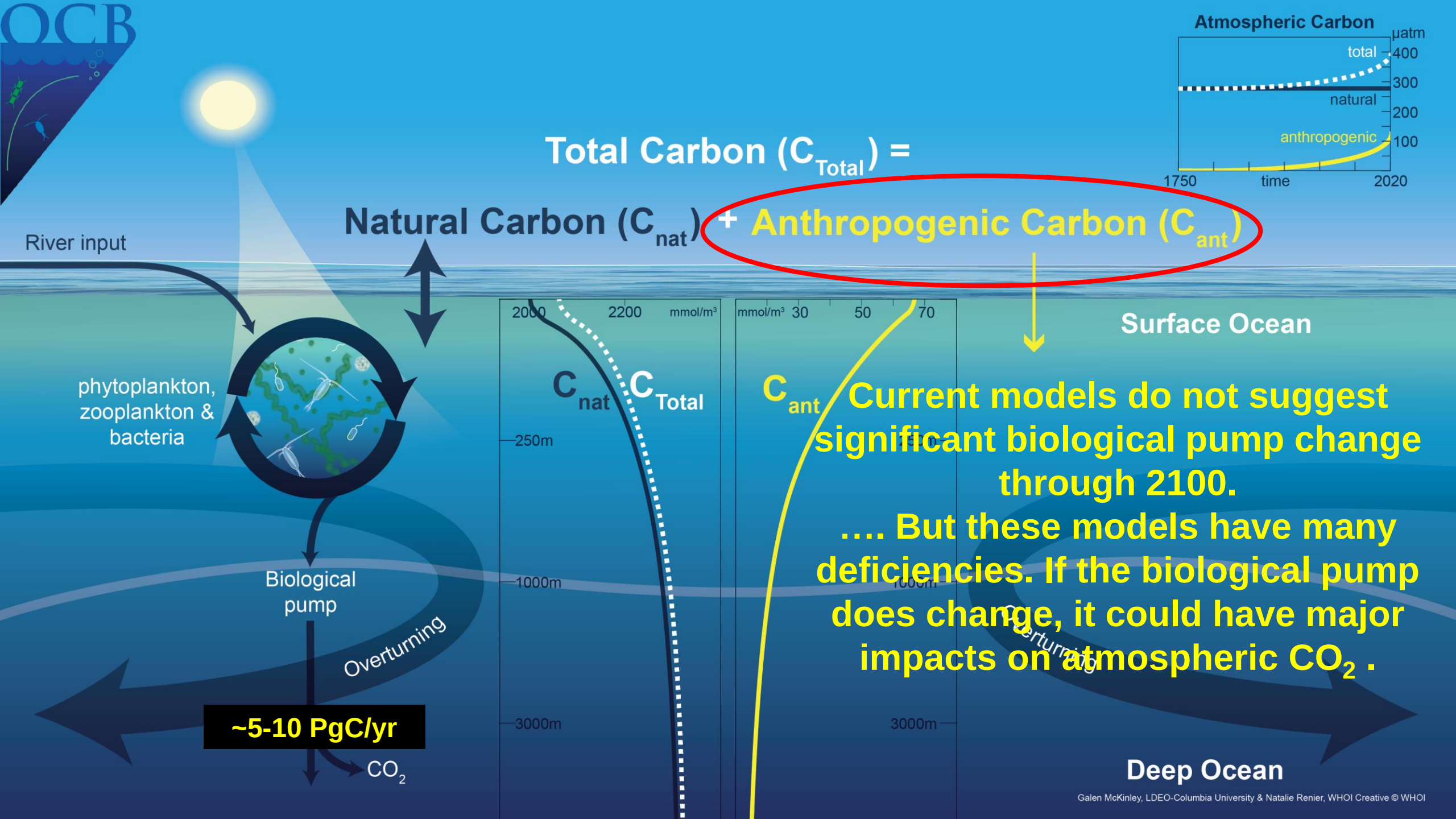
SSP1-1.9

Future Flux change



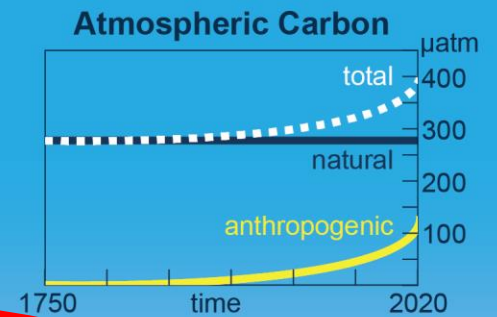
Take home messages

- Air-sea CO₂ flux mean and seasonality can be reconstructed from sparse pCO₂ data, but longer-timescale variability remains uncertain
- Hindcast and Earth System Models have significant mean-state biases



Total Carbon (C_{Total}) =

Natural Carbon (C_{nat}) + Anthropogenic Carbon (C_{ant})



River input

phytoplankton,
zooplankton &
bacteria

Biological
pump

Overturning

$\sim 5\text{-}10 \text{ PgC/yr}$

CO_2

2000 2200 mmol/m^3 mmol/m^3 30 50 70

C_{nat} C_{Total}

C_{ant}

Surface Ocean

Current models do not suggest
significant biological pump change
through 2100.

.... But these models have many
deficiencies. If the biological pump
does change, it could have major
impacts on atmospheric CO_2 .

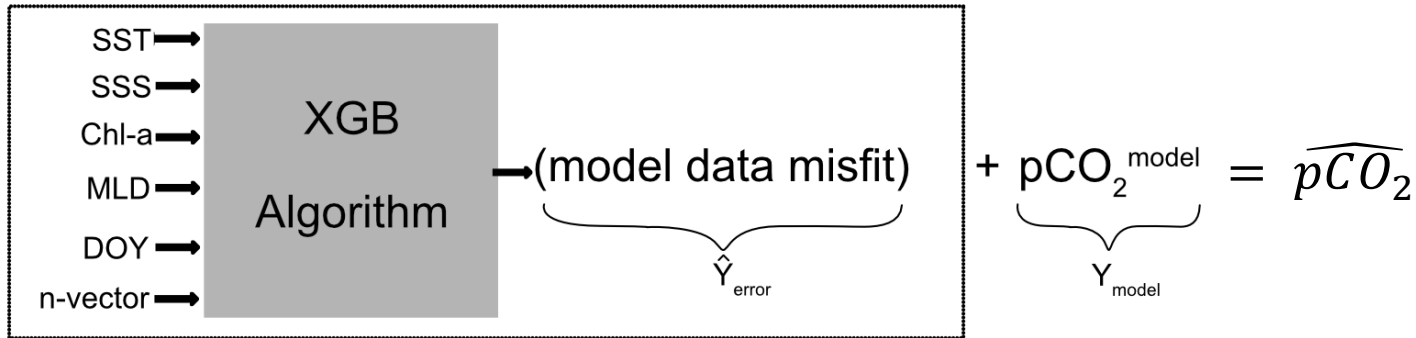
Deep Ocean

Outline

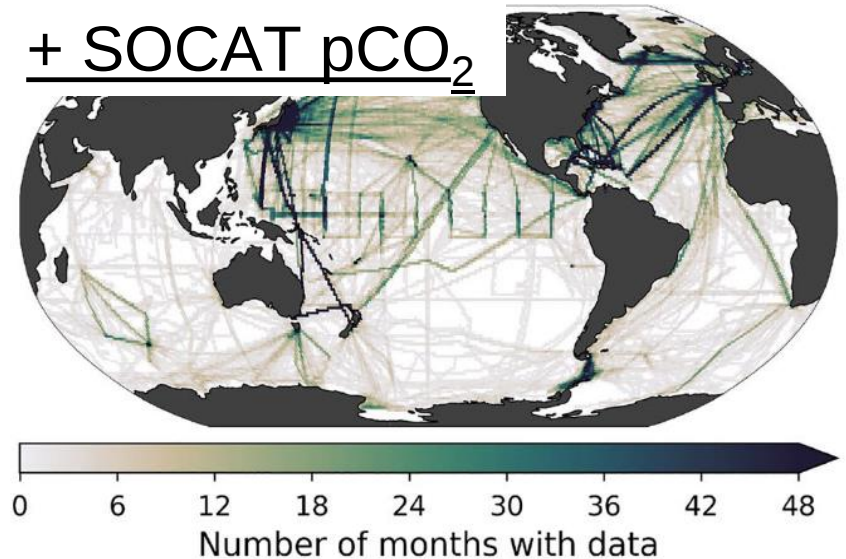
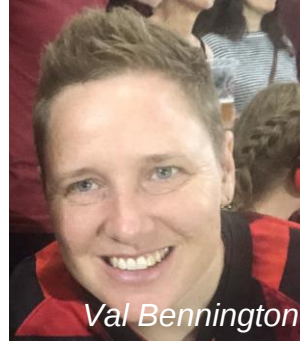
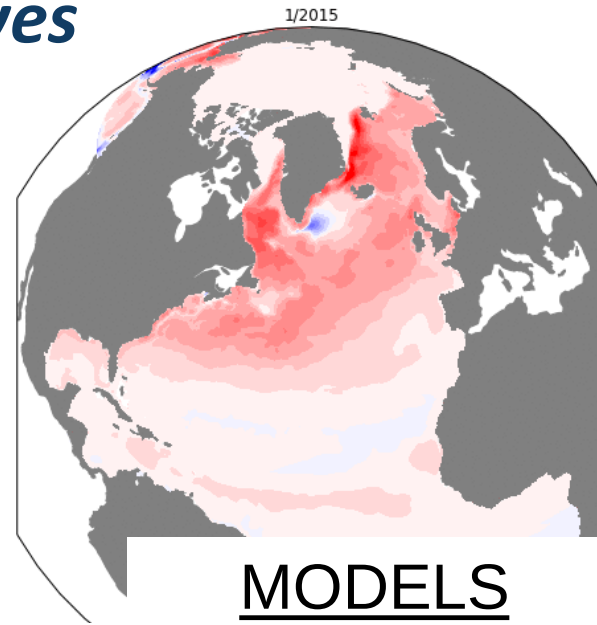
- Key processes of the ocean carbon cycle
- Improving quantification of the ocean carbon sink
 - *Models*: evaluate ML-based data product skill, given sampling
 - *pCO₂ data products*: apply to identify model mean-state biases
 - ***Models and data: combine using ML***
 - *marine Carbon Dioxide Removal (mCDR)* – quantifying “additionality”?

Using ML to combine models and data improves surface flux estimates (LDEO-HPD)

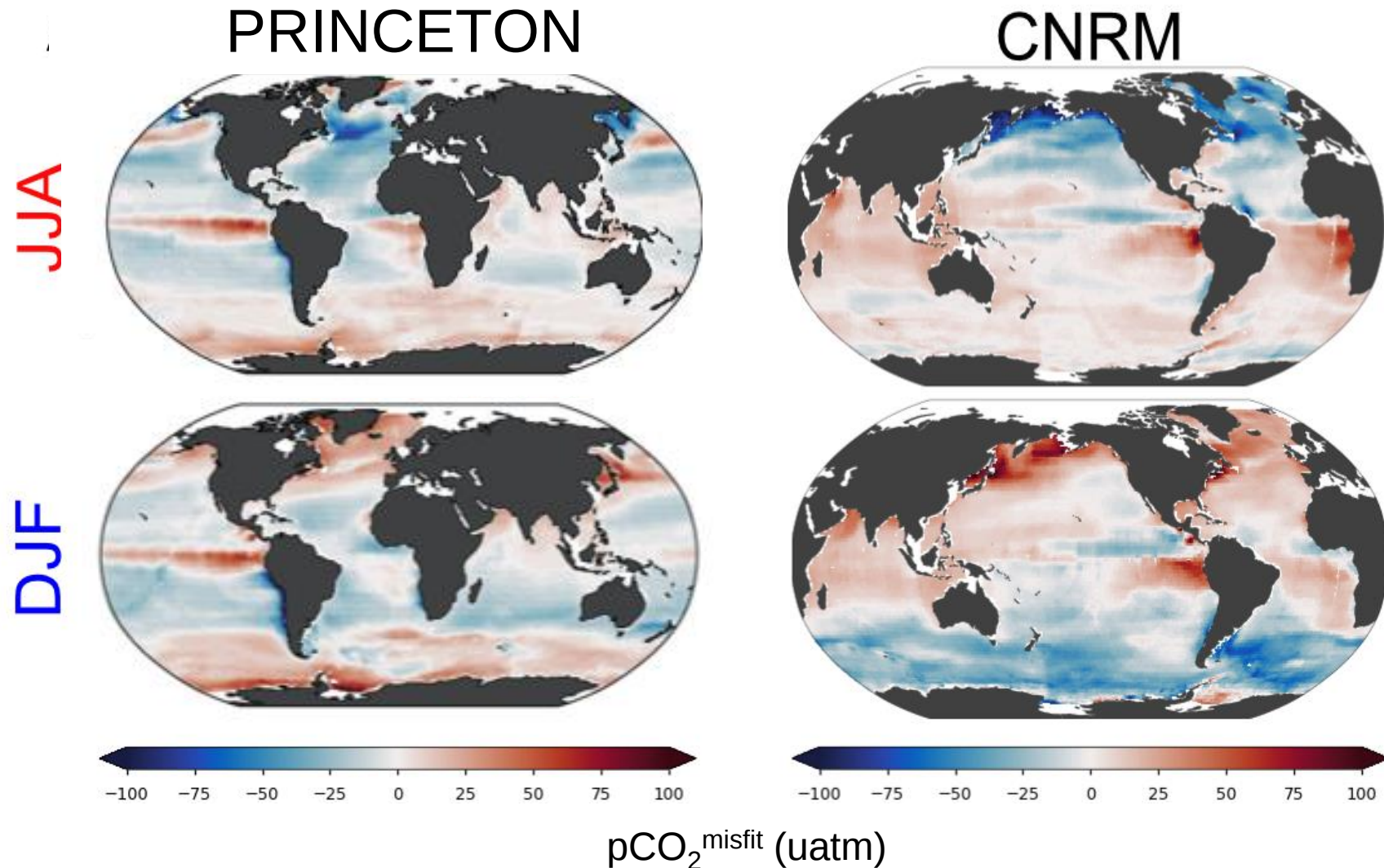
- $p\text{CO}_2^{\text{model}}$ as priors
- Reconstruct $p\text{CO}_2^{\text{misfit}} = (p\text{CO}_2^{\text{model}} - p\text{CO}_2^{\text{SOCAT}})$
(driver data: SST, SSS, Chl-a, MLD, time, location)
- eXtreme Gradient Boost (XGB) algorithm
- Misfits to correct N models; average for final $p\text{CO}_2$



Gloege et al. 2022 JAMES (N=9), Bennington et al. 2022 GRL (N=8)



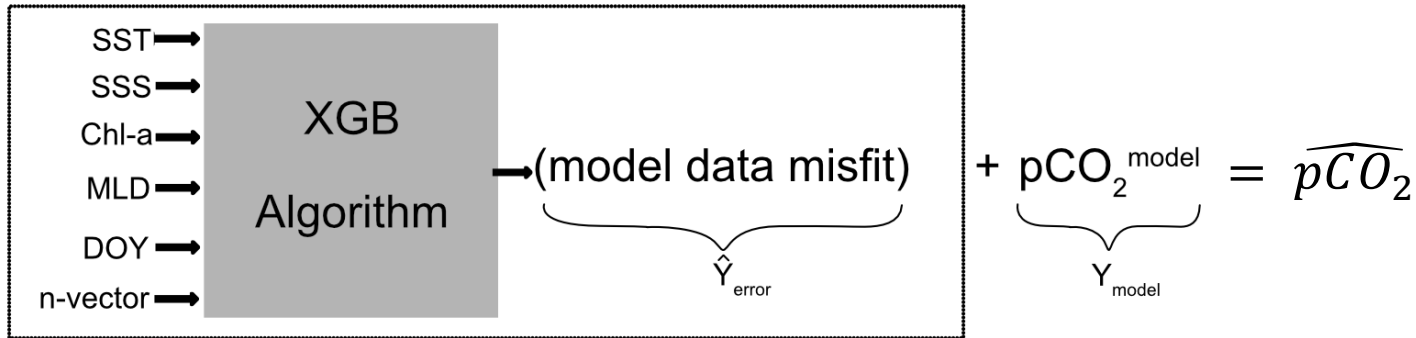
Full-coverage, monthly misfits illustrate substantial model biases



