

Subsea cable observing system design with regional models

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1. SMART cables

2. Bottom pressure anomaly

3. Observing system assessment

Deep Ocean Observing



Figure 1: Proposed SMART sensor locations [Howe et al., 2019].

- Ocean observing system: *in situ* (buoys, moorings, ships) and remote (satellite)
- Deep ocean relatively unobserved/undersampled
- Proposed global network of scientific sensors "piggybacking" telecommunications cables on the ocean floor

SMART Cables



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Bottom pressure anomaly and wind

- Wind stress dominates seasonal bottom pressure changes relative to other atmospheric controls [Fukumori et al., 2015][Chen et al., 2023].
- During bottom pressure data assimilation, wind stress controls will receive the strongest adjustments
- In turn, constrain geostrophic transport

Consider the Qol of bottom pressure:

$$\mathcal{J}(t) = p_b$$

and it's monthly mean anomaly

$$\delta \mathcal{J}(t) = p_b - \overline{p_b}$$

Taylor series reconstruction using gradients w.r.t weekly forcing anomalies ϕ_i

$$\widetilde{\delta \mathcal{J}}(t) = \sum_{i} \sum_{\mathbf{x}} \sum_{\Delta t} \frac{\partial \mathcal{J}}{\partial \phi_i(\mathbf{x}, \Delta t)} \delta \phi_i(\mathbf{x}, t - \Delta t)$$

Q: Can we demonstrate the wind stress-bottom pressure connection at higher frequencies?

Consider the Qol of bottom pressure:

$$\mathcal{J}(t) = p_b$$

and it's monthly daily mean anomaly

$$\delta \mathcal{J}(t) = p_b - \overline{p_b}$$

Taylor series reconstruction using gradients w.r.t weekly hourly forcing anomalies ϕ_i

$$\widetilde{\delta \mathcal{J}}(t) = \sum_{i} \sum_{\mathbf{x}} \sum_{\Delta t} \frac{\partial \mathcal{J}}{\partial \phi_i(\mathbf{x}, \Delta t)} \delta \phi_i(\mathbf{x}, t - \Delta t)$$

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Observing System Simulated Experiments (OSSEs)

- Nature run
 - High resolution, high fidelity
 - Source of synthetic SMART data
 - Instrument, representation error
- Base/data assimilation model
 - Coarser, differing physics
 - Optimizable



Figure 2: Regional of interest encompassing islands of Vanuatu and New Caledonia.

Base model-nature run comparison

	base model	nature run
Horiontal grid spacing [degrees]	1/6	1/48
Vertical levels	50	90
Surface level thickness [meters]	10	1
Atmospheric forcing	6-hourly ECMWF analysis 0.14-degree grid bulk for- mulae/relative wind	6-hourly ERA-interim anal- ysis 0.7-degree grid bulk formulae
Atmospheric load	No	Yes
Tides	No	Yes
Barotropic time-stepping	Adams-Bashforth	Crank-Nicolson
Time step [seconds]	120s	25s

Table 1: Comparison of regional model to nature run [Gallmeier et al., 2023].

Synthetic data: vertical representation



Synthetic data: horizontal representation

Nature Run Field



Figure 4: Representing a nature run observation (single pixel) in a coarser model (grey squares) requires we carry some notion of representation uncertainty.

Observing System Simulated Experiments (OSSEs)



 $\mathbf{D}\mathbf{A}$ _ _ _ _ _ _



150°E 155°E 160°E 165°E 170°E 175°E

Example OSSE: Two temperature sensors



Example OSSE: Two temperature sensors



OSSE result: Skill score comparison

$$\mathsf{MSD}(x,y) = \frac{1}{n} \sum_{i=1}^{N} \left[(x_i - \langle x \rangle) - (y_i - \langle y \rangle) \right]^2$$

Normalized MSD relative to reference experiment is then

$$\mathsf{MSD}_{\mathsf{NORM}} = \frac{\mathsf{MSD}_{\mathsf{TRIAL}}}{\mathsf{MSD}_{\mathsf{REF}}}$$

and a skill score is assigned:

 $S = 1 - MSD_{NORM}$.

For our example, we compute MSD of SST

$$MSD_{NORM} = \frac{MSD(SST_{TRIAL}, SST_{NR})}{MSD(SST_{REF}, SST_{NR})}$$



Figure 6: Regional of interest encompassing islands of Vanuatu and New Caledonia.

Uncertainty Quantification

- Endow parameters with notion of (e.g. Gaussian) uncertainty
- Curvature of misfit cost function provides data-informed directions
- Project Qol sensitivities onto data-informed subspace
- Compute dynamic proxy potential, i.e. Qol uncertainty reduction in the face of new data [Loose and Heimbach, 2021]

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