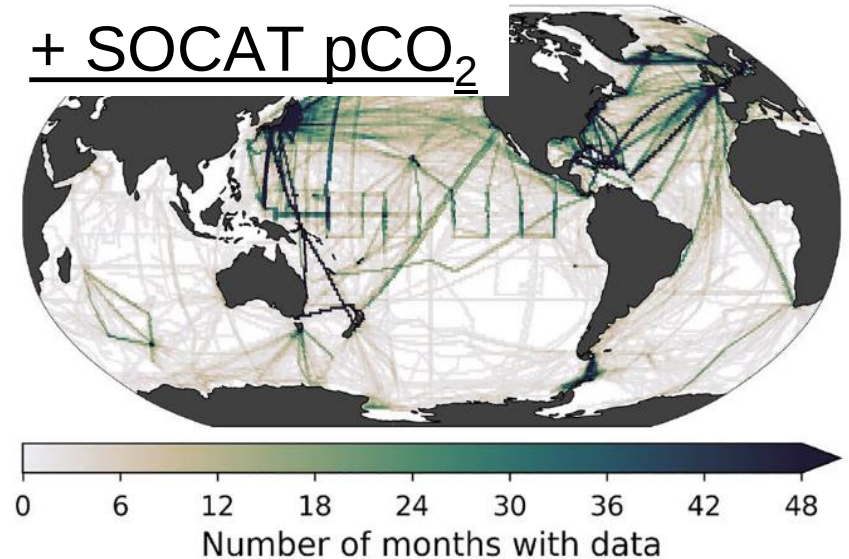
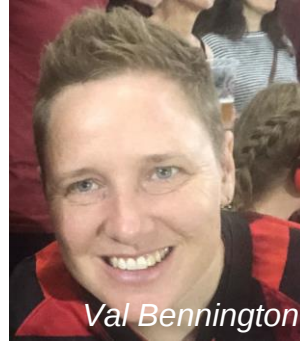
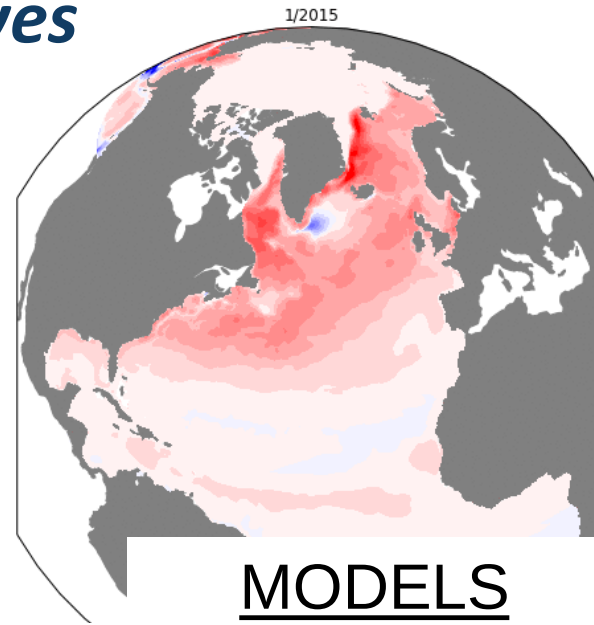
Gloege et al. 2022, JAMES

Using ML to combine models and data improves surface flux estimates (LDEO-HPD)

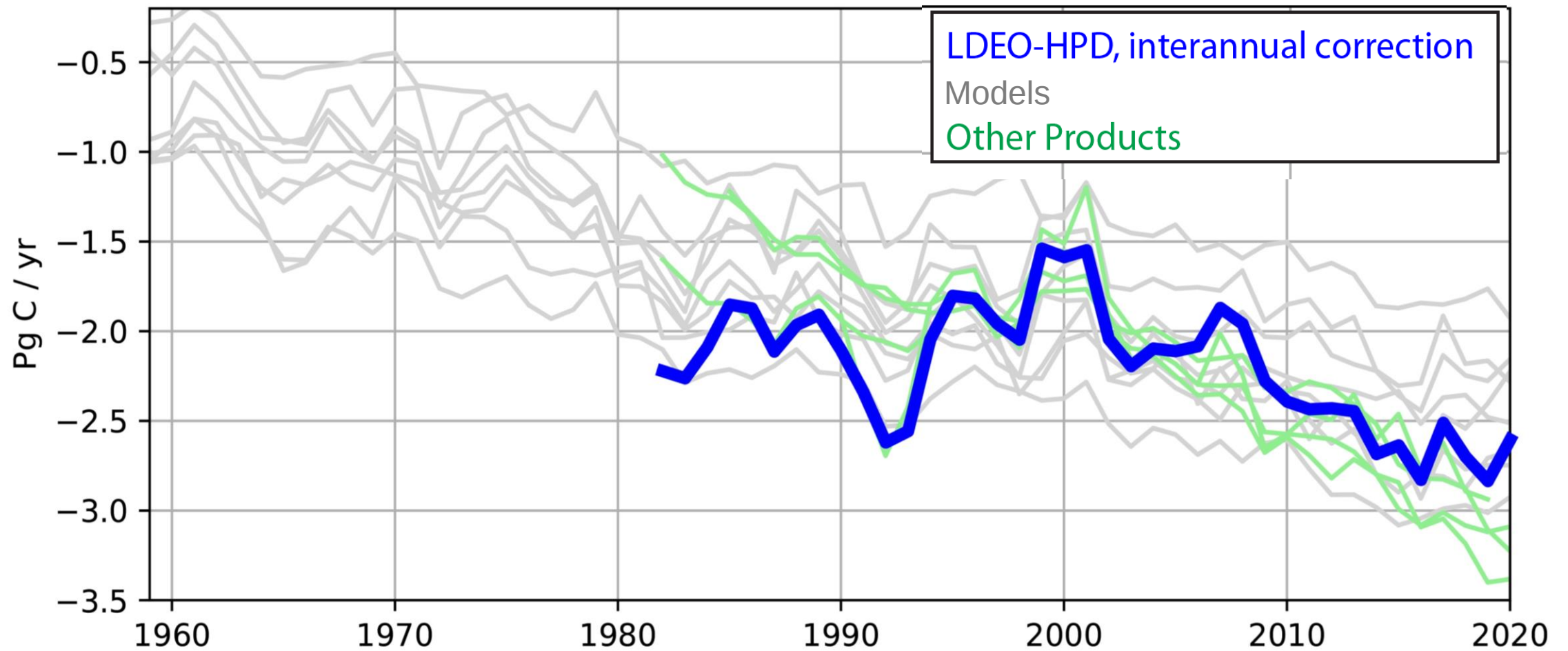
- $p\text{CO}_2^{\text{model}}$ as priors
- Reconstruct $p\text{CO}_2^{\text{misfit}} = (p\text{CO}_2^{\text{model}} - p\text{CO}_2^{\text{SOCAT}})$
(driver data: SST, SSS, Chl-a, MLD, time, location)
- eXtreme Gradient Boost (XGB) algorithm
- **Misfits to correct N models; average for final $p\text{CO}_2$**



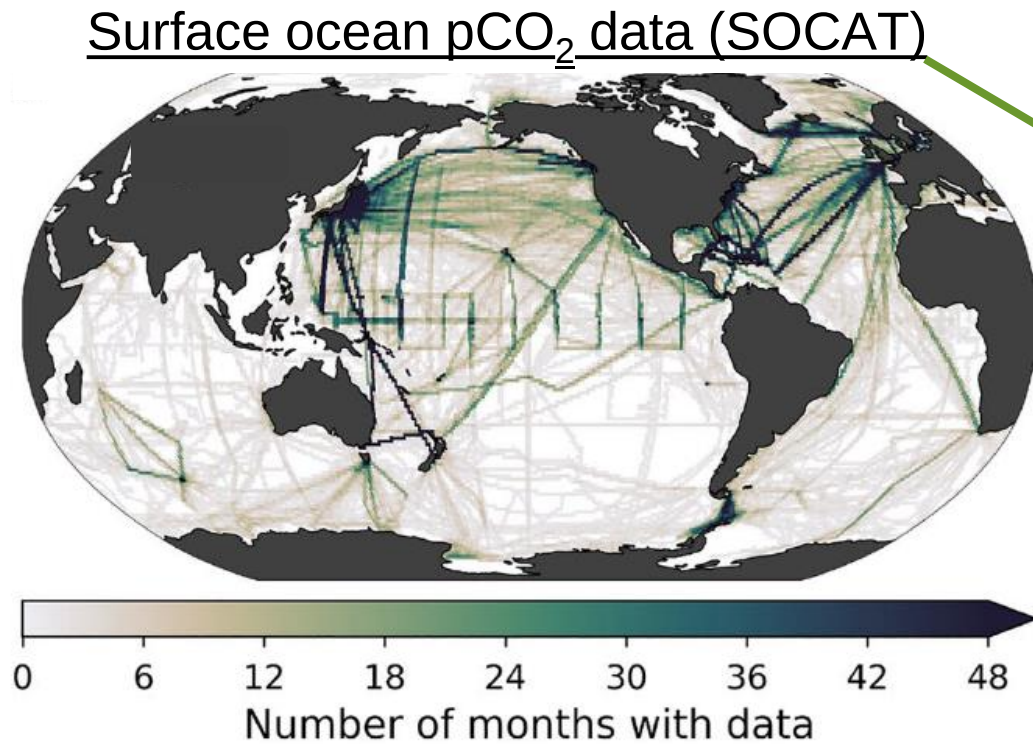
Gloege et al. 2022 JAMES (N=9), Bennington et al. 2022 GRL (N=8)



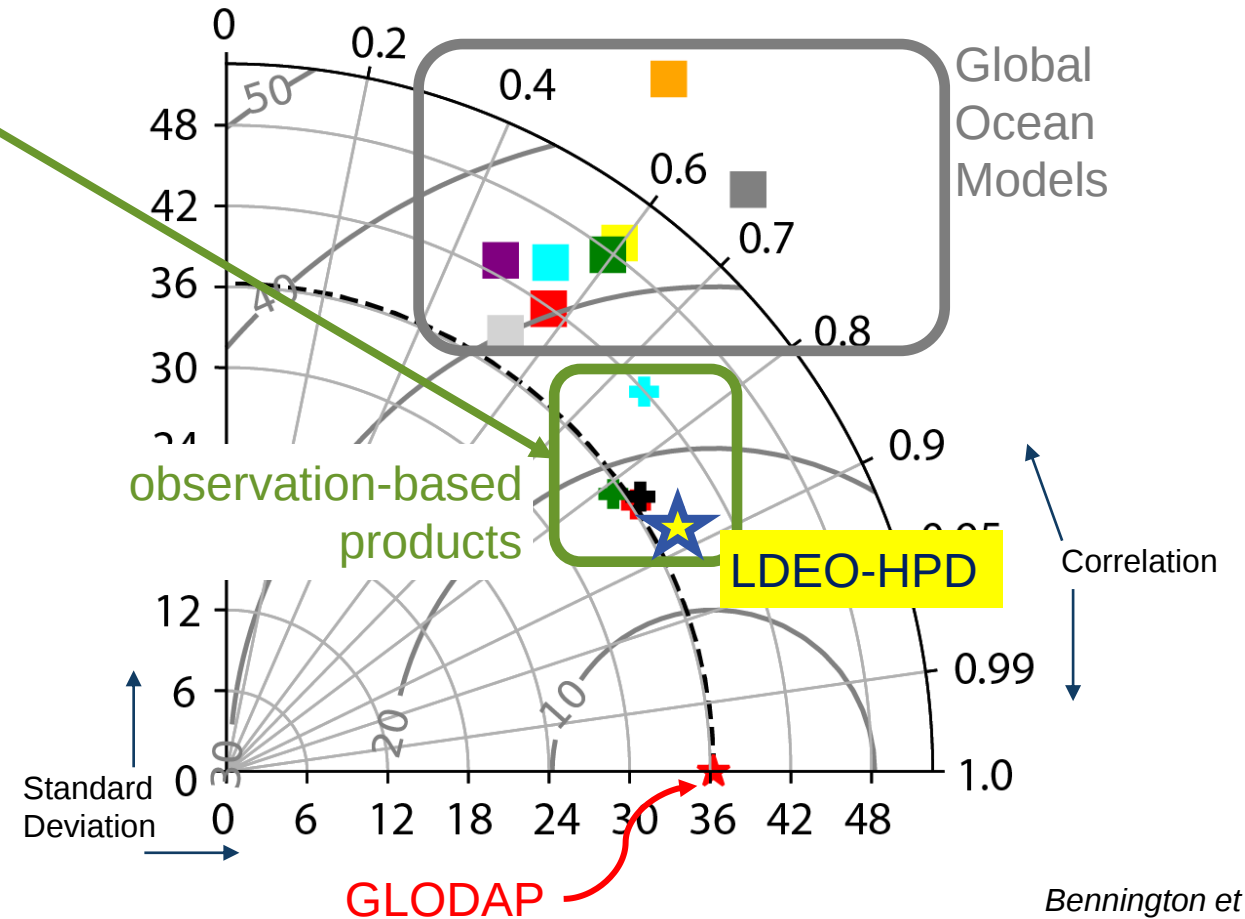
Global air-sea CO₂ flux



LDEO-HPD hybrid model-data approach offers improved skill



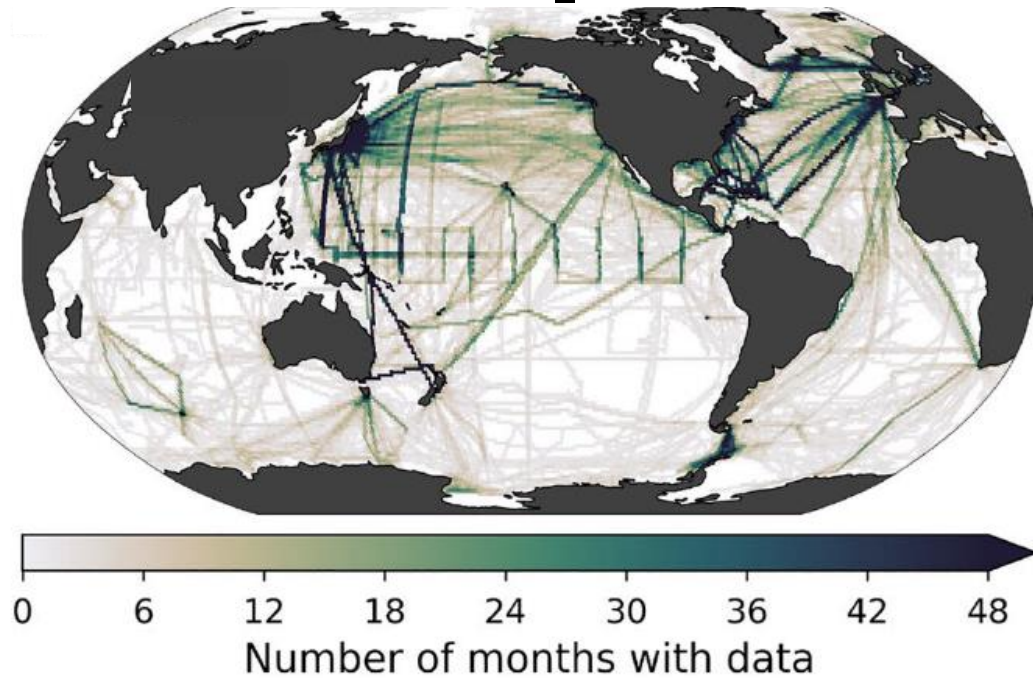
Model + product pCO₂ validation with GLODAP



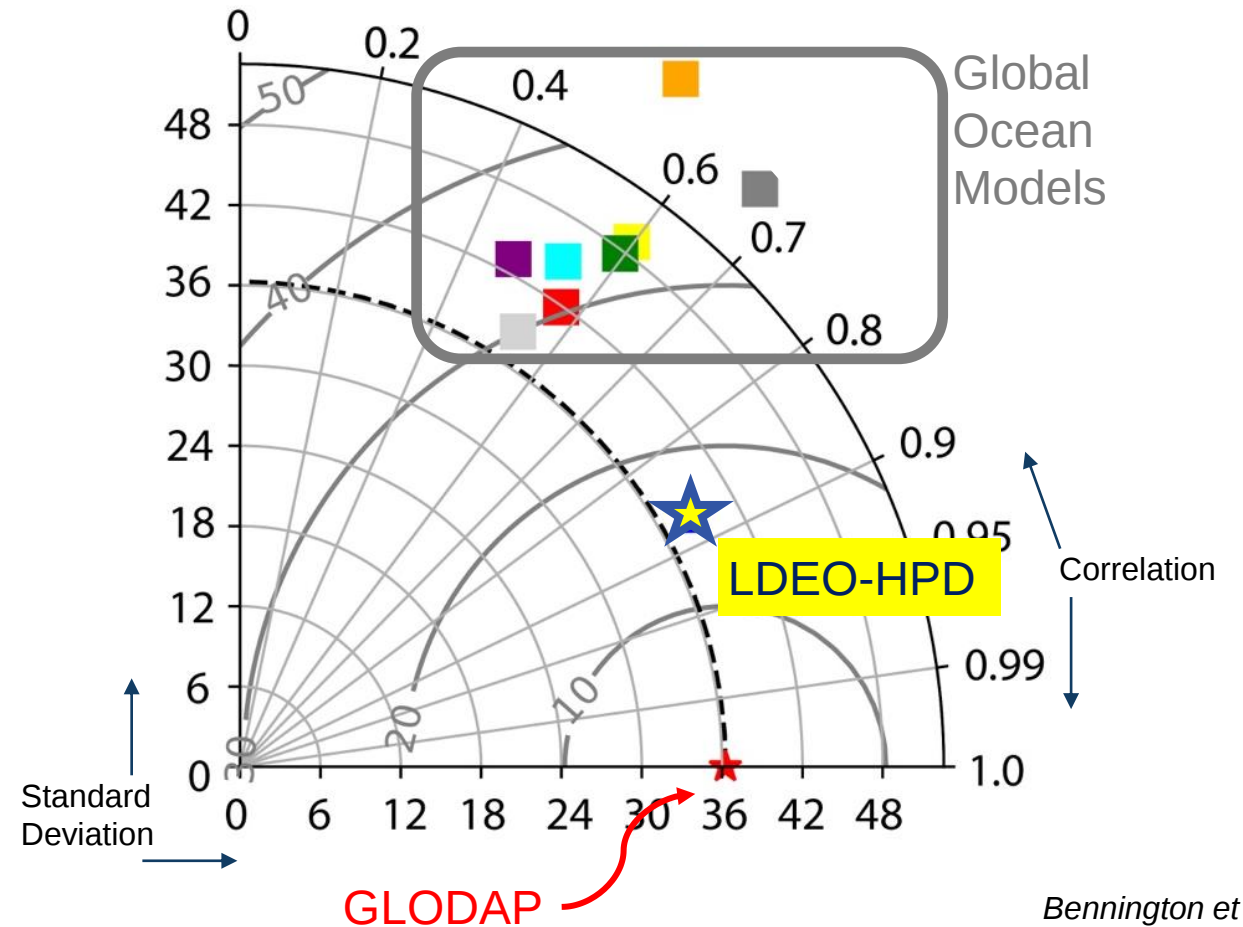
Bennington et al., 2022
Gloege et al. 2021, Fay and McKinley 2021

From where does this enhanced skill arise?

Surface ocean pCO₂ data (SOCAT)

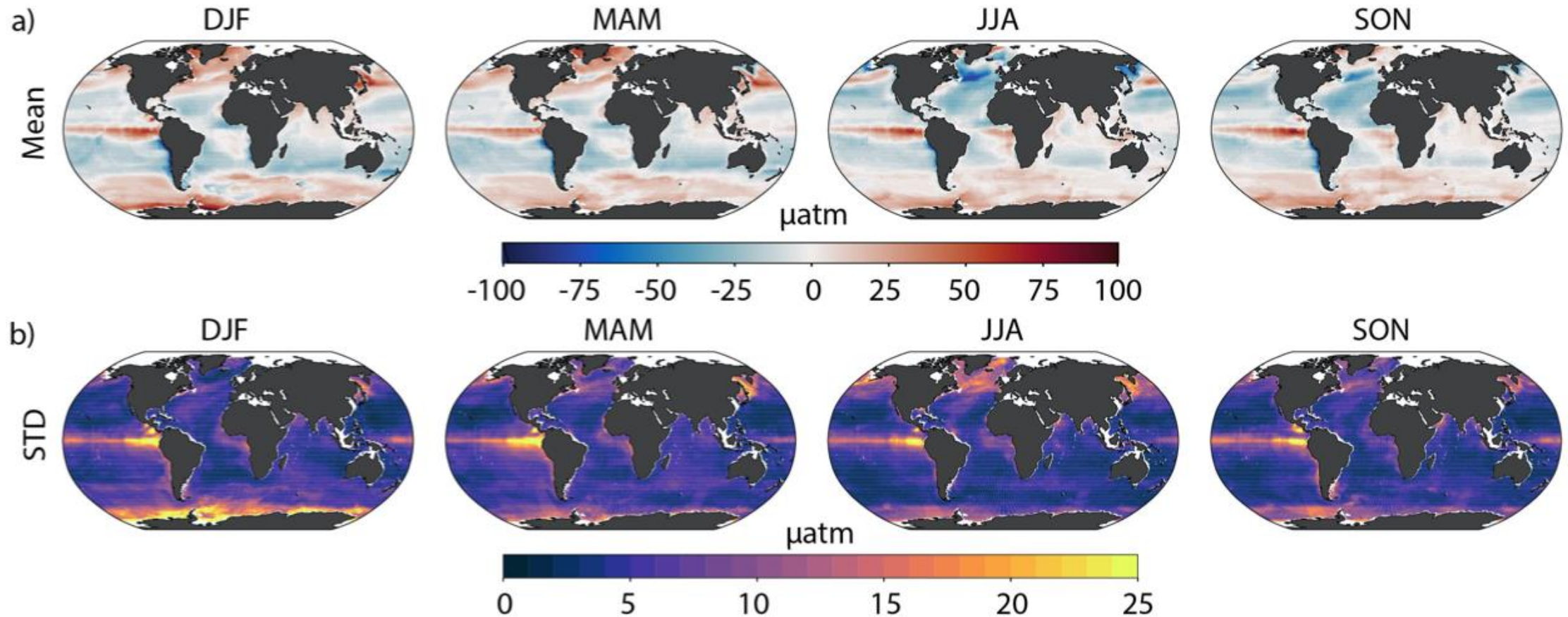


Model + product pCO₂ validation with GLODAP



Bennington et al., 2022
Gloeger et al. 2021, Fay and McKinley 2021

Climatological model-data misfits much larger than interannual

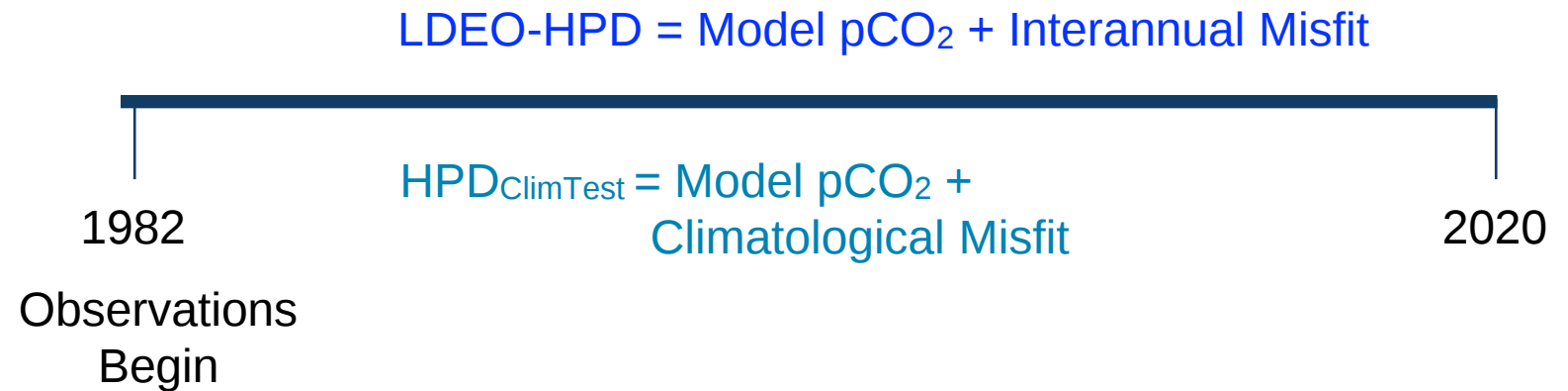


Princeton Model, others similar

Bennington et al. 2022 GRL

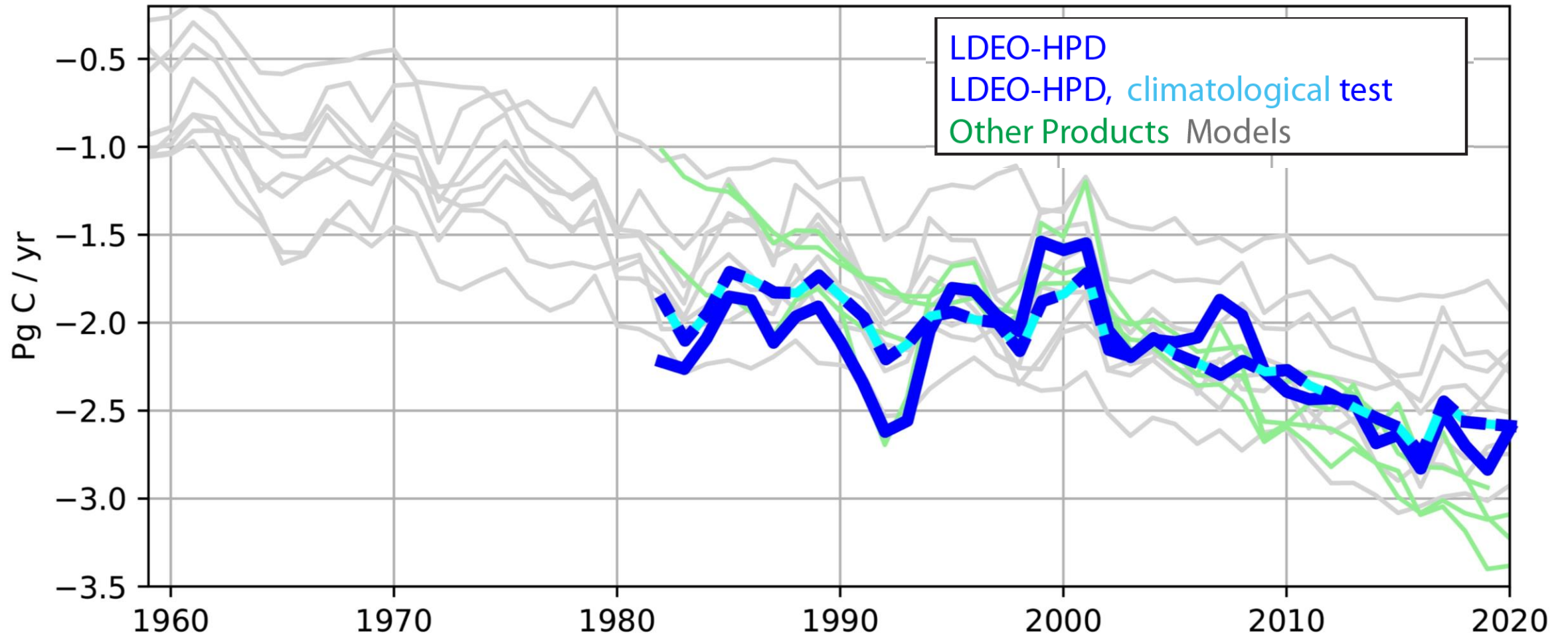
Since climatological misfit dominates, how much skill is gained by applying only this as correction, as opposed to interannual correction?

- $\text{HPD}_{\text{ClimTest}}$ applies the 2000-2020 climatology of the model-observation misfit



Bennington et al. 2022 GRL

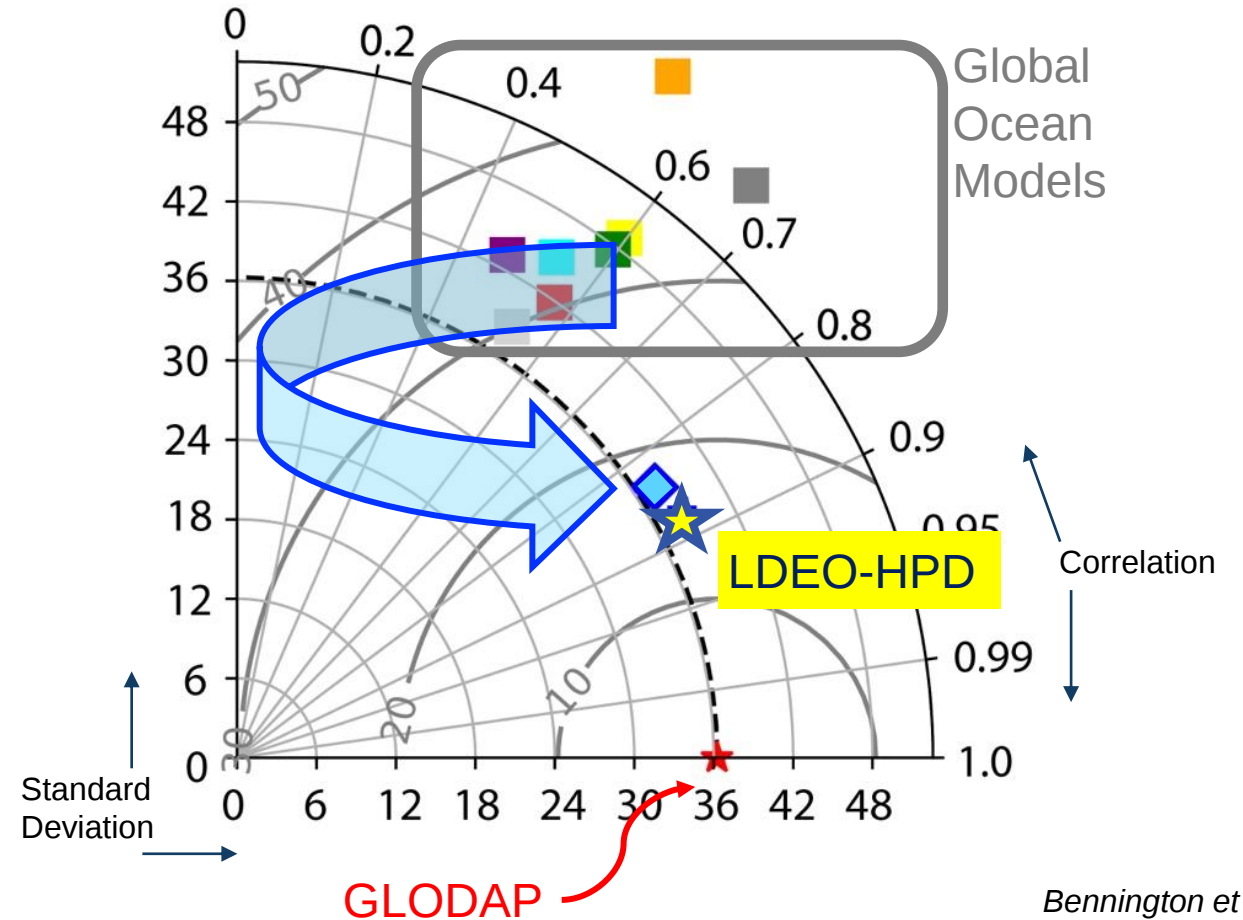
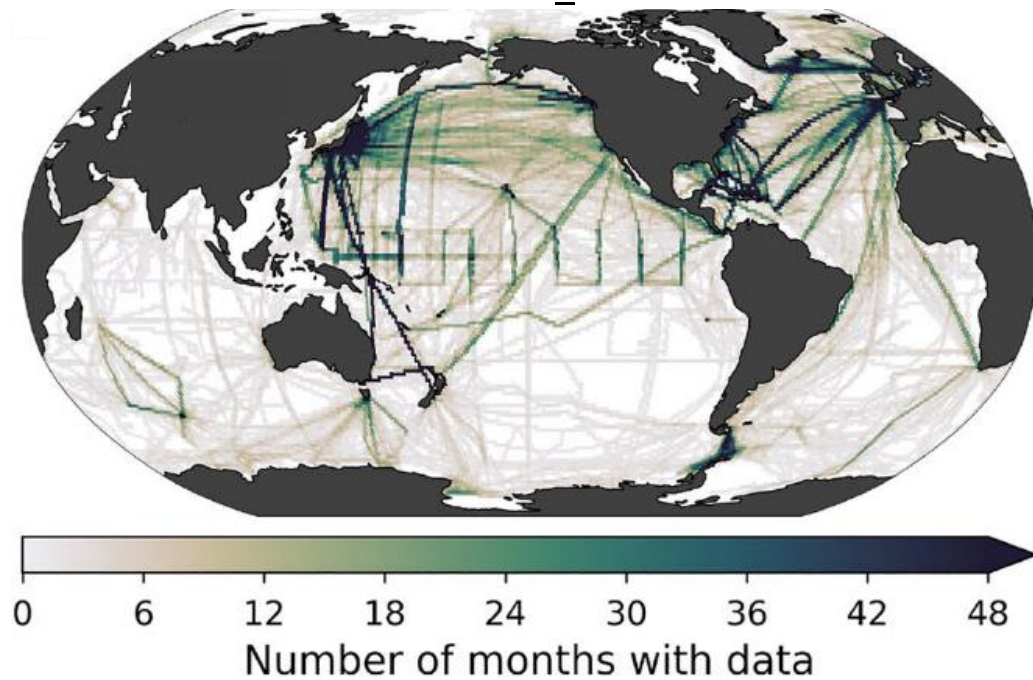
Global air-sea CO₂ flux



Most of the improvement comes from the climatological correction

Model + product pCO₂ validation with GLODAP

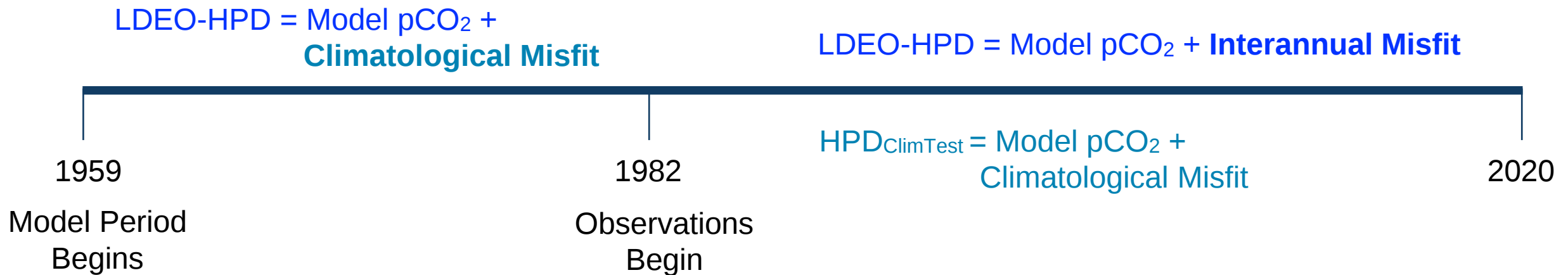
Surface ocean pCO₂ data (SOCAT)



Bennington et al., 2022
Gloeger et al. 2021, Fay and McKinley 2021

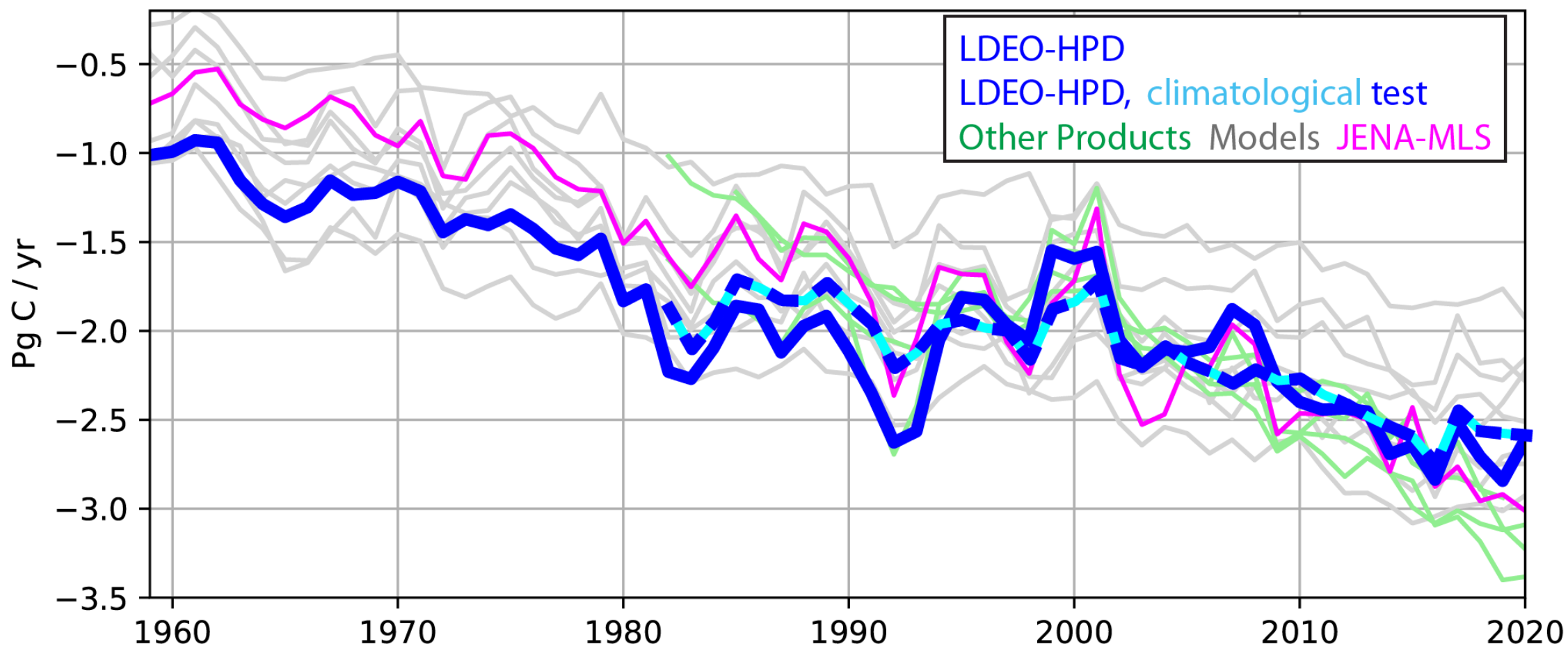
Taking advantage of the climatological misfit being the biggest source of model error, we extend back to 1959

- HPD_{ClimTest} applies the 2000-2020 climatology of the model-observation misfit
- *Since the climatological correction provides most of the additional skill, we use it to correct models in the pre-observed period*



Bennington et al. 2022 GRL

Global air-sea CO₂ flux

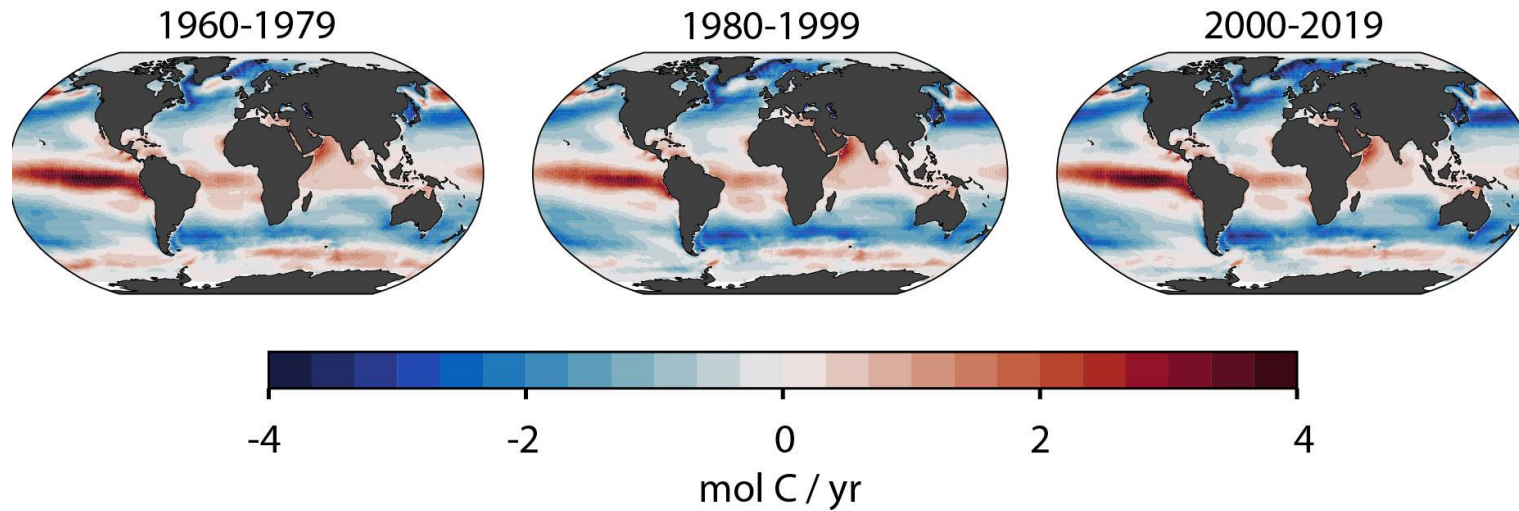


Climatological correction
1959-1982

Interannually varying correction
1982-2020

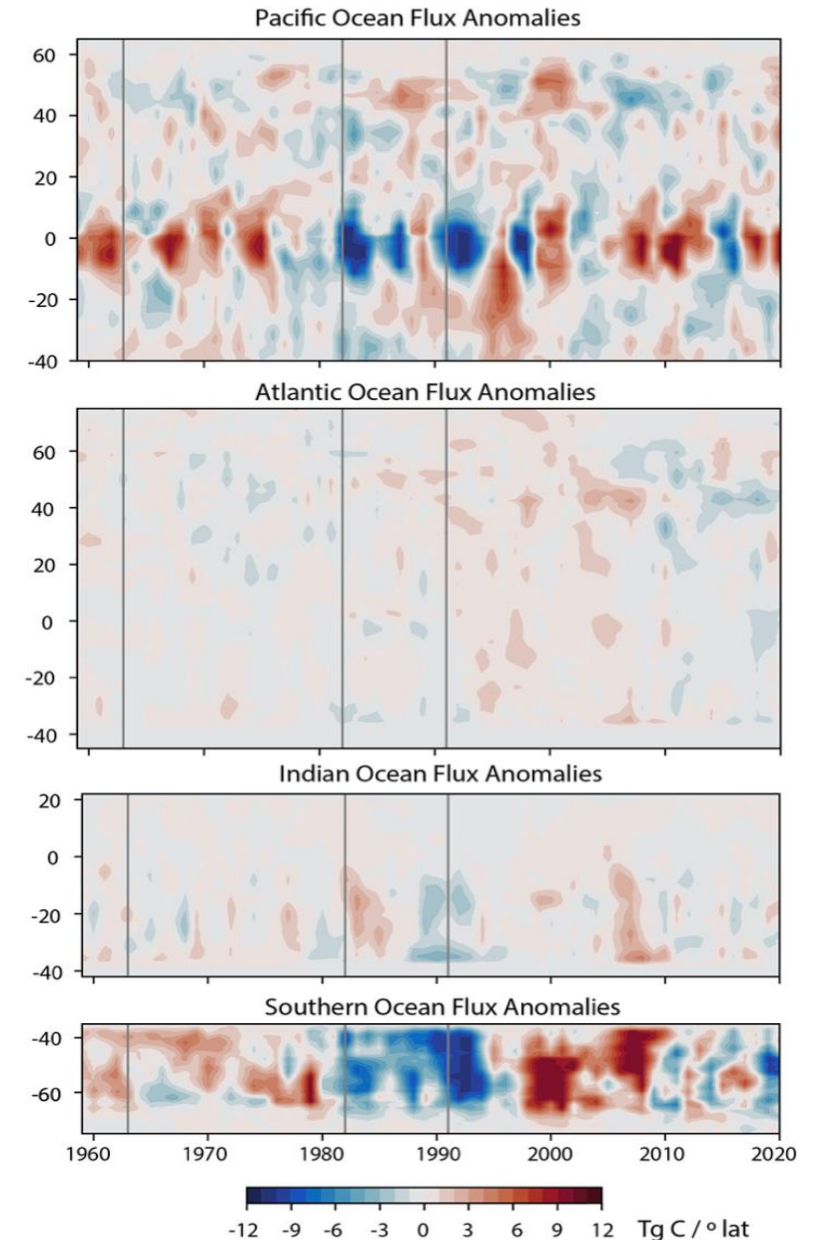
Bennington et al. 2022 GRL

LDEO-HPD: 60+ years of air-sea CO₂ fluxes



- 60+ years of monthly 1x1 air-sea CO₂ fluxes
- Significant decadal variations; coherent between equatorial Pacific and Southern Ocean

Detrended Basin Flux Anomalies



Bennington et al. 2022 GRL

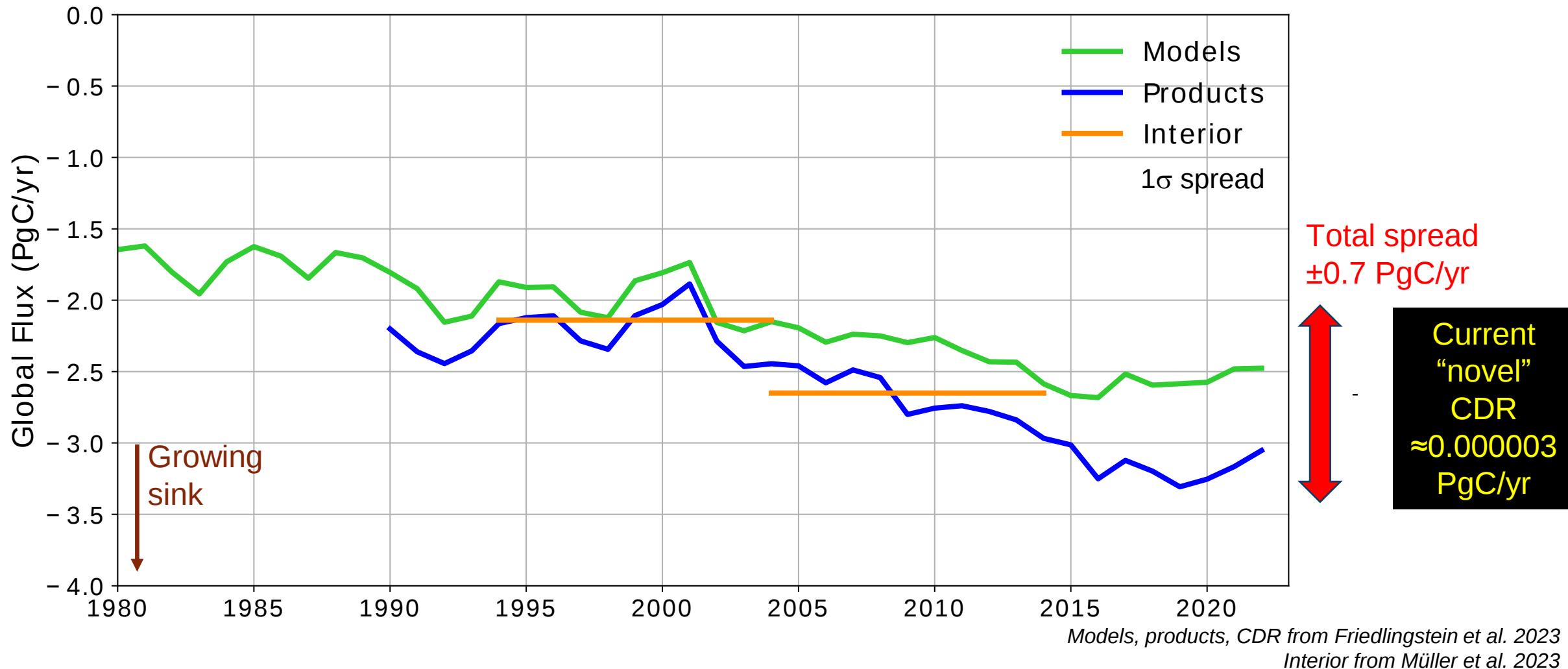
Take home messages

- Air-sea CO₂ flux mean and seasonality can be reconstructed from sparse pCO₂ data, but longer-timescale variability remains uncertain
- Hindcast and Earth System Models have significant mean-state biases
- Correcting models with data offers improved skill and 60+ years of reconstructed monthly surface ocean pCO₂

Outline

- Key processes of the ocean carbon cycle
- Improving quantification of the ocean carbon sink
 - *Models*: evaluate ML-based data product skill, given sampling
 - *pCO₂ data products*: apply to identify model mean-state biases
 - *Models and data*: combine using ML
 - ***marine Carbon Dioxide Removal (mCDR)*** – quantifying “additionality”?

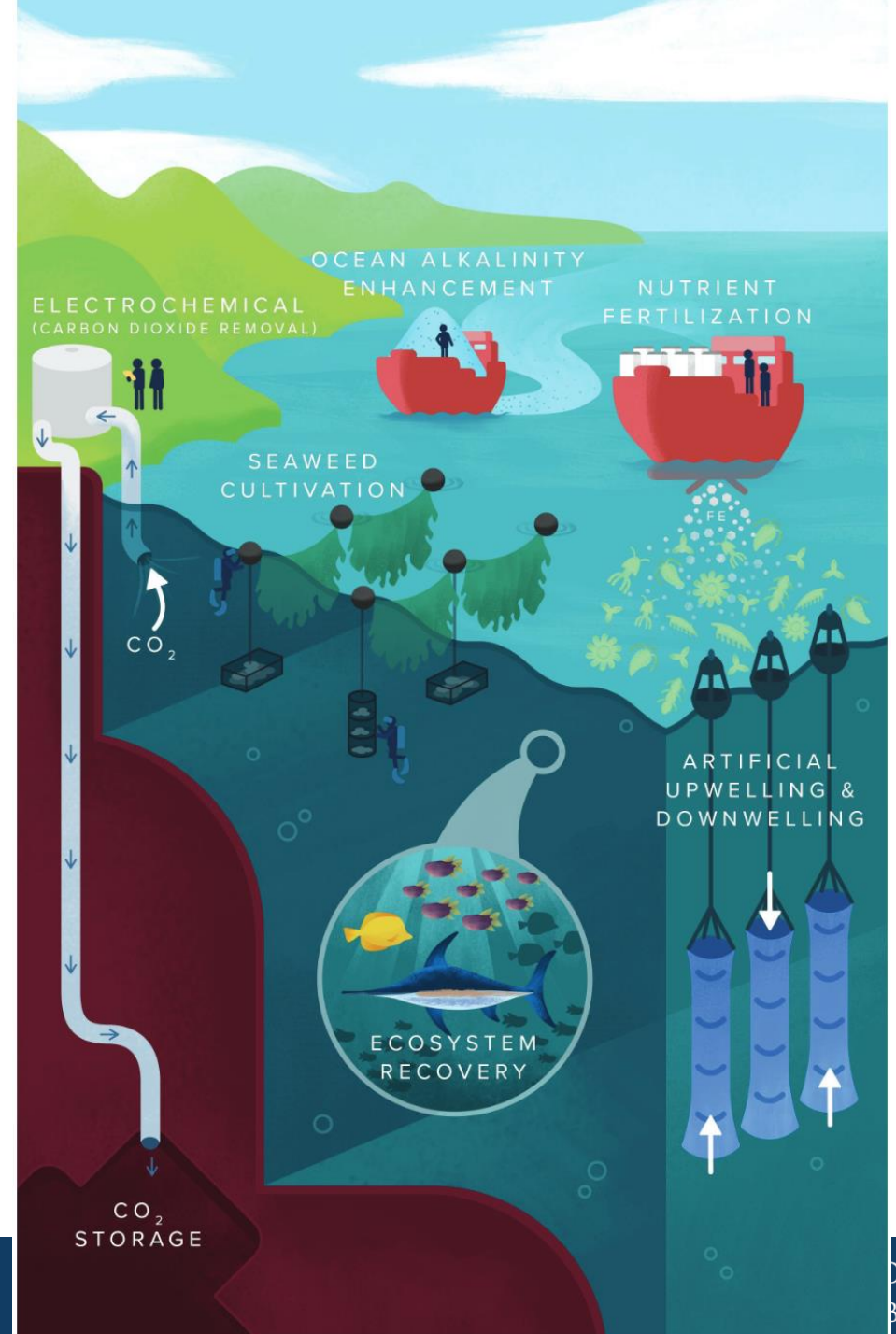
“novel” CDR is imperceptible with respect to the ocean sink



Marine Carbon Dioxide Removal *mCDR*

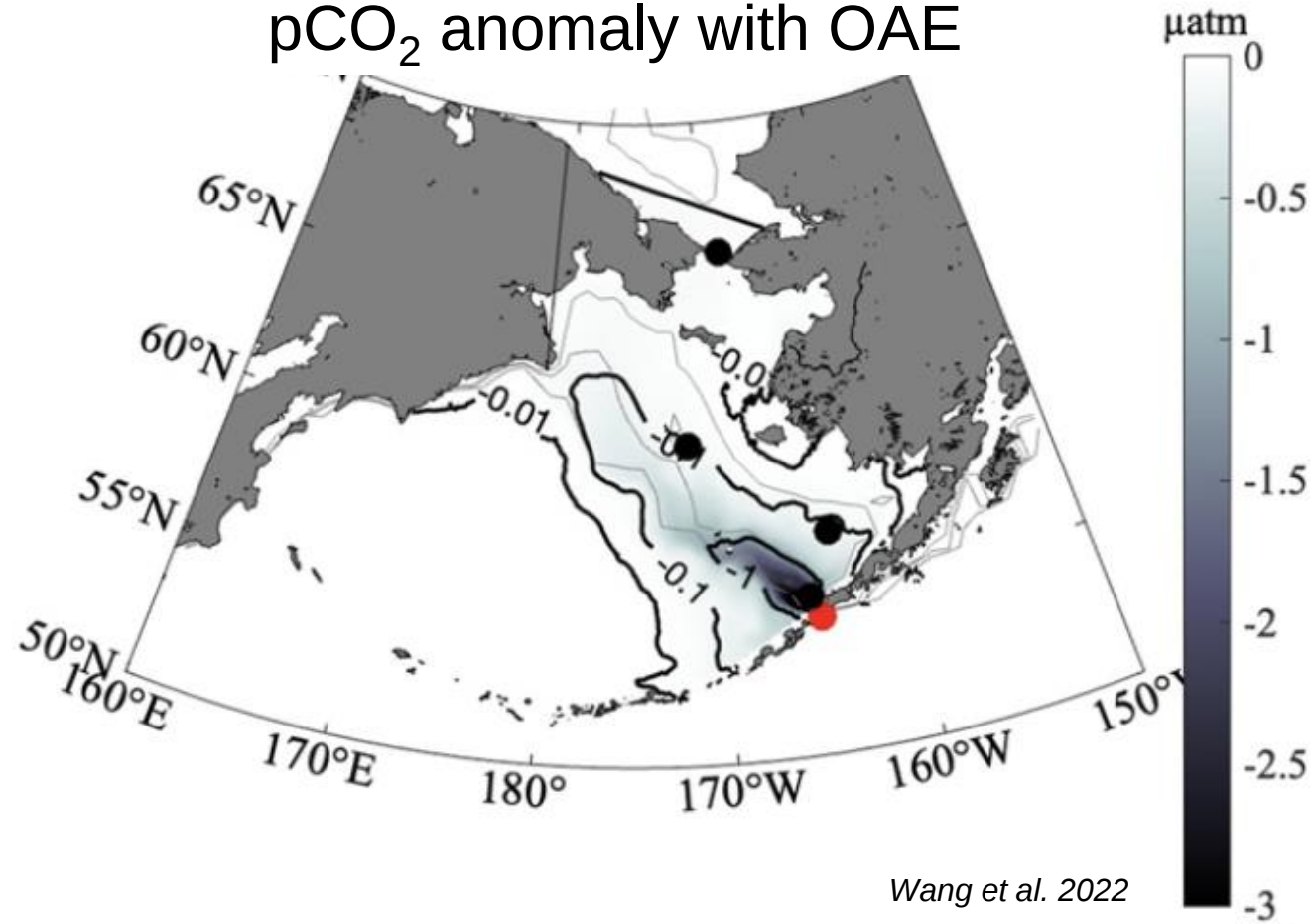
- Efficacy?
- Safety?
- Durability?
- **Additionality?**

NASEM 2021



mCDR pCO₂ impacts

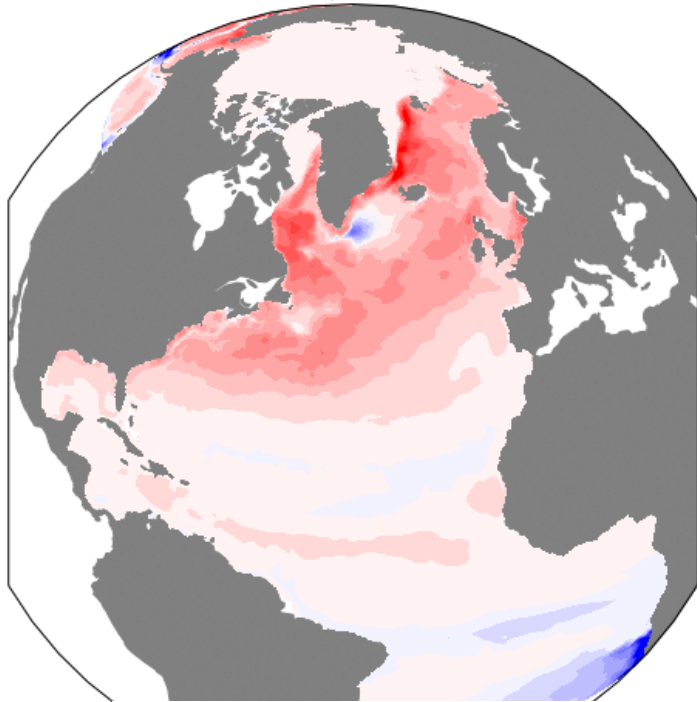
10-year mean modeled
pCO₂ anomaly with OAE



Wang et al. 2022

Global approaches potential for time-space resolved baseline

Models

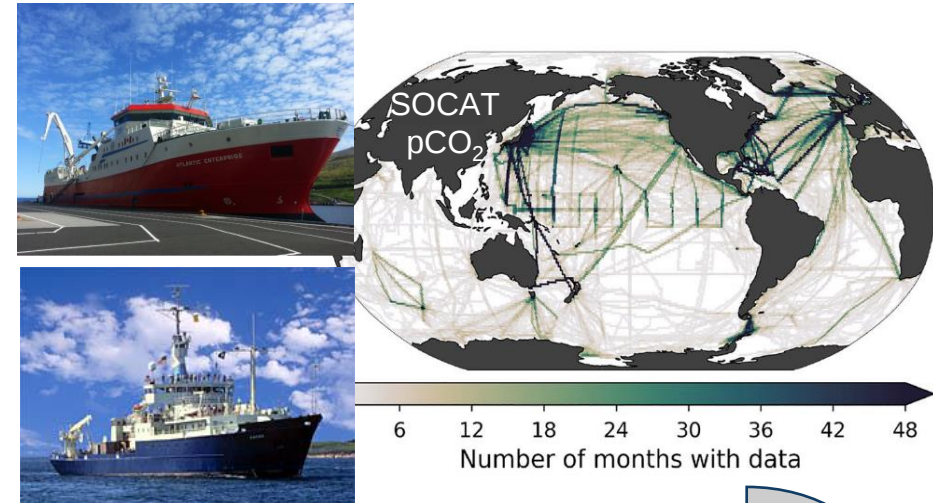


ASTE-BGC
air-sea CO₂ flux

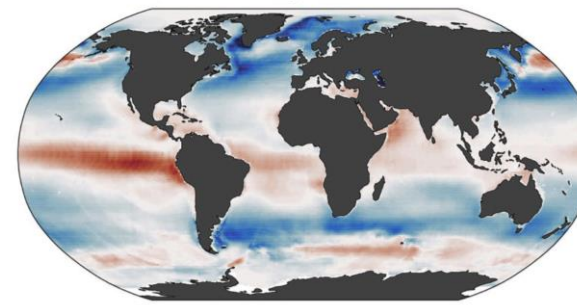
Surface pCO₂ data products

Air-sea fluxes (~monthly)

Model validation



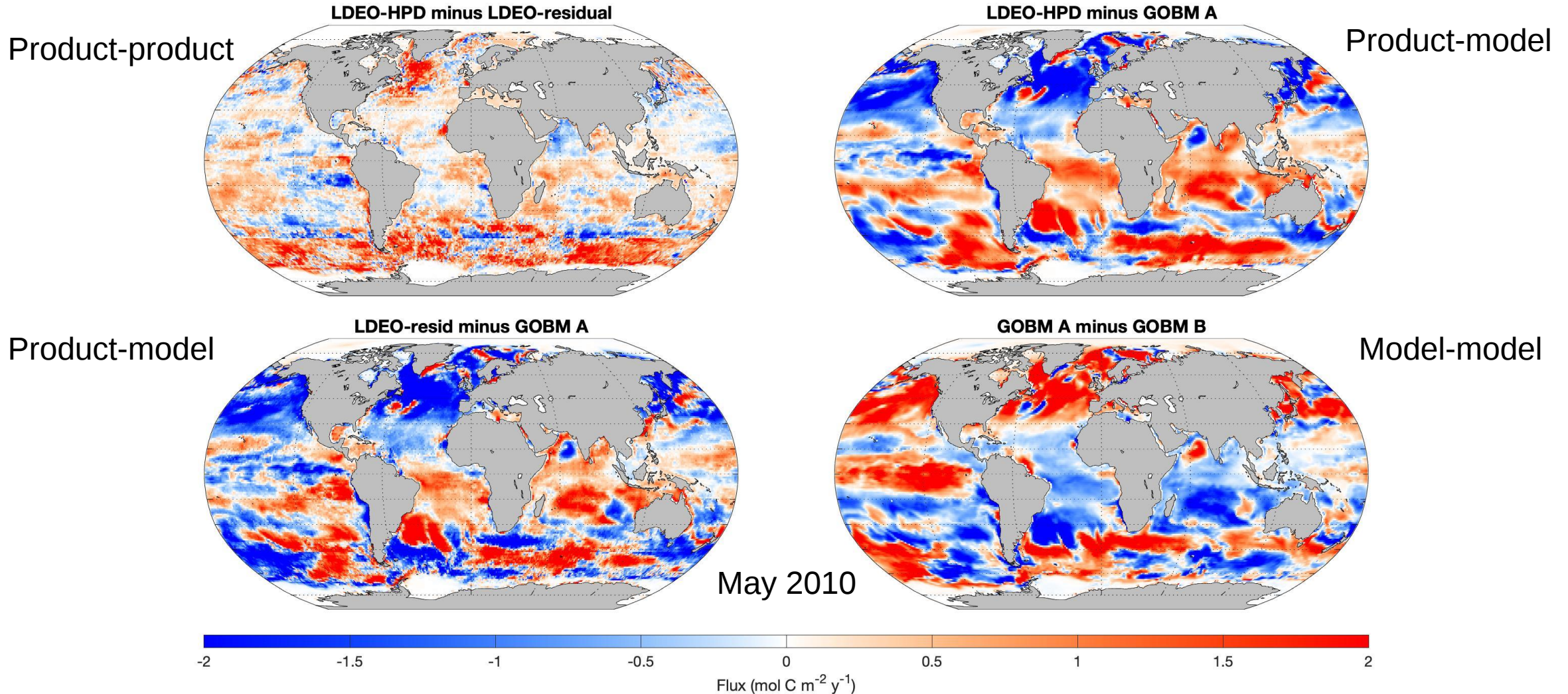
a) CO₂ Flux



AI/ML

Air-sea flux uncertainties at smaller space-time scales are huge

Uncertainties are estimated here as difference between individual estimates

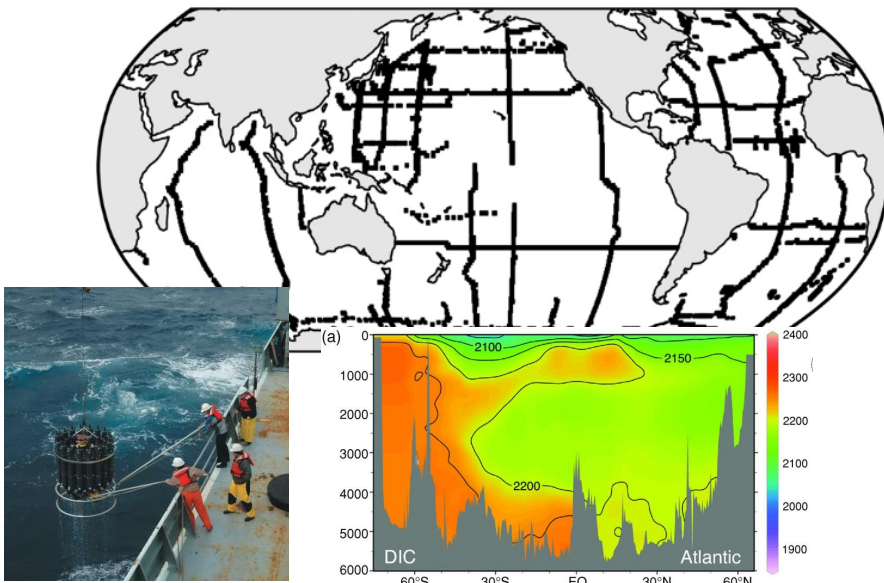


Local differences are same magnitude as long-term mean flux

Independent $p\text{CO}_2$ to assess product and model skill, 2000-2023

Hydrography

GLODAP

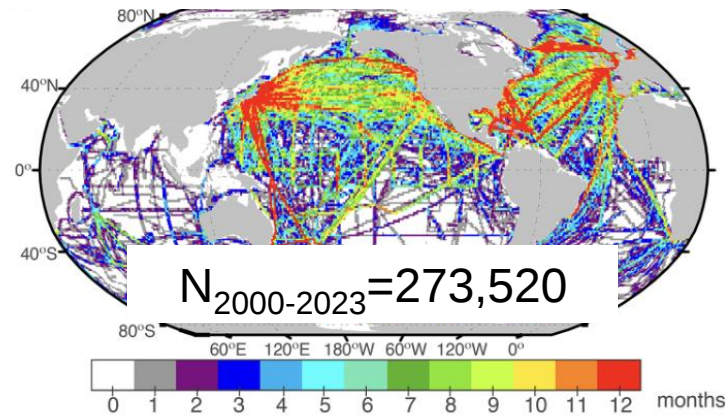


$N_{2000-2023}=1,043$

McKinley et al. in prep

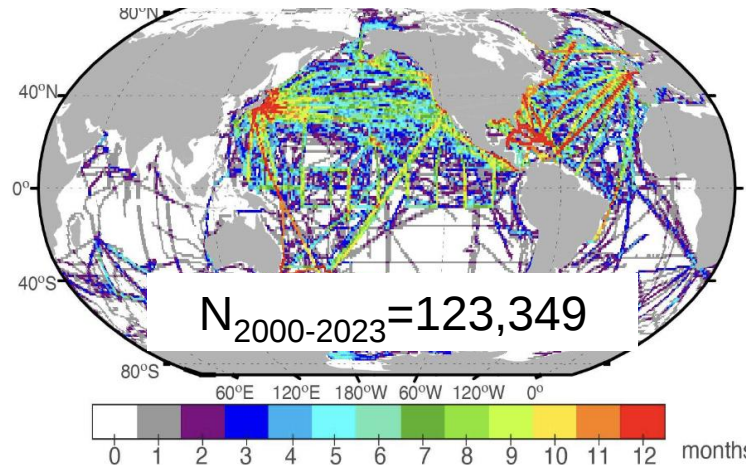
Surface $p\text{CO}_2$

SOCAT $p\text{CO}_2$, 1980-present



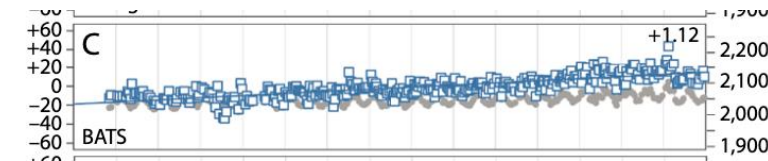
$N_{2000-2023}=273,520$

LDEO $p\text{CO}_2$, 1980-2019



$N_{2000-2023}=123,349$

Timeseries

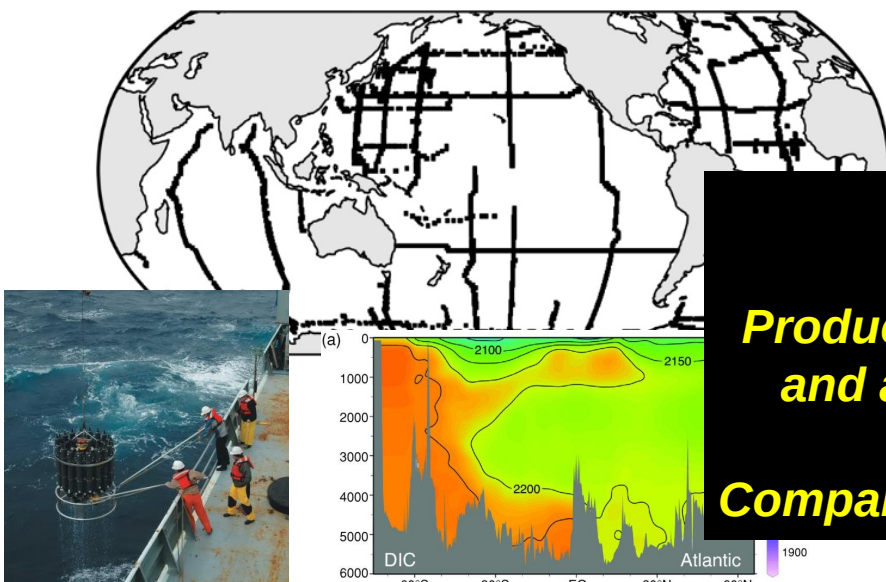


$N^{\text{BATS}}_{2000-2023}=225$

Independent $p\text{CO}_2$ to assess product and model skill

Hydrography

GLODAP

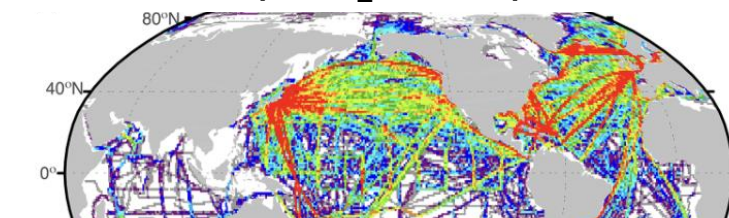


$N_{2000-2023}=1,043$

McKinley et al. in prep

Surface $p\text{CO}_2$

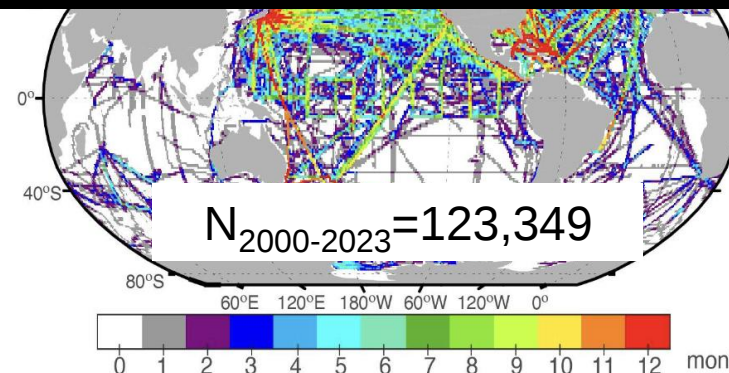
SOCAT $p\text{CO}_2$, 1980-present



Models can be compared to all

Products are trained with SOCAT, so remove this and any SOCAT points in LDEO from analysis

Comparison to 10 models, 10 products of GCB2024



$N_{2000-2023}=123,349$

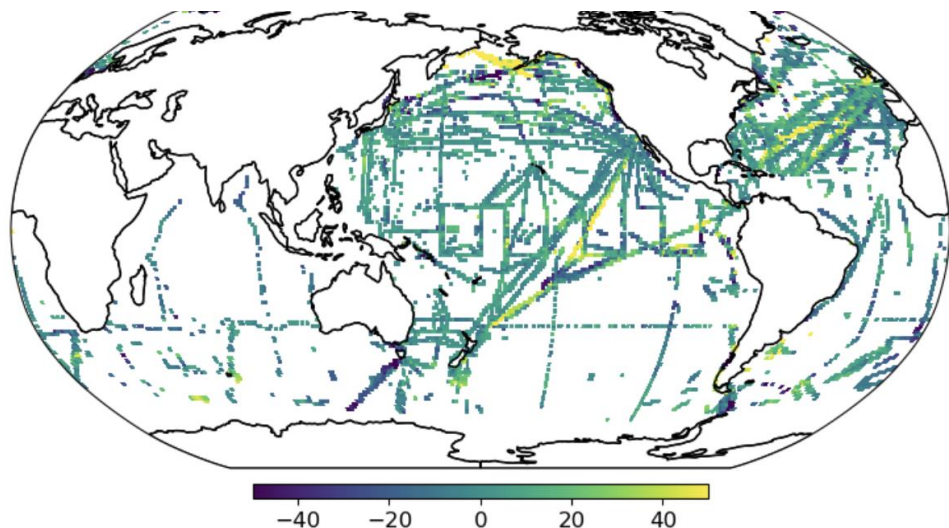
Timeseries



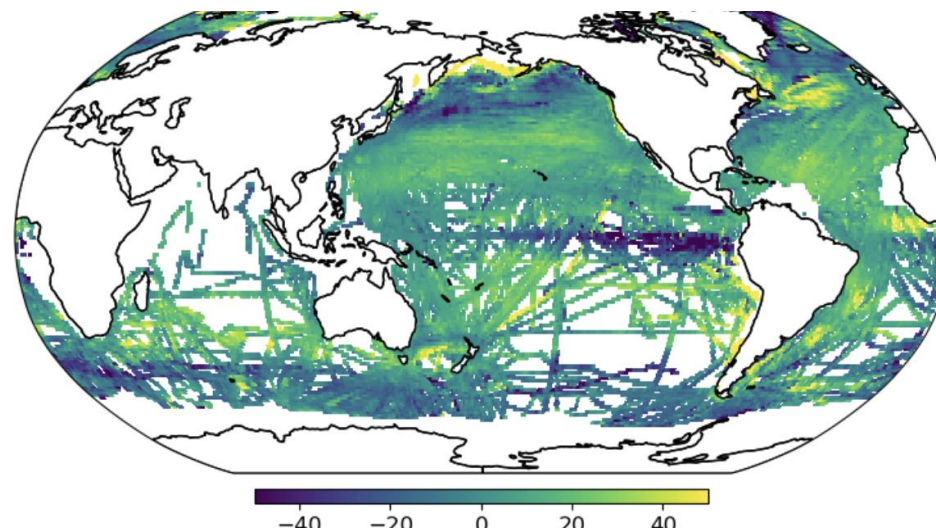
$N^{\text{BATS}}_{2000-2023}=225$

Global $p\text{CO}_2$: Models vs. products, and vs. independent data

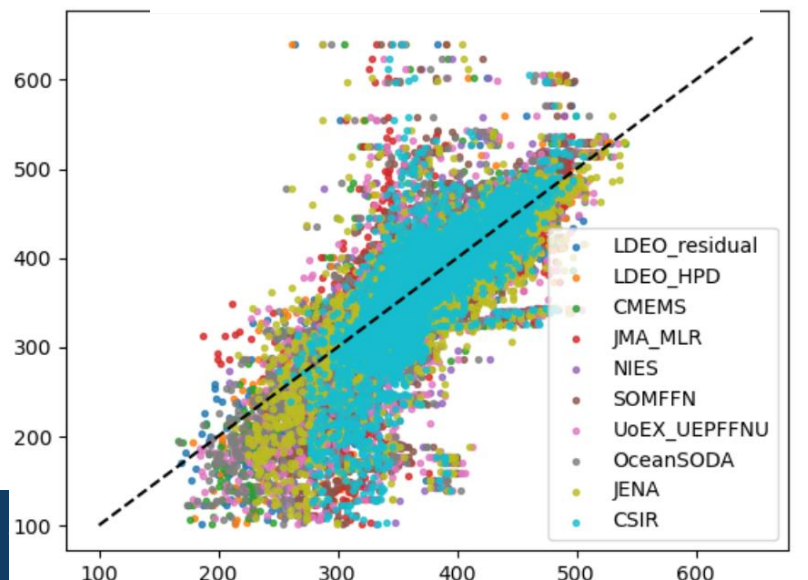
One Product - Data



One Model - Data



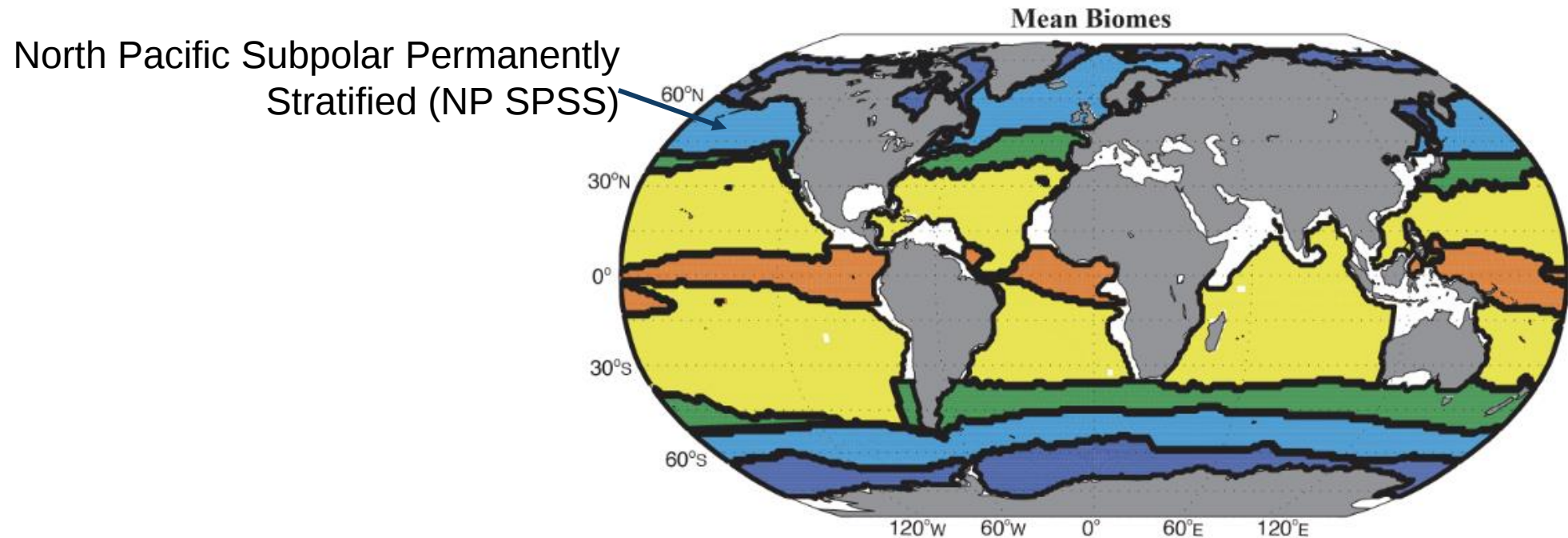
LDEO+GLODAP vs. Products



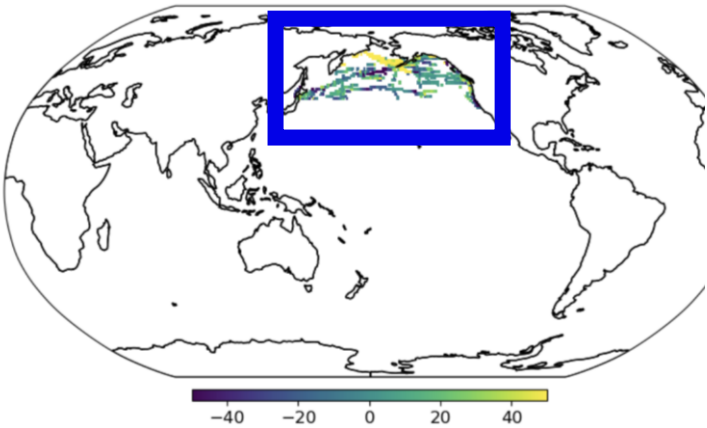
RMSE and correlation

	10 Product mean	10 Model mean
2000-2023 LDEO + GLODAP (n=12,590)	28.0 μatm $r=0.80$	
SOCAT + GLODAP (n=285,256)		32.9 μatm $r=0.63$

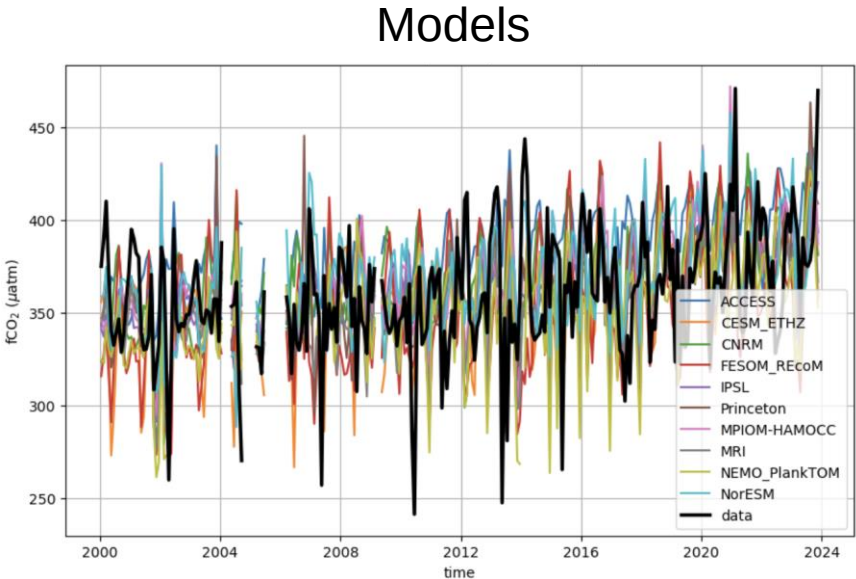
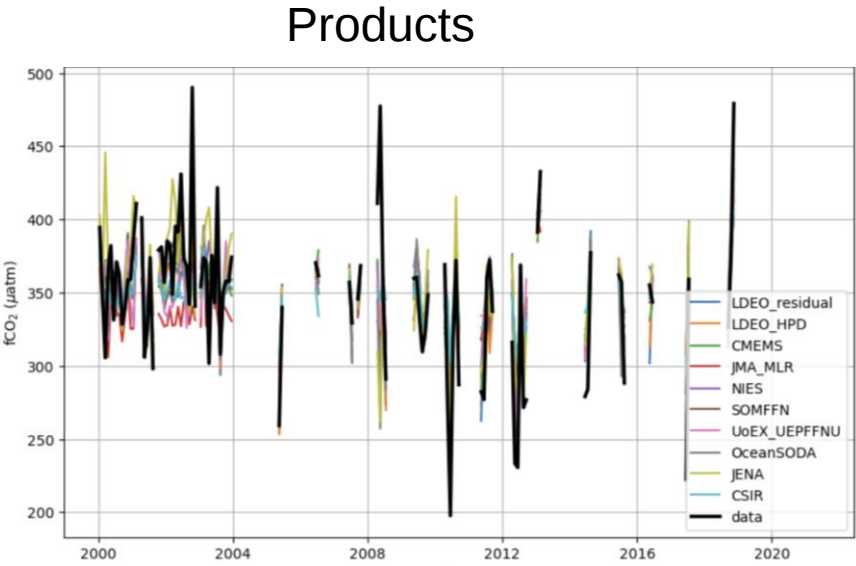
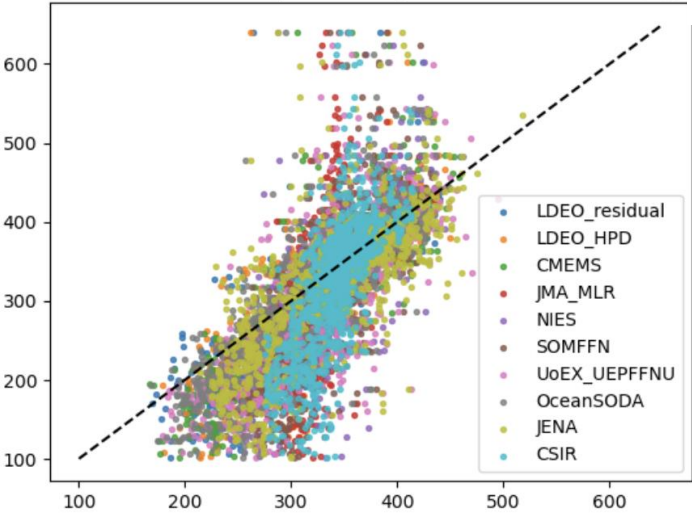
Biomes are mechanistically similar; use these for temporal decomposition



N. Pacific SPSS: Models and products vs. independent data



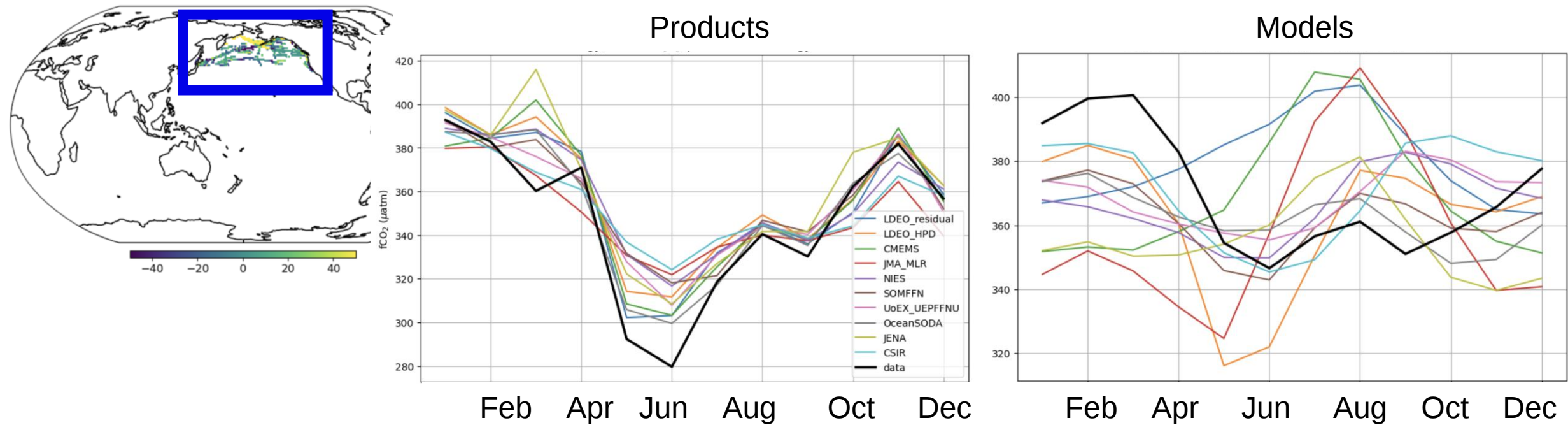
LDEO+GLODAP vs. Products



RMSE and correlation

2000-2023	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,240)	54.3µatm r=0.77	
SOCAT + GLODAP (n=23,109)		54.4µatm r=0.30

N. Pacific SPSS: Climatological

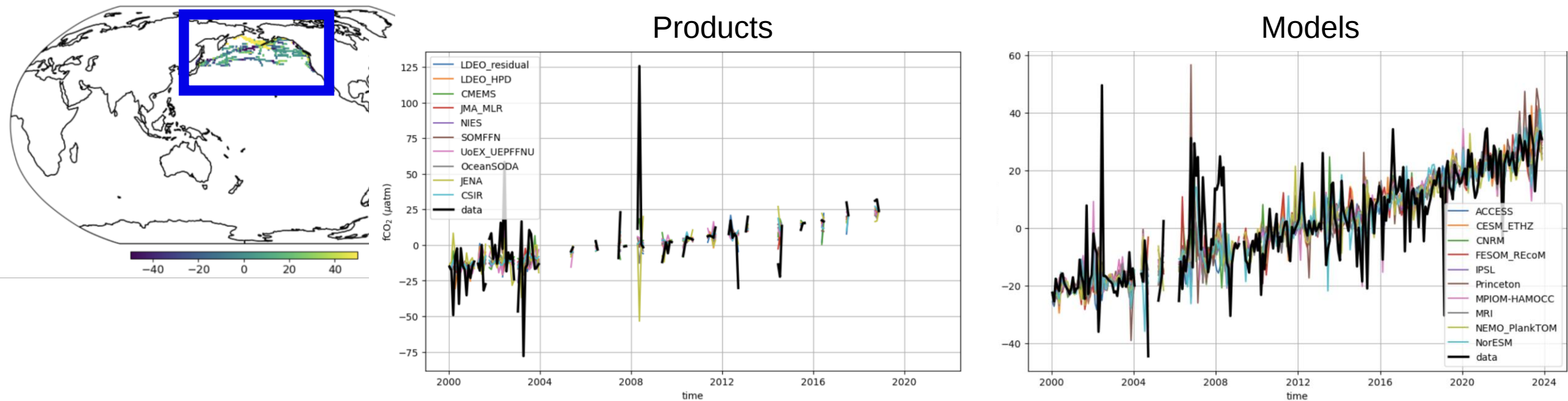


RMSE and correlation

	10 Product mean	10 Model mean
2000-2023 LDEO + GLODAP (n=1,240)	50.7 μatm r=0.81	
SOCAT + GLODAP (n=23,109)		47.4 μatm r=0.28

McKinley et al. in prep

N. Pacific SPSS: Remove climatology



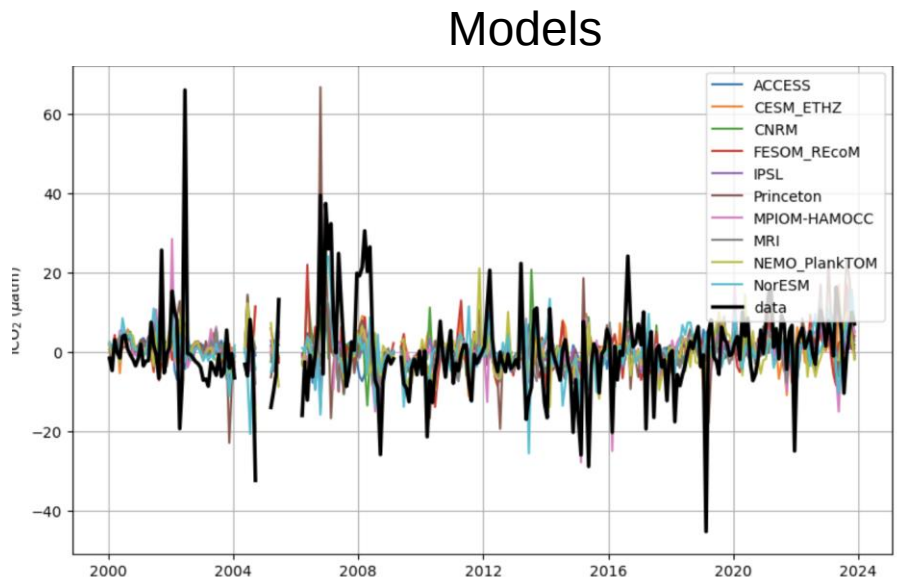
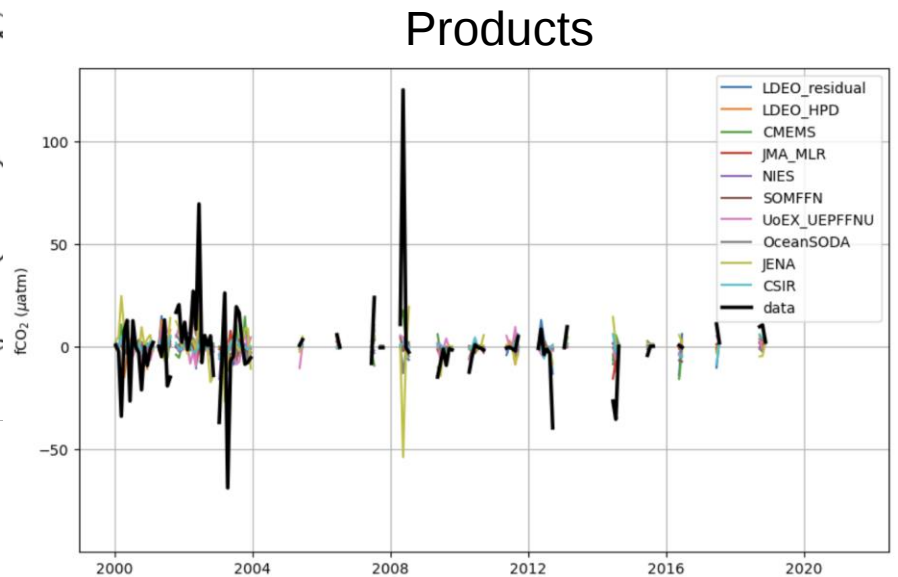
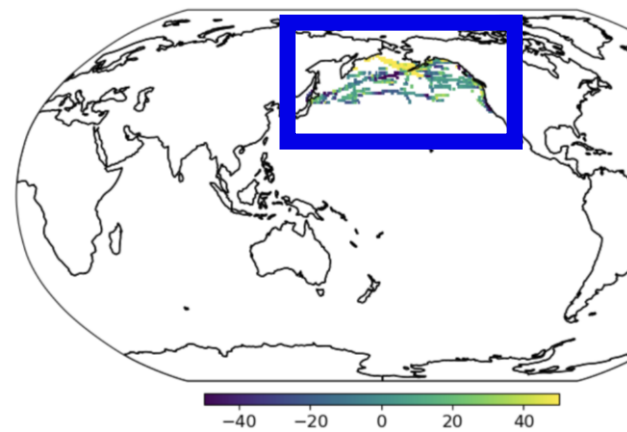
Product, Model, Datasets
Trends = 1.9-2.1 uatm/yr

RMSE and correlation

	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,240)	18.9µatm r=0.50	
SOCAT + GLODAP (n=23,109)		23.2µatm r=0.54

McKinley et al. in prep

N. Pacific SPSS: Remove trend



RMSE and correlation

	10 Product mean	10 Model mean
2000-2023 LDEO + GLODAP (n=1,240)	18.9µatm r=0.23	
SOCAT + GLODAP (n=23,109)		23.2µatm r=0.14

McKinley et al. in prep

Take home messages

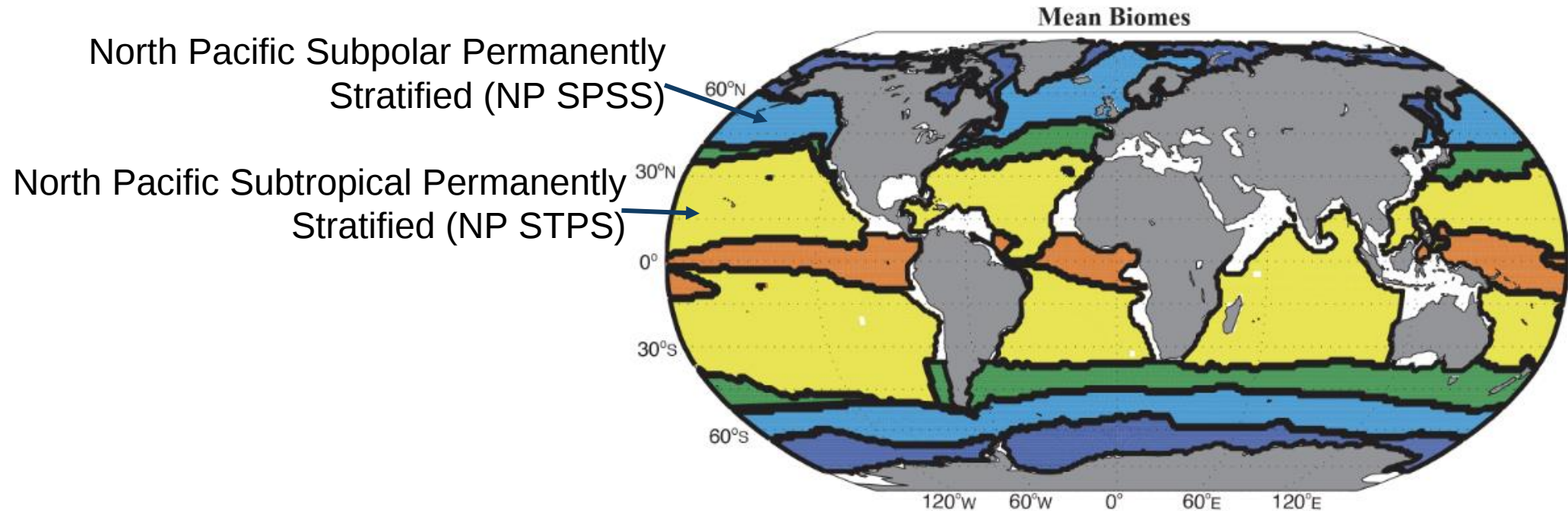
- Air-sea CO₂ flux mean and seasonality can be reconstructed from sparse pCO₂ data, but longer-timescale variability remains uncertain
- Hindcast and Earth System Models have significant mean-state biases
- Correcting models with data offers improved skill and 60+ years of reconstructed monthly surface ocean pCO₂
- Current models and data products have local uncertainties orders of magnitude larger than expected mCDR impacts

Thank you!

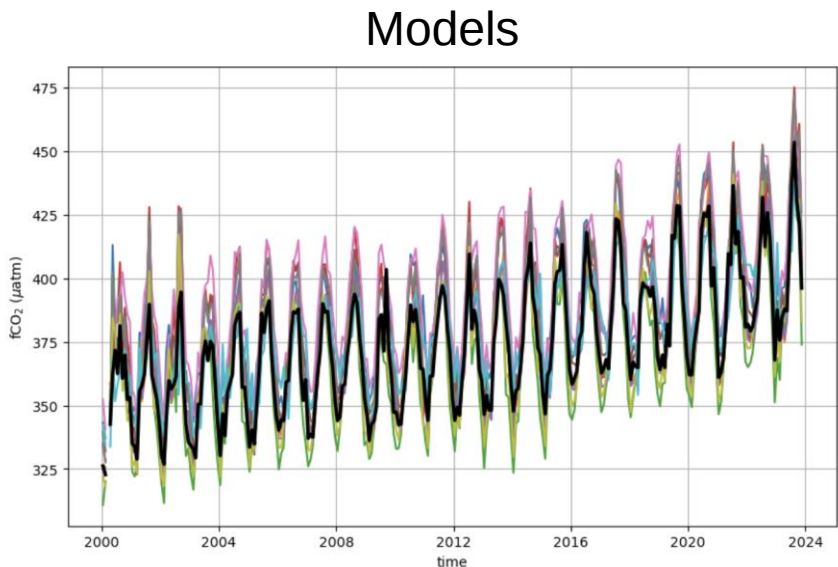
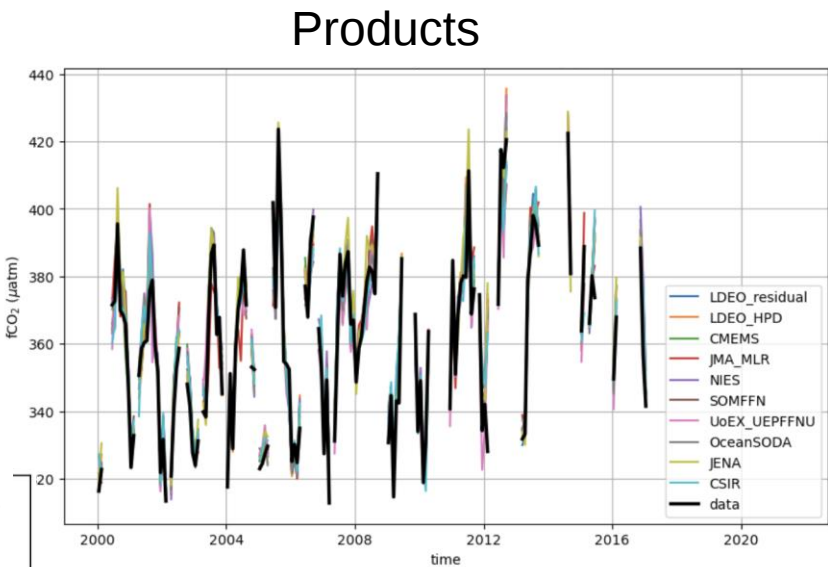
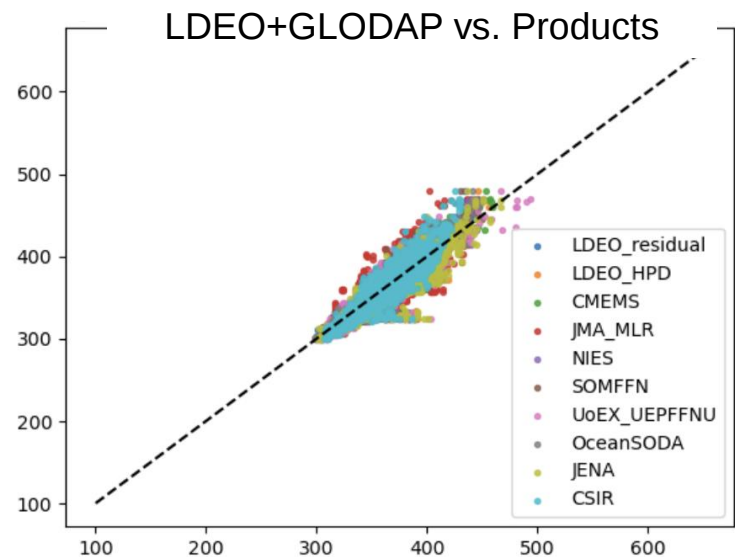
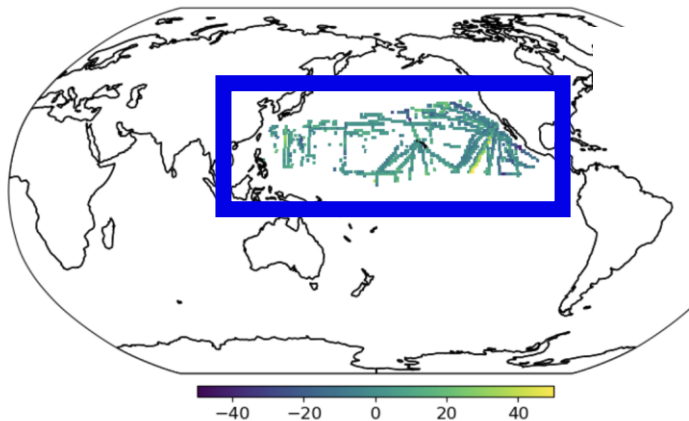


Viviana Acquaviva (professor, CUNY), Ce Bian (postdoc), Galen McKinley
Thea Heimdal (scientist), Abby Shaum (PhD student), Amanda Fay (senior scientist)

Biomes are mechanistically similar; use these for temporal decomposition



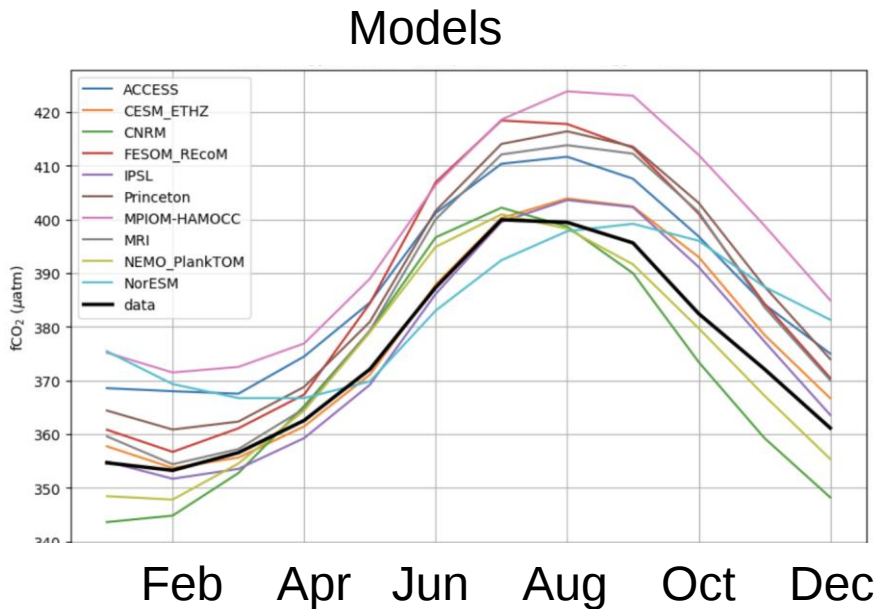
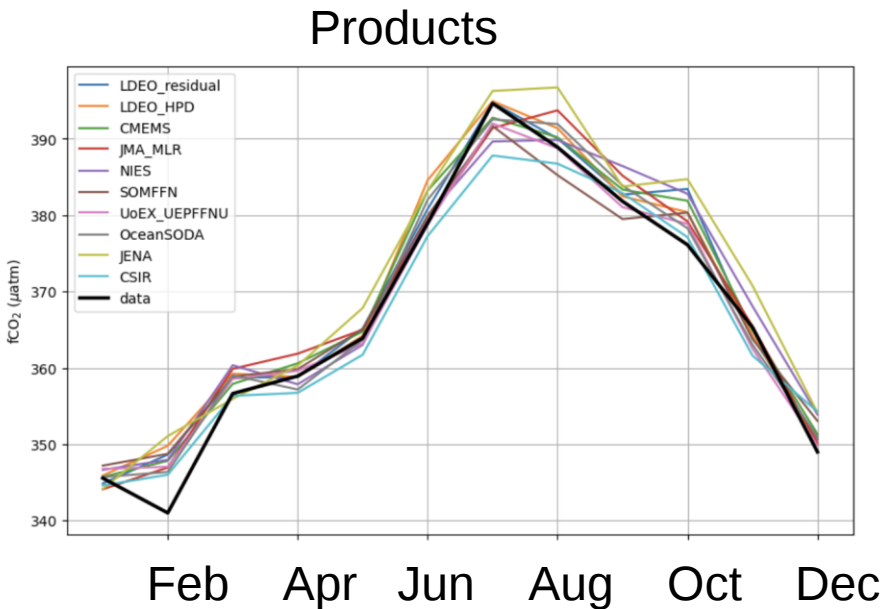
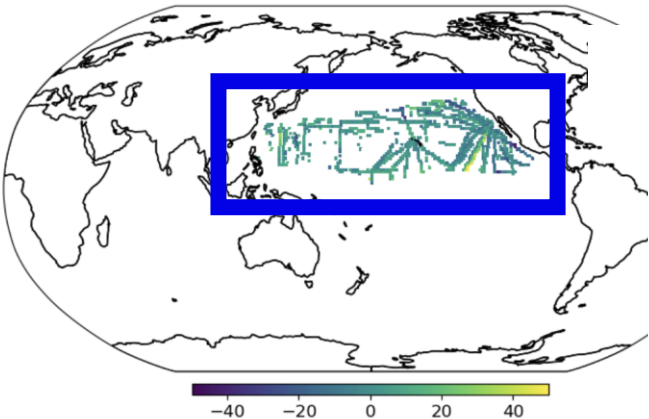
N. Pacific STPS: Models and products vs. independent data



RMSE and correlation

	2000-2023	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,931)		12.3µatm r=0.91	
SOCAT + GLODAP (n=41,820)			18.3µatm r=0.88

N. Pacific STPS: Climatological

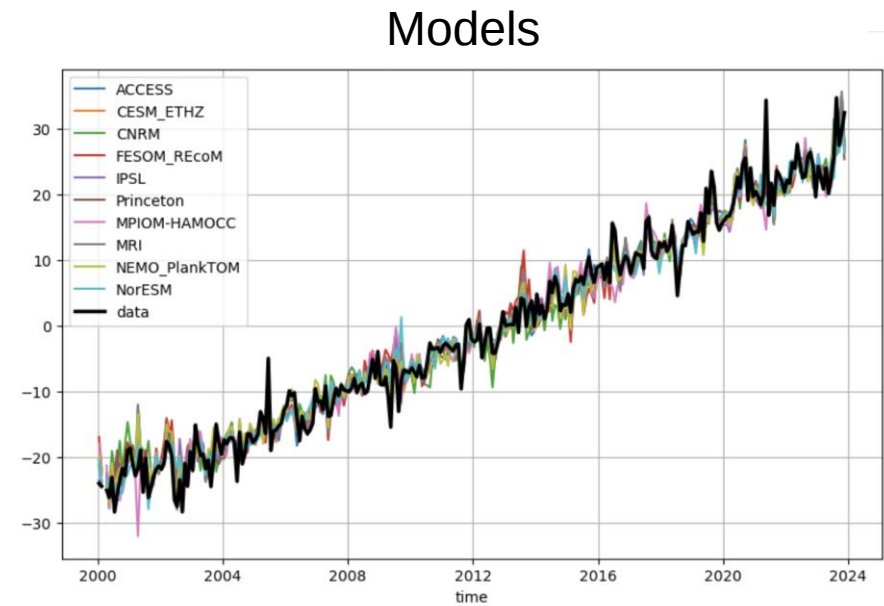
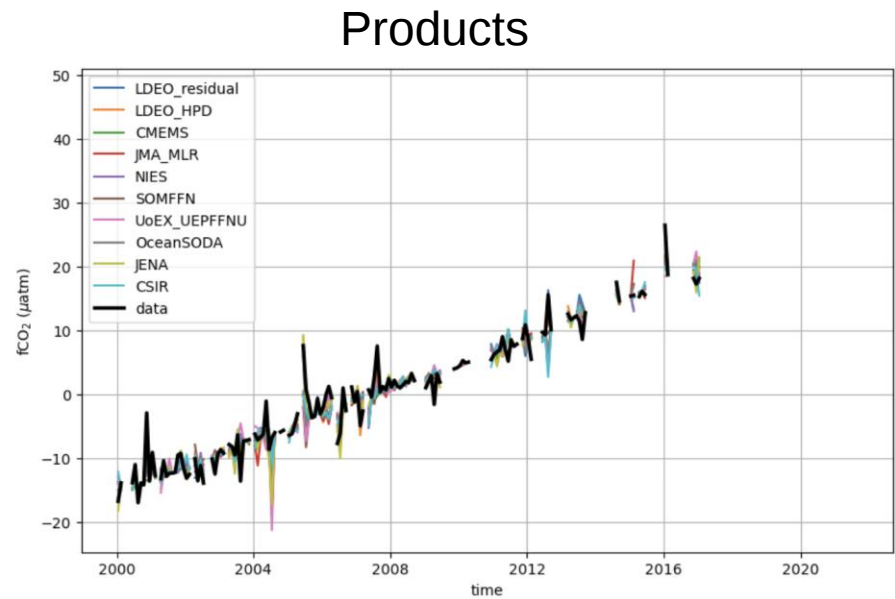
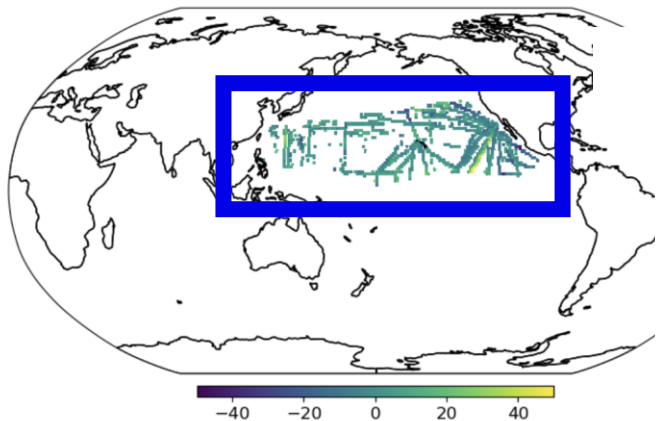


RMSE and correlation

	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,931)	12.2µatm r=0.89	
SOCAT + GLODAP (n=41,820)		17.7µatm r=0.85

McKinley et al. in prep

N. Pacific STPS: Remove climatology



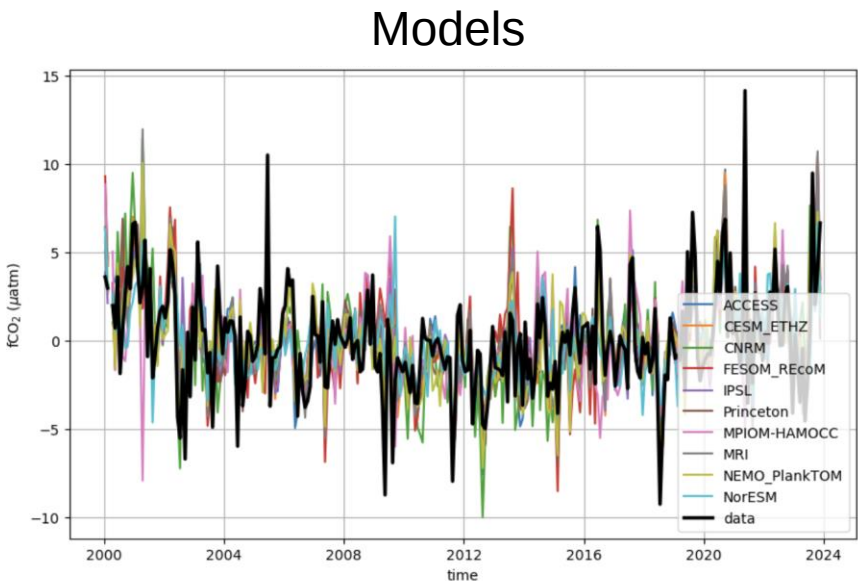
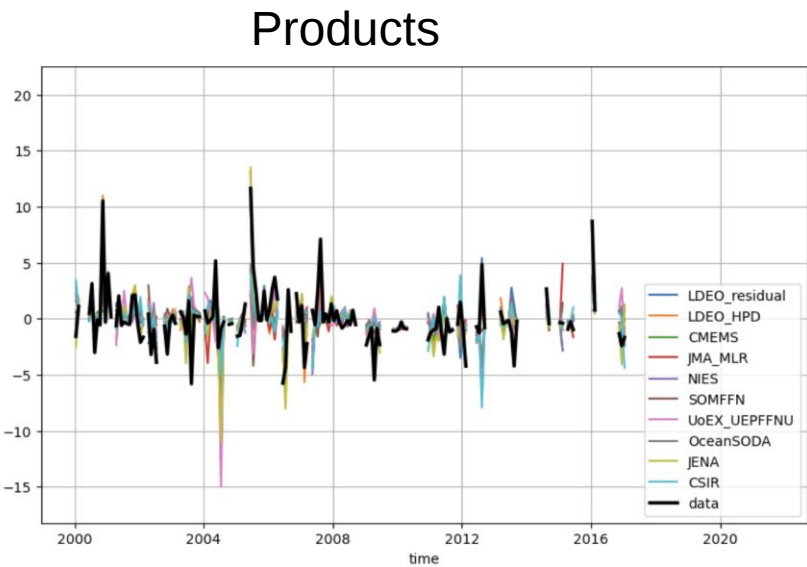
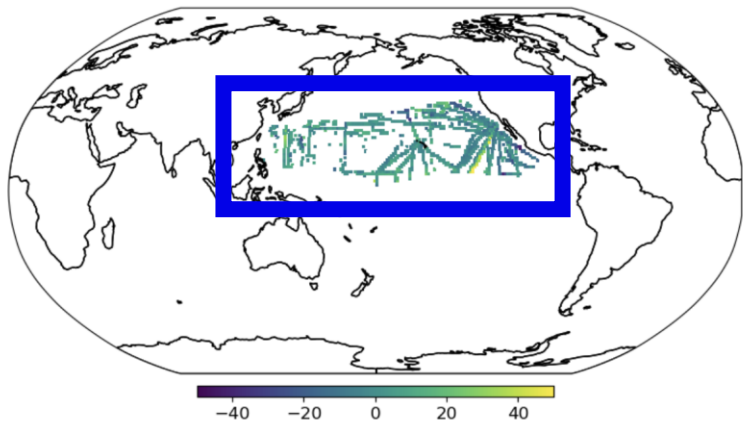
RMSE and correlation

2000-2023	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,931)	3.7μatm r=0.92	
SOCAT + GLODAP (n=41,820)		7.2μatm r=0.89

Product, Model, Datasets
Trend = 2.1-2.2 uatm/yr

McKinley et al. in prep

N. Pacific STPS: Remove trend



RMSE and correlation

2000-2023	10 Product mean	10 Model mean
LDEO + GLODAP (n=1,931)	3.7µatm r=0.40	
SOCAT + GLODAP (n=41,820)		7.2µatm r=0.47

McKinley et al. in prep